

How scale-dependent is surface-atmosphere exchange?

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University of Wisconsin-Madison

14 Apr 2015, Chaos and Complex Systems



Why is this so damn hard to model?



What does it have to do with scale?

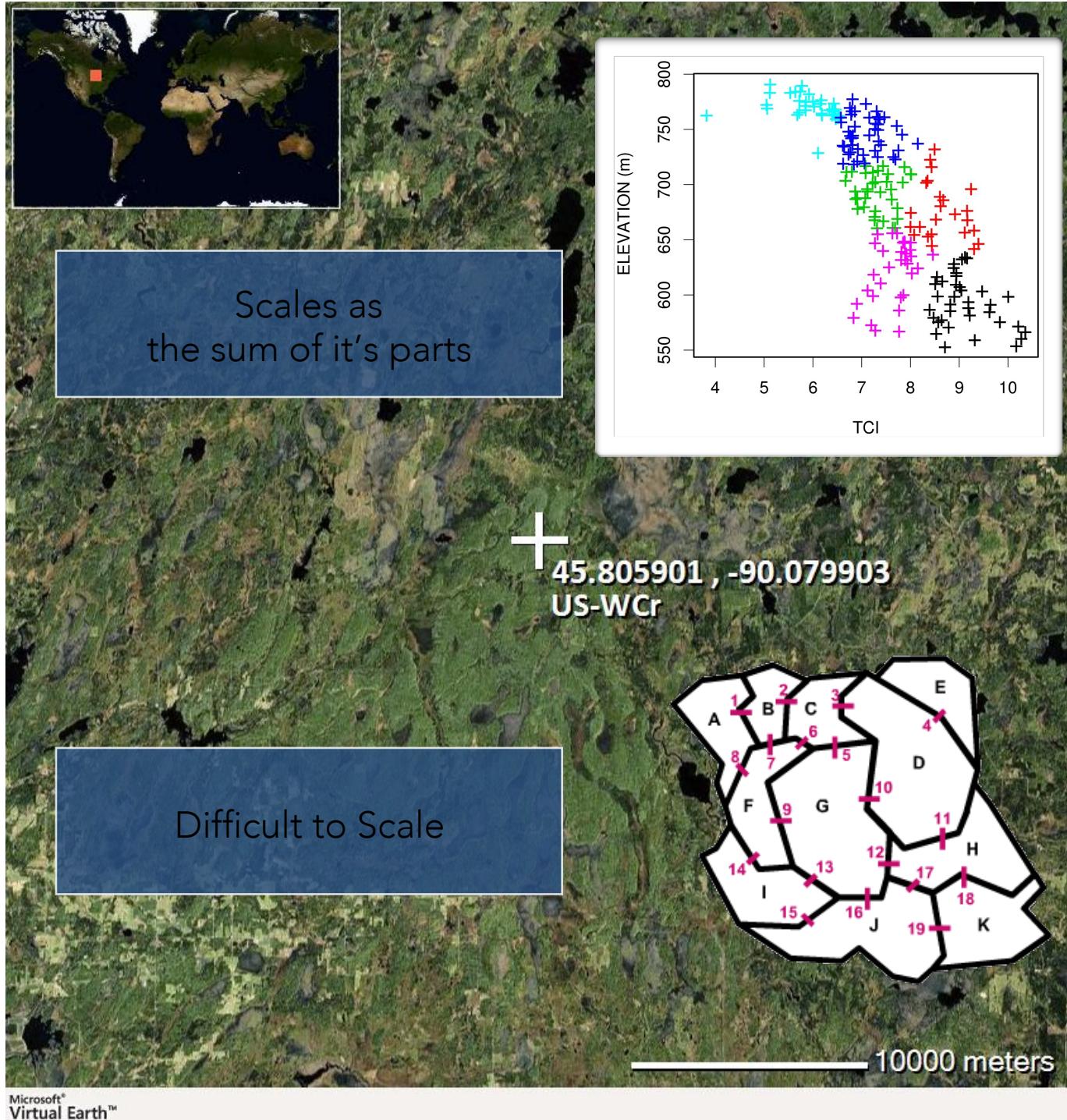
We face a fundamental scale mismatch



Between observations &
models

Between the atmosphere &
ecosystems



