The Math Behind the Weather

Ankur Desai, Dept of Atmospheric and Oceanic Sciences UW-Madison, 15 March 2018



Columbia Pictures http://flux.aos.wisc.edu desai@aos.wisc.edu @profdesai

Who Am I?

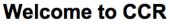
- I was born and raised in New Jersey
- I live in Madison with my wife and three daughters
- I am a climate scientist who has spent that past 2 decades studying how plants, climate, and weather all influence each other

THE CENTER FOR **CLIMATIC RESEARCH**

THE NELSON INSTITUTE FOR ENVIRONMENTAL STUDIES | UNIVERSITY OF WISCONSIN-MADISON ABOUT CCR NEWS RESEARCH RESOURCES SUPPORT CC

Member of the US LTER Network

Welcome to NTL-LTER





CCR researchers are investigating global and regional biogeochemistry, with a particular focus on the carbon cycle of the land biosphere a oceans and Great Lakes. Using data and elucidate natural carbon fluxes and the controlling them, and work to use this i improve predictive models.







North Temperate Lakes Long Term Ecological Rese

North Temperate Lak sites established by t and changing land us present, future).

Our primary study sit their surrounding lan Limnology at the Unit

Climate Impacts

- Land Surface Processes
- Oceanography and Limnology
- Past Climates

Department of Atmospheric and **Oceanic Sciences** ROALINI S



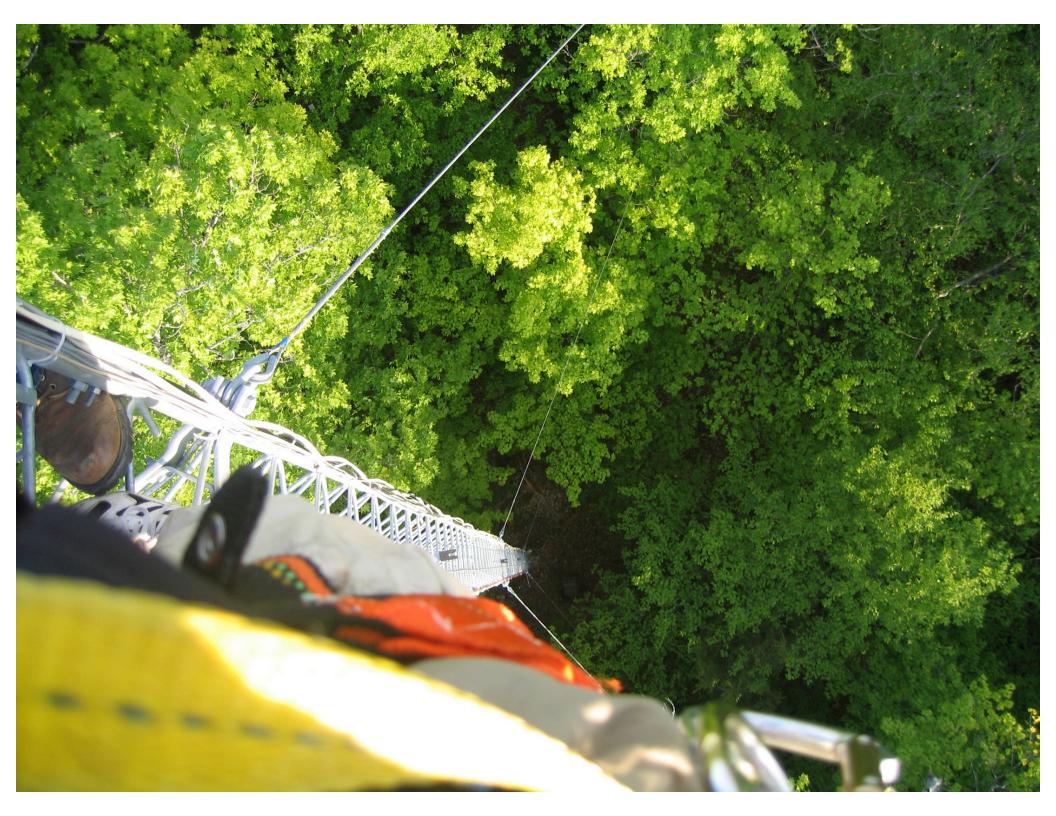
Since 1948 we have grown into one of the leading departments in our field of Atmospheric and Oceanic Sciences. We have strong graduate and undergraduate programs which are nationally recognized. We graduate about 15 Ph.D. and M.S. students each year; our graduates are active in research labs and universities around the world. We graduate approximately 20 B.S. students each year; they choose options allowing a focus on weather systems or general atmospheric science.

Our faculty of 15 has long maintained breadth and special strength in three areas:

- Climate systems, including the ocean
- Satellite and remote sensing
- Weather systems, including synoptic-dynamic meteorology

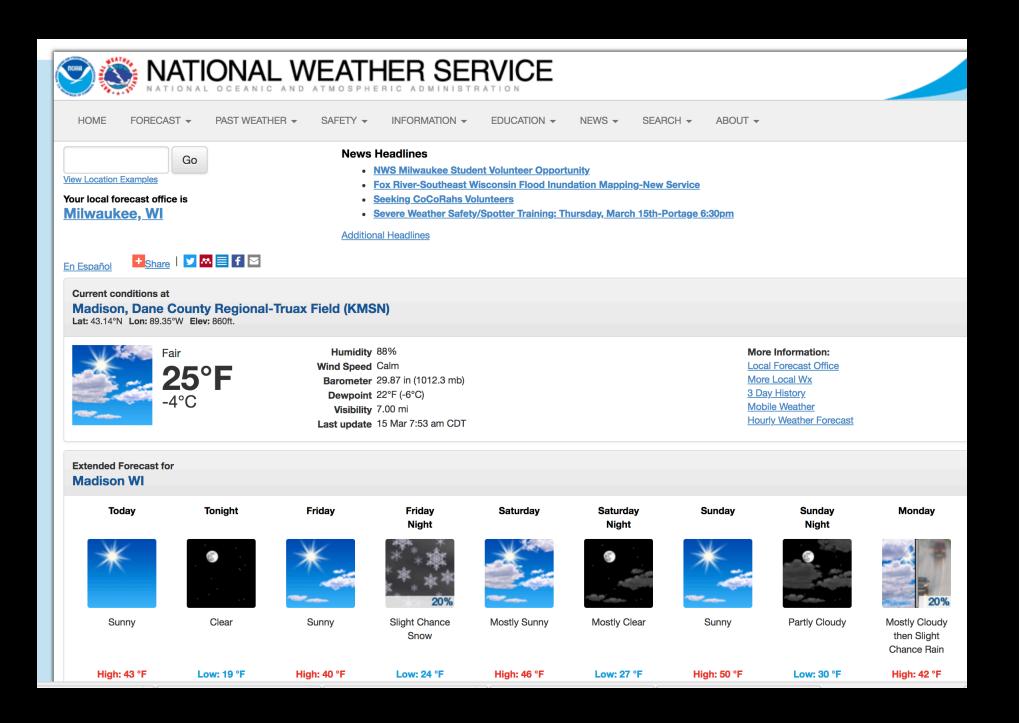


Who We Are









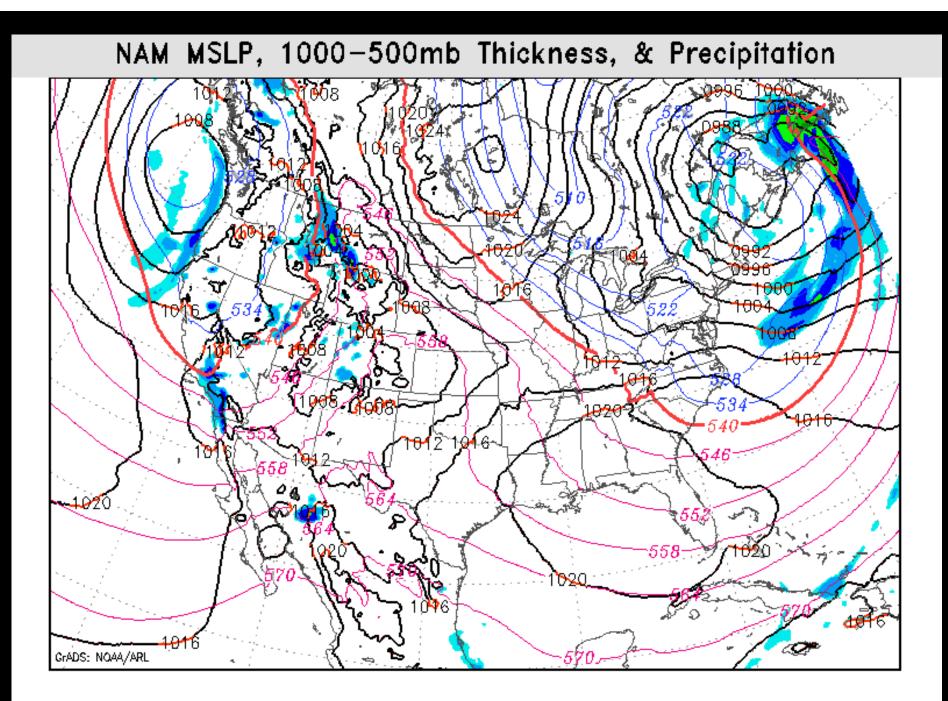
Friday and Saturday...Forecast confidence is medium.

A **shortwave trough** will lose **amplitude** while progressing from the central High Plains to the Ohio River valley. Rapid cyclogenesis will commence along the Colorado Front Range on Friday, with the occluding surface cyclone reaching the mid Mississippi and Ohio Valleys by late Saturday. Forecast models are in good agreement with keeping the low, and its associated tight **baroclinic zone**, to our south. However, there is a window Friday night where column moisture increases (Precipitable Water values ~ 0.5" in southwest Wisconsin), juxtaposed with robust lift and modest moisture within the dendritic (snow) growth zone. The GFS ensemble members are now in better agreement with the surface low track, with most solutions bringing measurable precipitation into southwest Wisconsin on Friday night. This compares favorably with the latest **ECMWF** and Canadian deterministic runs, which suggest the same. The **<u>NAM</u>** appears a rather wet outlier (with > 0.25" of **QPF** southwest of Madison) and is discounted for now. All of this is to say that we're seeing enough model consistency to increase rain/snow chances on Friday night, particularly for areas southwest of Madison and towards the Illinois border where light snow accumulations are possible.

Sunday through Thursday...Forecast confidence is medium.

<u>Shortwave</u> troughing ejects out of the Four Corers region on Sunday, before reaching the lower Missouri and Ohio Valleys by Monday night. The associated surface low passes through the mid Mississippi

https://forecast.weather.gov/product.php?site=MKX&issuedby=MKX&product=A FD&format=CI&version=1&glossary=1

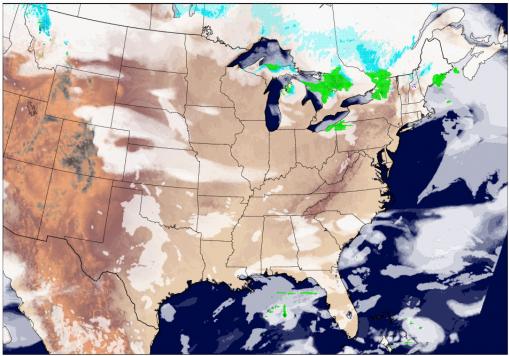


Analysis: 12Z Thu 15 MAR, 2018

SLP (mb-1000), 1000-500mb Thk (dam), 3hr Precip.(mm)

North American Mesoscale Model

Sunday 0900 EDT, November 10, 2013

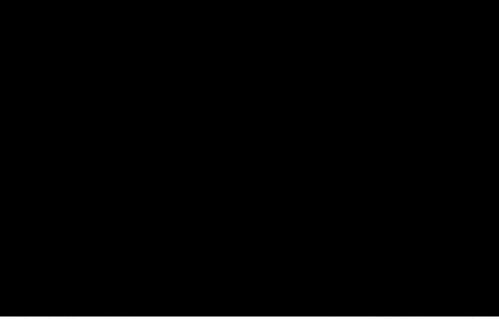


Mixed

Snow

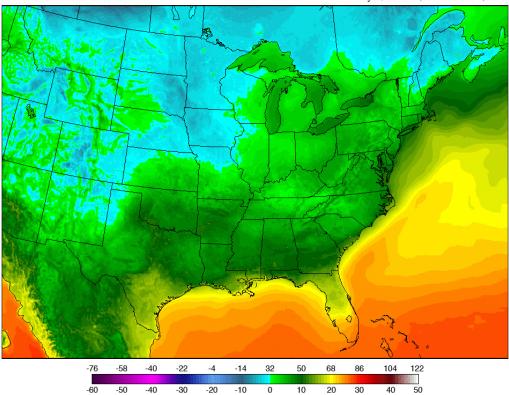
Rain

HThe Climate Reanalyzer | cci-reanalyzer.org



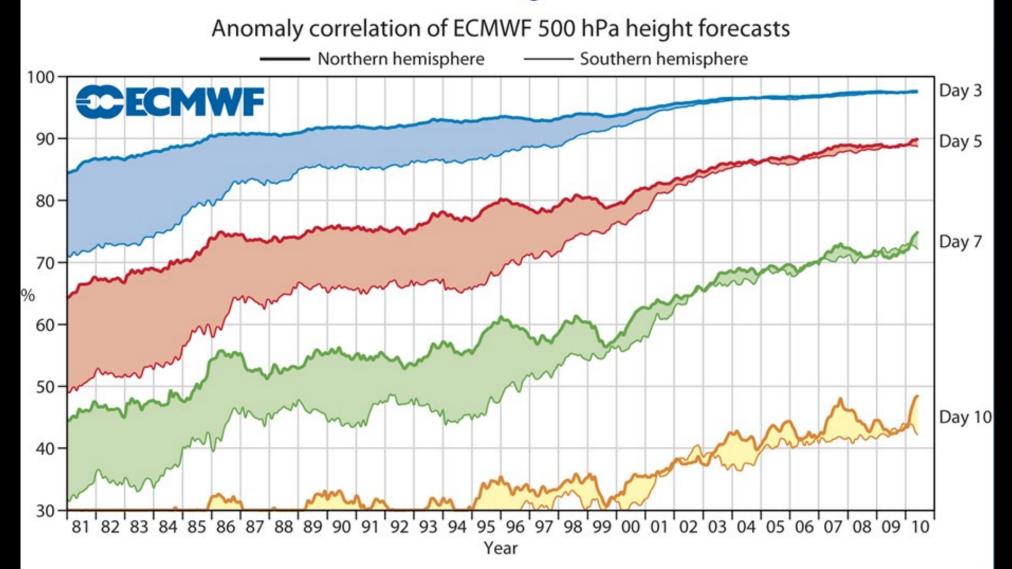
North American Mesoscale Model

Sunday 0900 EDT, November 10, 2013



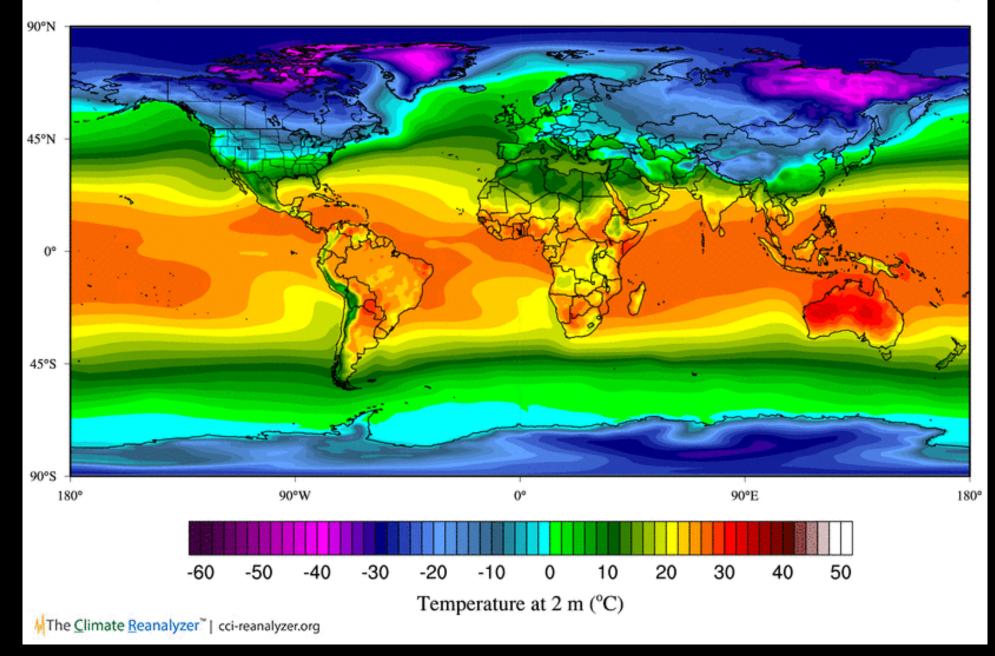
Temperature at 2 meters (°F/°C)

Advances in Global and Regional Weather Forecasts



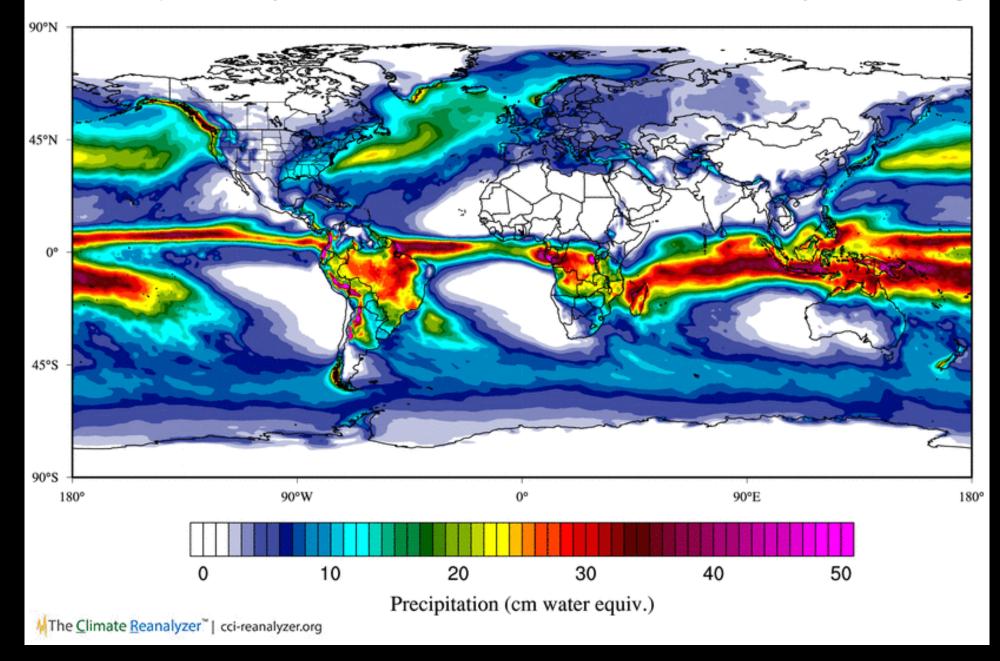
ERA-Interim | Climate Reanalyzer

January 15 1979-2000 Average

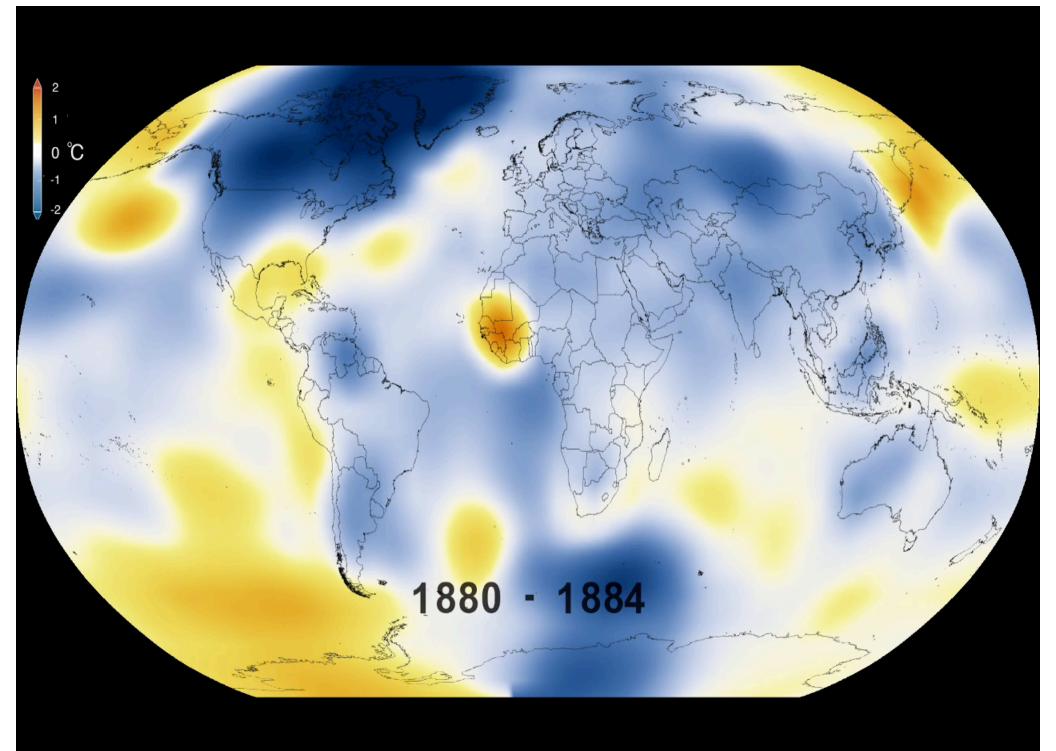


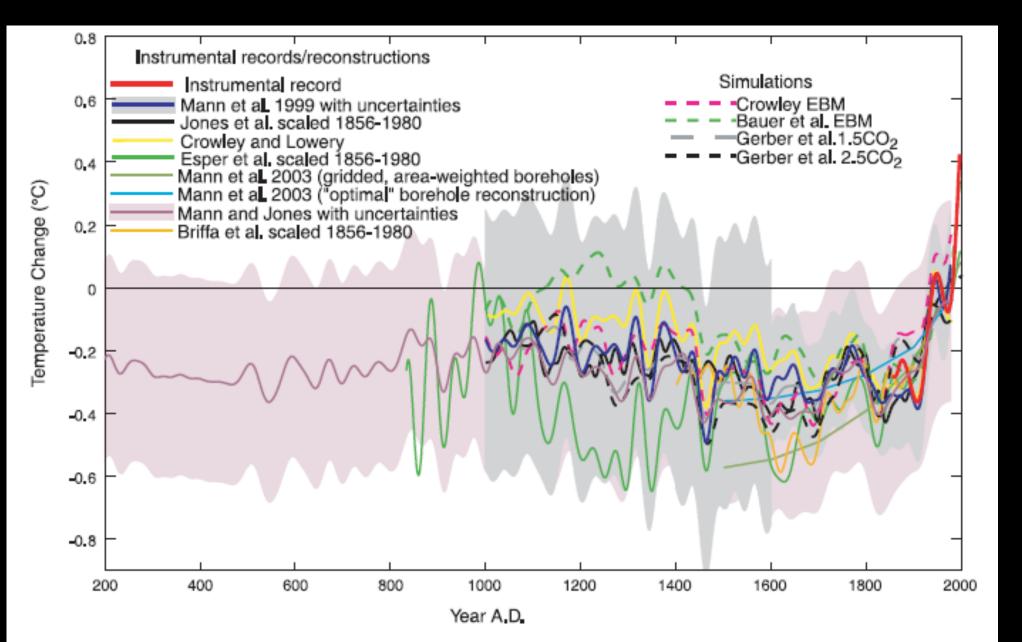
http://cci-reanalyzer.org/

ERA-Interim | Climate Reanalyzer



http://cci-reanalyzer.org/





Mann et al., 2003, EOS

Climate Model = Weather Model

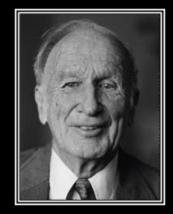
Major math innovations related to meteorology

- Navier-Stokes equation in turbulent, rotating reference frame
- Chaos in Non-linear dynamical systems
- Numerical solutions and computational approaches to non-linear PDEs
- Analytical geometry
- Statistical Bayesian data assimilation of Earth-atmosphere-ocean observations

Predictability in a deterministic nonperiodic flow

"Does the flap of a butterfly's wings in Brazil set off a tornado in Texas?" -(Lorenz 1972)





Deterministic Nonperiodic Flow¹

Edward N. Lorenz

Massachusetts Institute of Technology

(Manuscript received 18 November 1962, in revised form 7 January 1963)

Abstract

Finite systems of deterministic ordinary nonlinear differential equations may be designed to represent forced dissipative hydrodynamic flow. Solutions of these equations can be identified with trajectories in phase space. For those systems with bounded solutions, it is found that nonperiodic solutions are ordinarily unstable with respect to small modifications, so that slightly differing initial states can evolve into considerably different states. Systems with bounded solutions are shown to possess bounded numerical solutions. A simple system representing cellular convection is solved numerically. All of the solutions are found

to be unstable, and almost all of them are nonperiodic.

The feasibility of very-long-range weather prediction is examined in the light of these results.

Sensitive dependence to initial conditions

"Finite time for error in representation of small scales to affect accuracy of simulation of large scales, no matter how small in scale and hence amplitude this model error is"

 $\frac{dx}{dt} = \sigma(y-x)$

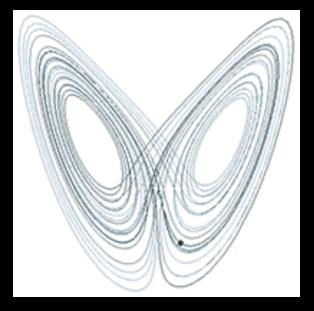
$$\frac{dy}{dt} = rx - y - xz$$

$$\frac{dz}{dt} = xy - bz$$

$$r = 28, \sigma = 10, and b = 8/3$$

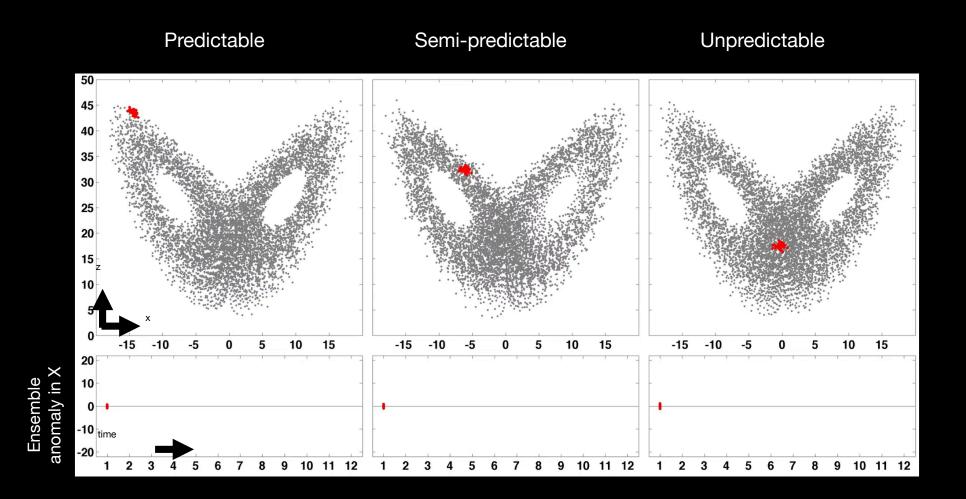
-(Lorenz 1969)

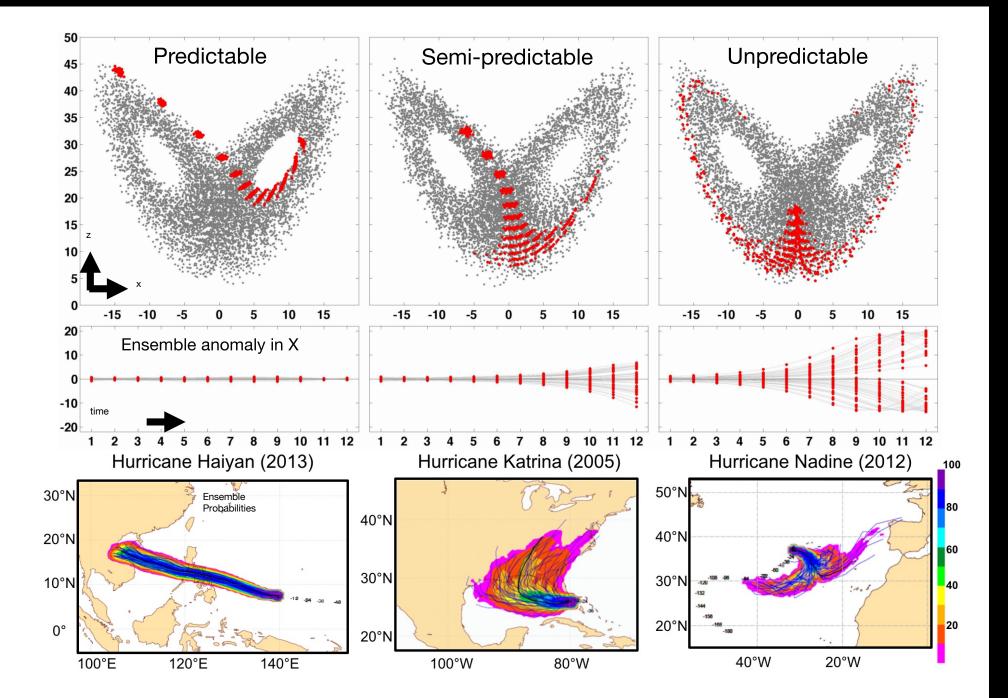




source: wikipedia

Ensemble Forecast with Initial Uncertainty





Hotter

What's Really Warming the World?

Skeptics of manmade climate change offer various natural causes to explain why the Earth has warmed 1.4 degrees Fahrenheit since 1880. But can these account for the planet's rising temperature? Watch to see how much different factors, both natural and industrial, contribute to global warming, based on findings from NASA's Goddard Institute for Space Studies.



Based on an interactive by Bloomberg



Colder

Transition to Turbulence



Album of Fluid Motion (Van Dyke)

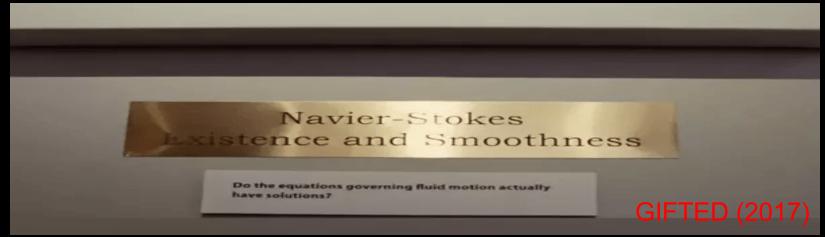


MODIS Jan 2017

Navier-Stokes a.k.a Newton's Second Law for a "Newtonian" Fluid

$$rac{\partial \mathbf{v}}{\partial t} + (\mathbf{v} \cdot
abla) \mathbf{v} = -
abla p +
u \Delta \mathbf{v} + \mathbf{f}(oldsymbol{x},t)$$

N-S Smoothness and Existence Problem is a Millennial Math Problem



1. $\mathbf{v}(x,t) \in \left[C^{\infty}(\mathbb{R}^3 \times [0,\infty))\right]^3$, $p(x,t) \in C^{\infty}(\mathbb{R}^3 \times [0,\infty))$ 2. There exists a constant $E \in (0,\infty)$ such that $\int_{\mathbb{R}^3} |\mathbf{v}(x,t)|^2 dx < E$ for all $t \ge 0$.

(A) Existence and smoothness of the Navier–Stokes solutions in \mathbb{R}^3

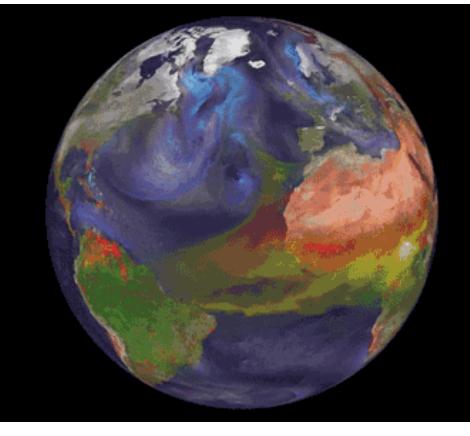
Let $\mathbf{f}(x,t) \equiv 0$. For any initial condition $\mathbf{v}_0(x)$ satisfying the above hypotheses there exist smooth and globally defined solutions to the Navier–Stokes equations, i.e. there is a velocity vector $\mathbf{v}(x,t)$ and a pressure p(x,t) satisfying conditions 1 and 2 above.

(B) Breakdown of the Navier–Stokes solutions in \mathbb{R}^3

There exists an initial condition $\mathbf{v}_0(x)$ and an external force $\mathbf{f}(x,t)$ such that there exists no solutions $\mathbf{v}(x,t)$ and p(x,t) satisfying conditions 1 and 2 above.

Going onto a rotating reference frame

$$ho rac{D \mathbf{u}}{D t} = -
abla ar{p} + \mu \,
abla^2 \mathbf{u} + rac{1}{3} \mu \,
abla (
abla \cdot \mathbf{u}) +
ho \mathbf{g} -
ho \left(2 \mathbf{\Omega} imes \mathbf{u} + \mathbf{\Omega} imes \mathbf{\Omega} imes \mathbf{x} + rac{d \mathbf{U}}{d t} + rac{d \mathbf{\Omega}}{d t} imes \mathbf{x}
ight).$$



N-S For Earth System

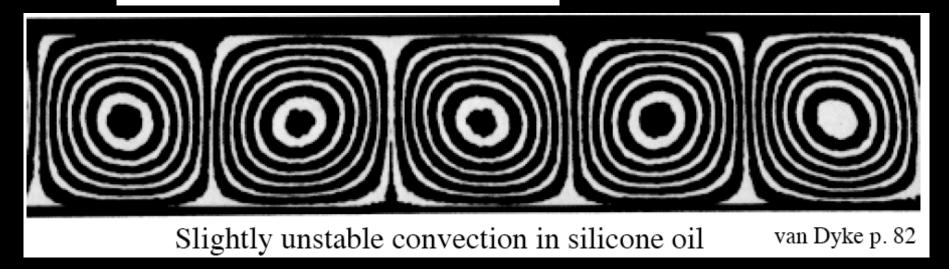
$$\begin{split} \frac{\partial \rho}{\partial t} &= -\frac{\partial \rho u_j}{\partial x_j}, \\ \frac{\partial \theta}{\partial t} &= -u_j \frac{\partial \theta}{\partial x_j} + S_{\theta}, \\ \frac{\partial u_i}{\partial t} &= -u_j \frac{\partial u_i}{\partial x_j} - \frac{1}{\rho} \frac{\partial p}{\partial x_i} - g \delta_{i3} - 2\epsilon_{ijk} \Omega_j u_k, \\ \frac{\partial q_n}{\partial t} &= -u_j \frac{\partial q_n}{\partial x_j} + S_{q_n}, \qquad n = 1, 2, 3 \end{split}$$

One Solution to N-S: Convective instability

- Requires density to increase with height
- Instability occurs when Rayleigh number reaches critical threshold

$$Ra = h^3 \Delta B / \nu \kappa > 1700$$

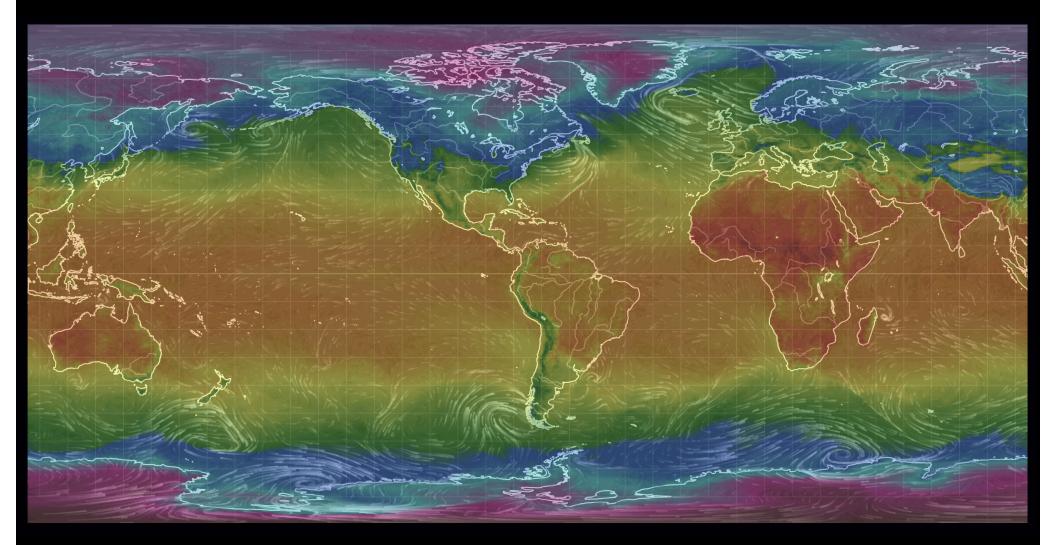
$$\Delta B = g \Delta T / T$$



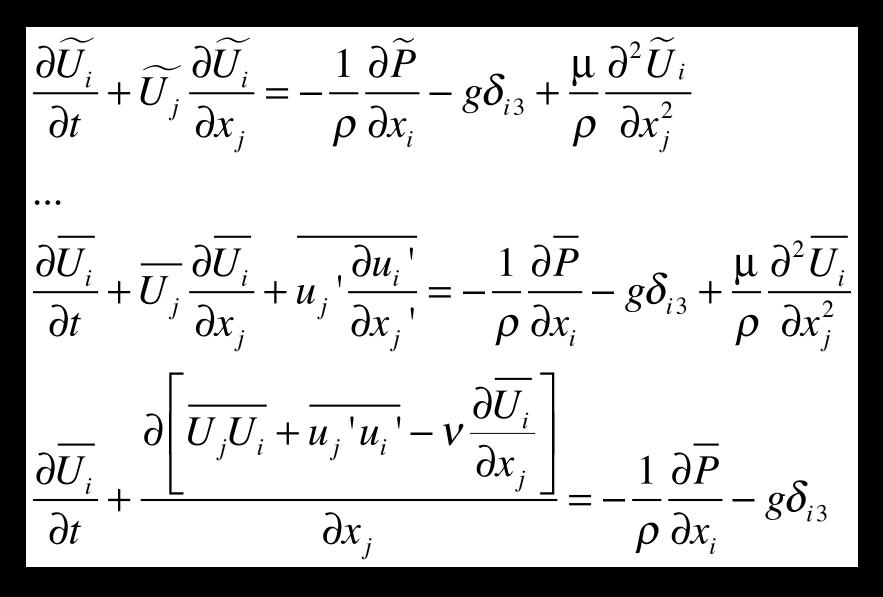
Cloud Streets: Convective Instability in the Real World



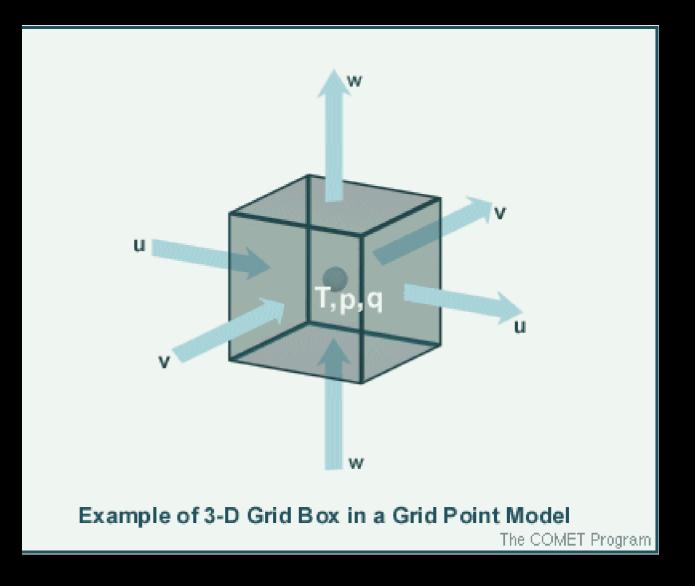
https://earth.nullschool.net/#current/wind/surface/level/overlay=temp/equirectangular



Turbulent Momentum Equation



Numerical Solutions are Necessary



http://www

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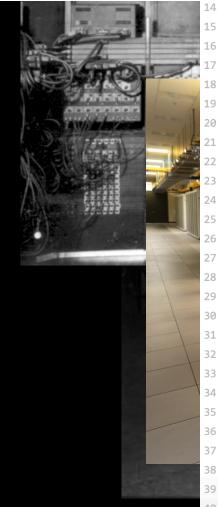
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Titan - Cray XK7, (9 4 NVIDIA K20x , Cray¹⁰ DOE/SC/Oak Ridge₁₂ United States



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SUBROUTINE	ADVE(NTSD,DT,DETA1,DETA2,PDTOP	&
&	,CURV,F,FAD,F4D,EM_LOC,EMT_LOC,EN,ENT,DX,DY	&
&	,HTM,HBM2,VTM,VBM2,LMH,LMV	&
&	,T,U,V,PDSLO,TOLD,UOLD,VOLD	&
&	,PETDT,UPSTRM	&
&	,FEW,FNS,FNE,FSE	&
&	,ADT,ADU,ADV	&
&	,N_IUP_H,N_IUP_V	&
&	,N_IUP_ADH,N_IUP_ADV	&
&	,IUP_H,IUP_V,IUP_ADH,IUP_ADV	&
&	,IHE,IHW,IVE,IVW,INDX3_WRK	&
&	,IDS,IDE,JDS,JDE,KDS,KDE	&
&	,IMS,IME,JMS,JME,KMS,KME	&
&	,ITS,ITE,JTS,JTE,KTS,KTE)	
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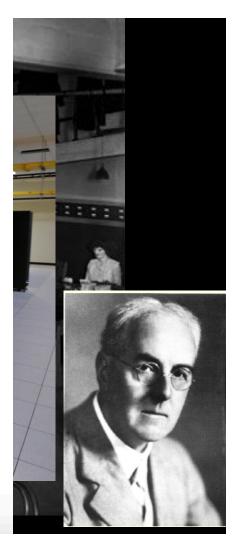
!\$\$\$ SUBPROGRAM DOCUMENTATION BLOCK

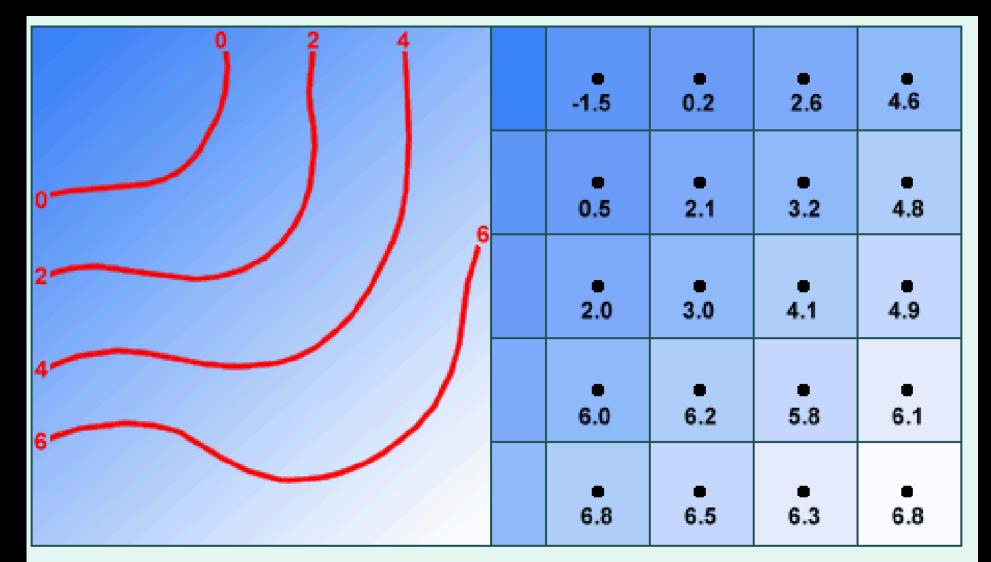
! SUBPROGRAM: ADVE	HORIZONTAL AND VERTICAL ADVECTION
PRGRMMR: JANJIC	ORG: W/NP22 DATE: 93-10-28

ABSTRACT:

22	!	ABSTRACT:			
23	!	ADVE CA	LCULATES THE	(CONTRIBUTION OF THE HORIZONTAL AND VERTICAL
24	!	ADVECTI	ON TO THE TE	N	DENCIES OF TEMPERATURE AND WIND AND THEN
25	!	UPDATES	THOSE VARIA	۱B	LES.
26	!	THE JAN	JIC ADVECTIO	DN	SCHEME FOR THE ARAKAWA E GRID IS USED
27	!	FOR ALL	VARIABLES 1	IN:	SIDE THE FIFTH ROW. AN UPSTREAM SCHEME
28	!	IS USED	ON ALL VAR	A	BLES IN THE THIRD, FOURTH, AND FIFTH
29	!	OUTERMO	ST ROWS. TH	łE	ADAMS-BASHFORTH TIME SCHEME IS USED.
30	!				
31	!	PROGRAM HIS	TORY LOG:		
32	!	87-06-??	JANJIC	-	ORIGINATOR
33	!	95-03-25	BLACK	-	CONVERSION FROM 1-D TO 2-D IN HORIZONTAL
34	!	96-03-28	BLACK	-	ADDED EXTERNAL EDGE
35	!	98-10-30	BLACK	-	MODIFIED FOR DISTRIBUTED MEMORY
36	!	99-07-	JANJIC	-	CONVERTED TO ADAMS-BASHFORTH SCHEME
37	!				COMBINING HORIZONTAL AND VERTICAL ADVECTION
38	!	02-02-04	BLACK	-	ADDED VERTICAL CFL CHECK
39	!	02-02-05	BLACK	-	CONVERTED TO WRF FORMAT
40	!	02-08-29	MICHALAKES	-	CONDITIONAL COMPILATION OF MPI

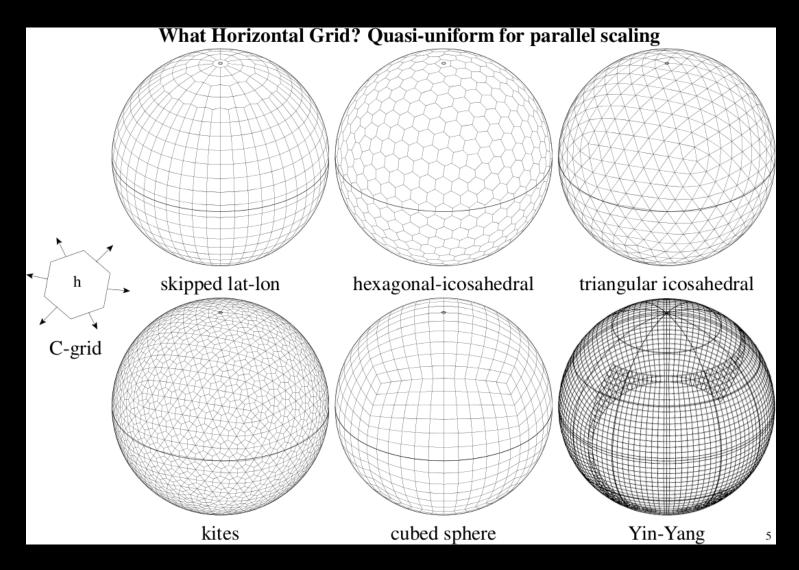




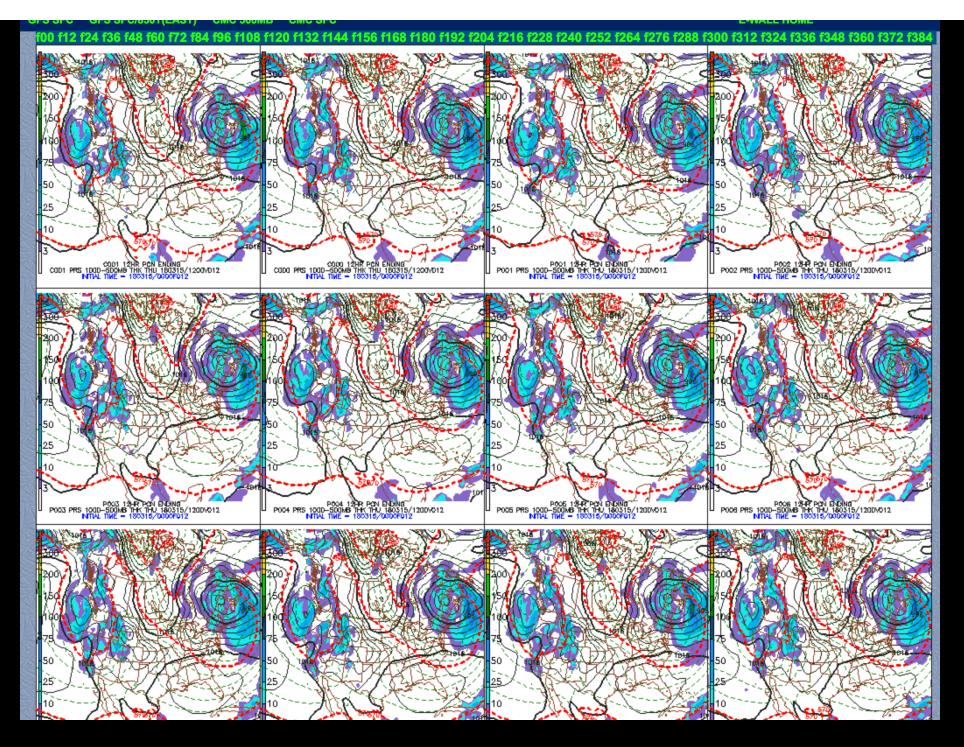


Actual smooth and continuous temperature field in degrees C (similar to spectral model representation) Grid point model representation of the same temperature field in degrees C

But the Earth is a sphere

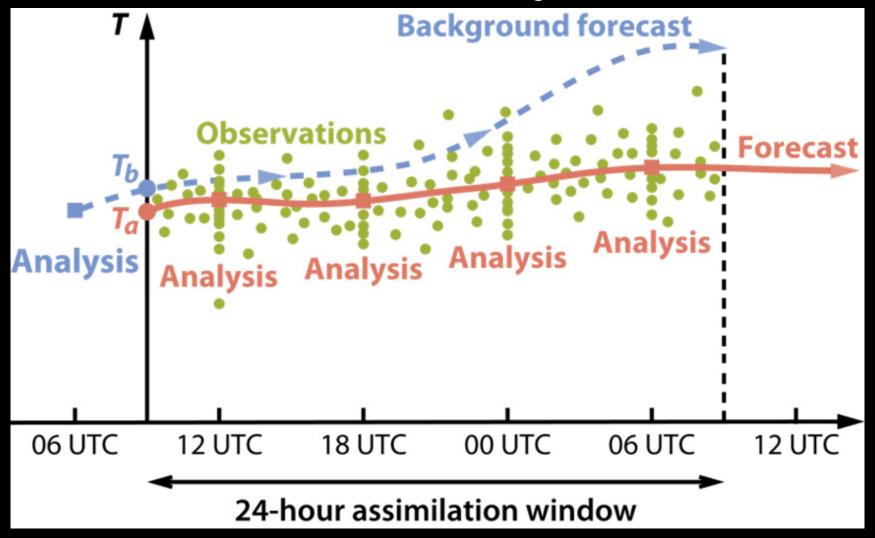


https://blogs.reading.ac.uk/wcd/2012/02/07/gungho-development-of-a-newdynamical-core-for-the-unified-model/



http://mp1.met.psu.edu/~fxg1/ENSPRS_12z/ensloop.html

Model spread needs to be constrained by data



Applications of Bayes' Rule

Likelihood

How probable is the evidence given that our hypothesis is true?

Prior

How probable was our hypothesis before observing the evidence?

$$P(H \mid e) = \frac{P(e \mid H) P(H)}{P(e)}$$

Posterior

How probable is our hypothesis given the observed evidence? (Not directly computable)

Marginal

How probable is the new evidence under all possible hypotheses? $P(e) = \sum P(e \mid H_i) P(H_i)$

Ensemble Kalman Filter

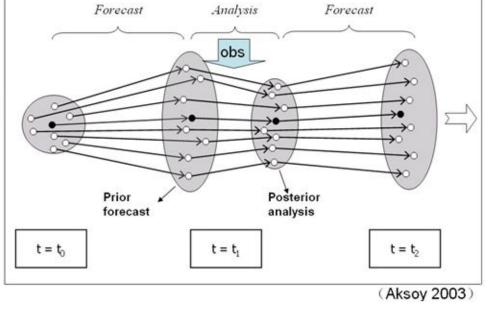
Ensemble Kalman filter is a Mont Carlo approximation of Kalman Filter. It samples the probability density function (PDF) of forecast and analysis using ensemble. (Evenson 1994).

 $X^a = X^f + K(y^o - HX^f)$

Before Ensemble Kalman filter

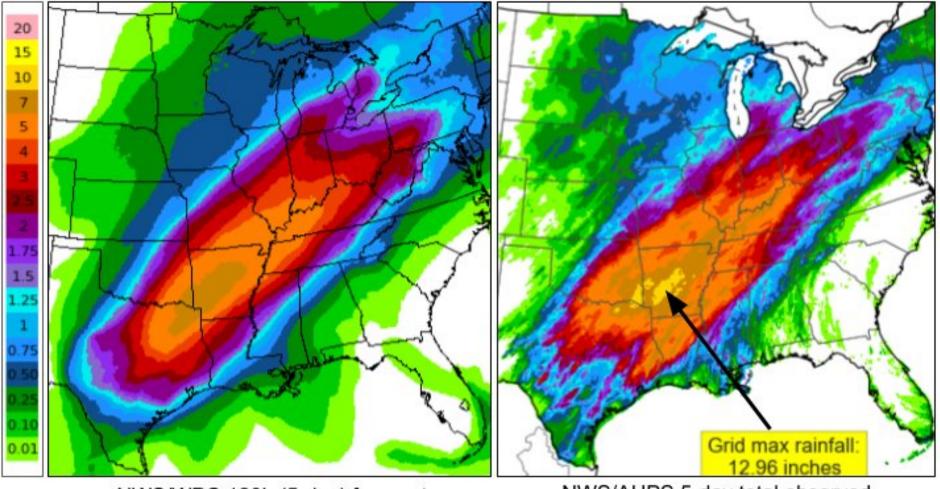
$$P^f \approx P_e^f = \langle x_i^{f'}, x_i^{f'} \rangle$$

After Ensemble Kalman filter $P^a \approx P_e^a = \langle x_i^{a'}, x_i^{a'} \rangle$

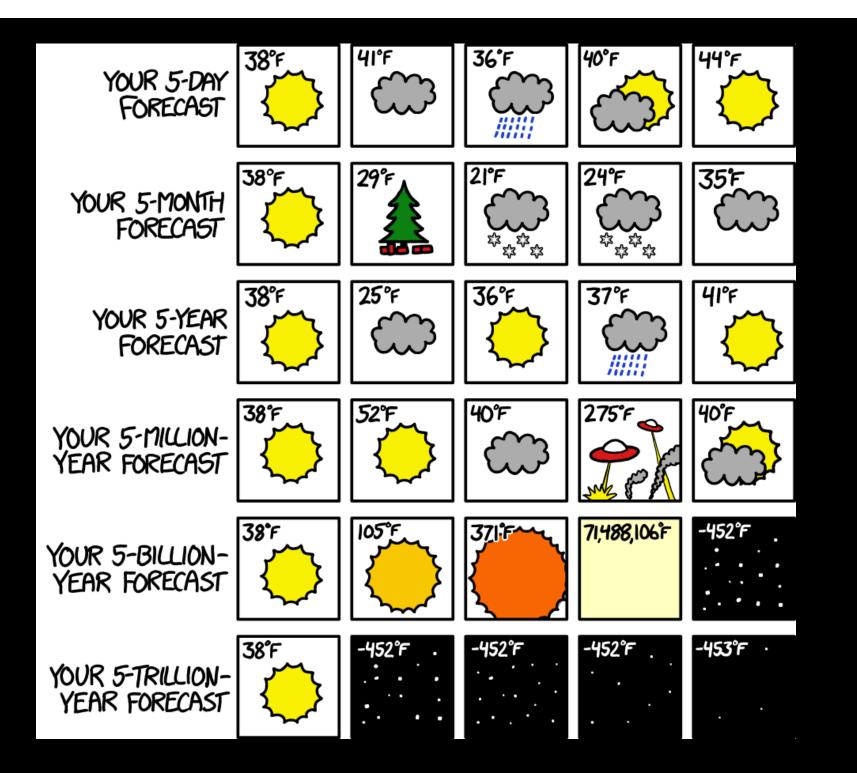


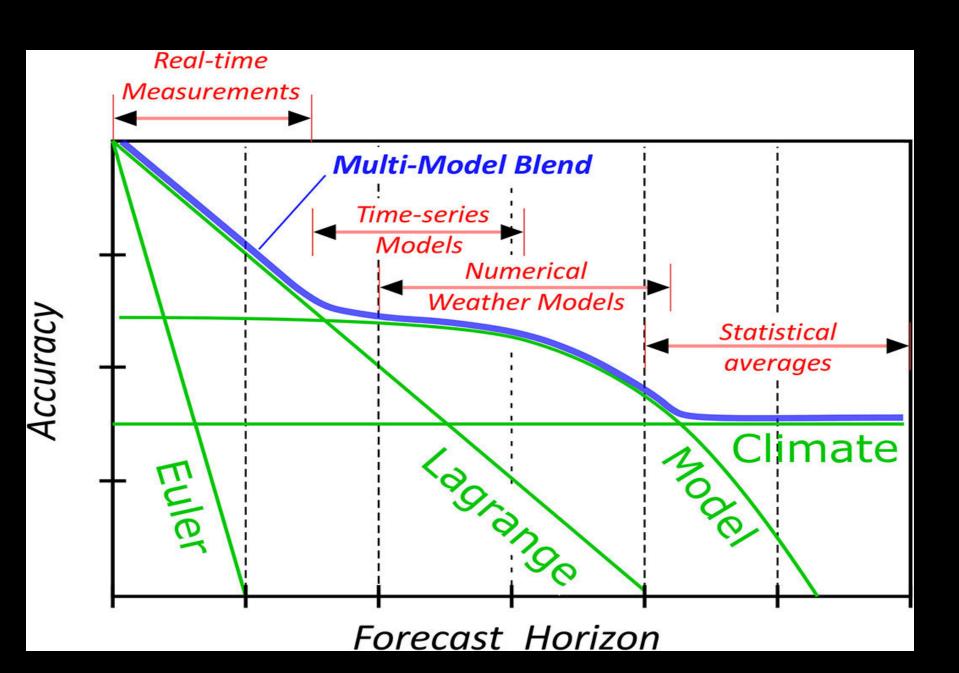
http://slideplayer.com/slide/9781007/ Stochastic filter and Deterministic filter

WPC 120h (5-day) Precipitation Forecast (left) vs. Observed Precipitation (NWS/AHPS, right)

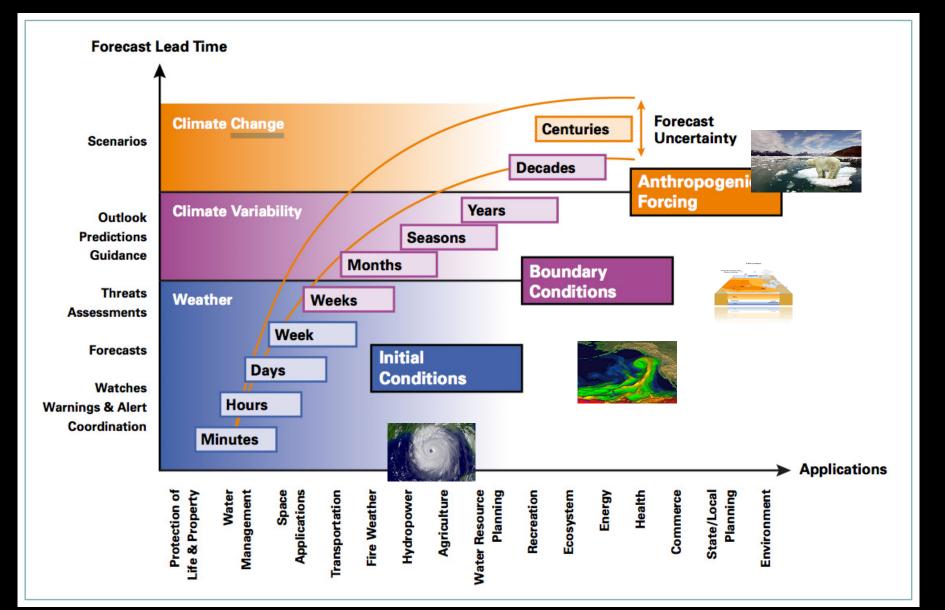


NWS/WPC 120h (5-day) forecast issued 0852 UTC Tue 2/20/18 NWS/AHPS 5-day total observed precipitation ending 12 UTC 2/25/18





http://spie.org/Images/Graphics/Newsroom/Imported-2015/006142/006142_10_fig1.jpg



source: https://www.wmo.int/

THANKS!

EDELRID

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Funding: NSF, DOE Ameriflux, NASA APRIL Group

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