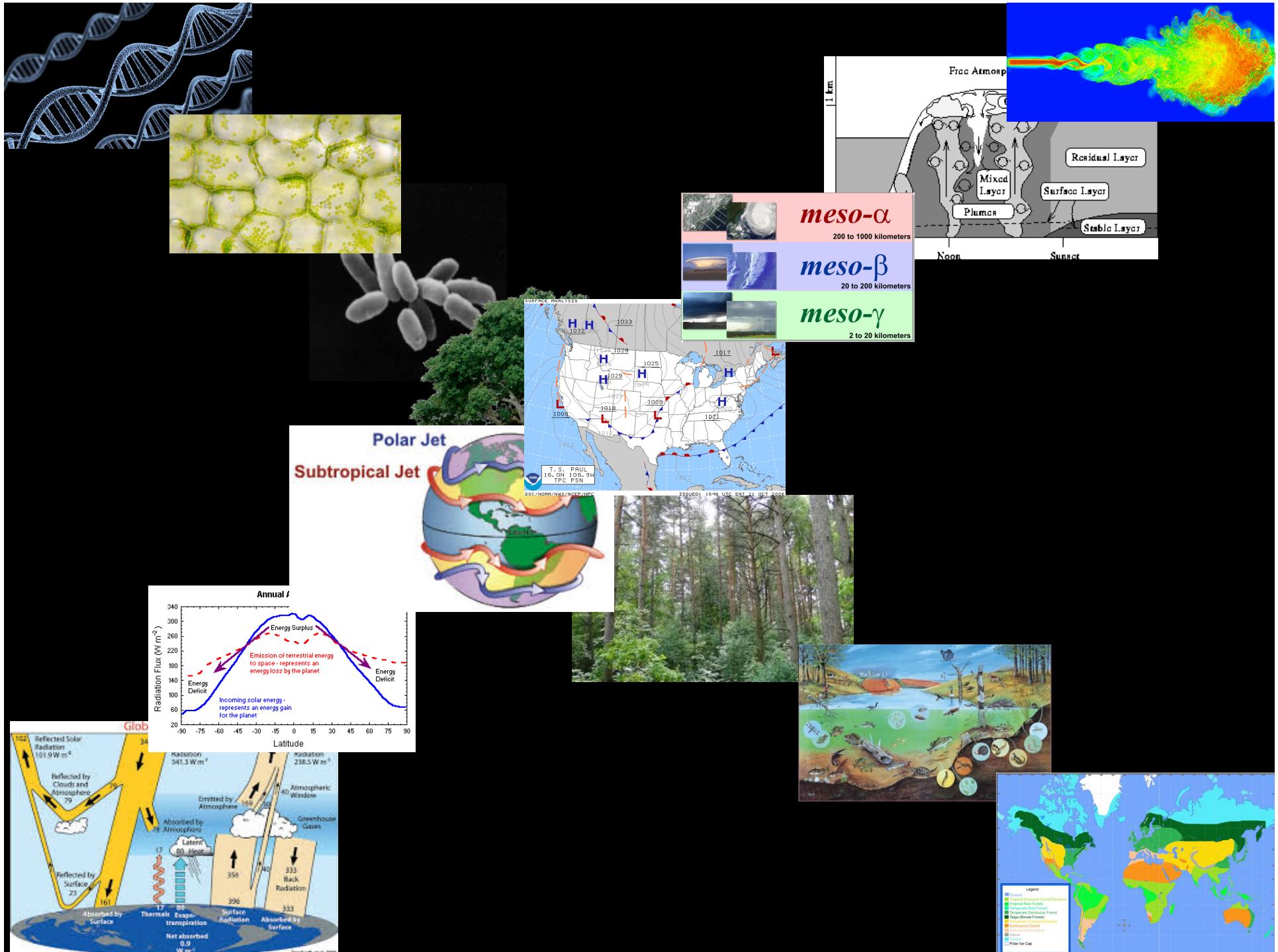


Spatial Heterogeneity

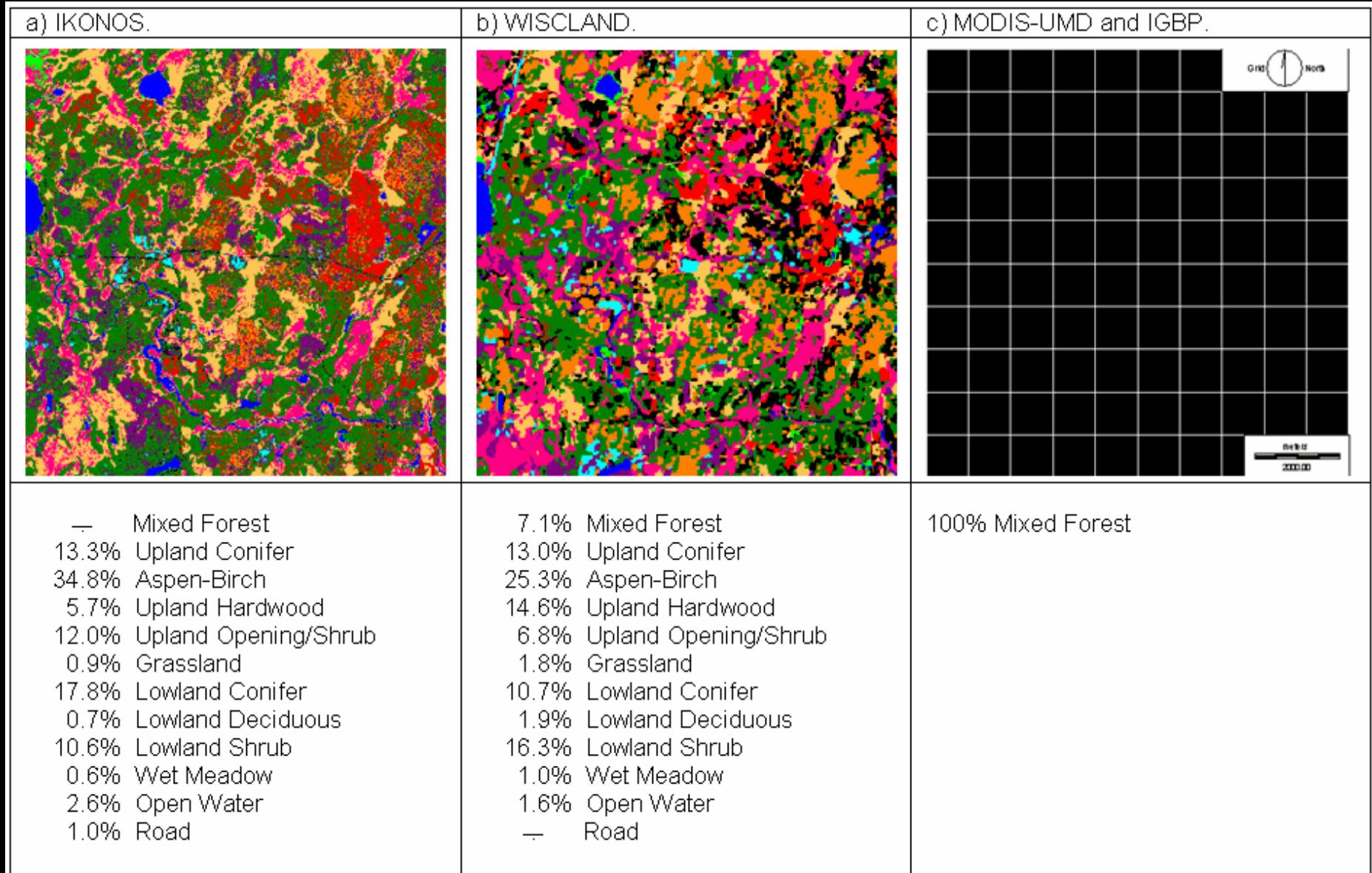
- Amount
- Frequency
- Distribution

Spatial Process

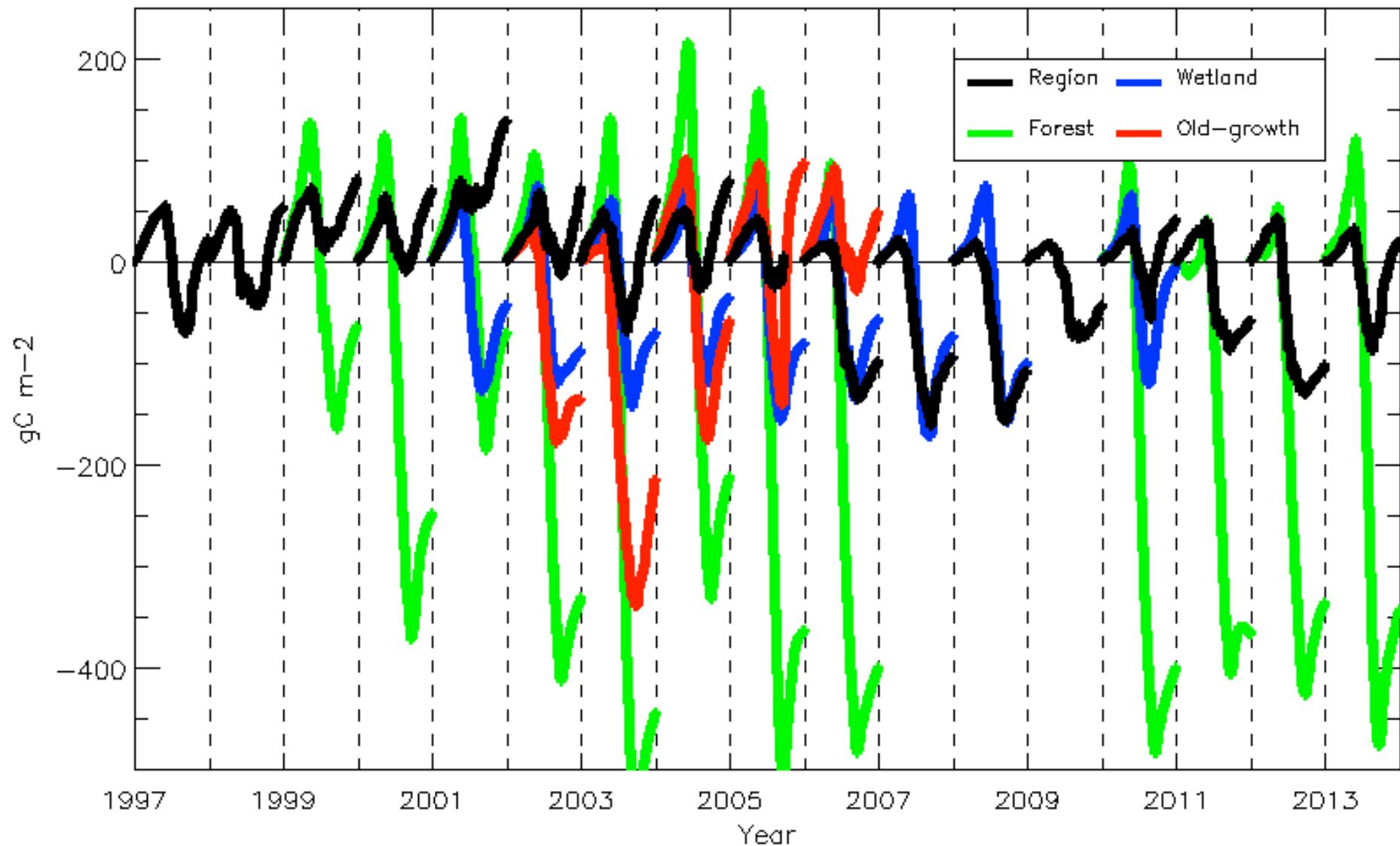
- Arrangement
- Location
- Distance



Complex Regions: 1+1≠2



Cumulative NEE



Global NPP 1983 version

FUNG ET AL.: BERN CO₂ SYMPOSIUM

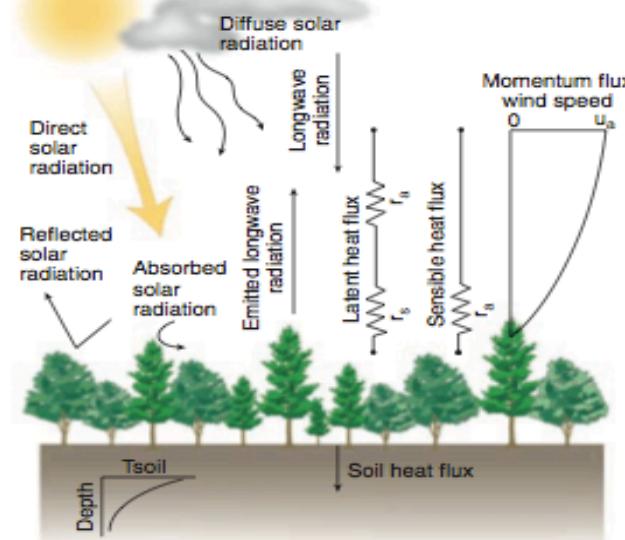
1285

	-180		-150		-120		-90		-60		-30		0		30		60		90		120		150																	
LAT	J	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36			
90.0	24																																							
82.2	23																																							
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66.5	21	4	1	14	23	18	16	14	7	6	5	2	5	0	2	0	0	2	0	0	5	17	24	14	18	19	17	20	24	23	27	28	22	18	22	19	12	0		
58.7	20	0	0	9	8	7	18	25	20	17	4	2	8	2	0	0	1	1	5	17	14	27	28	28	29	25	20	28	29	29	30	30	17	8	9	3	0			
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-74.3	3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	
-82.2	2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	
-90.0	1	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	
LAT	J	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36			
		-180	-150	-120	-90	-60	-30	0	30	60	90	120	150																											

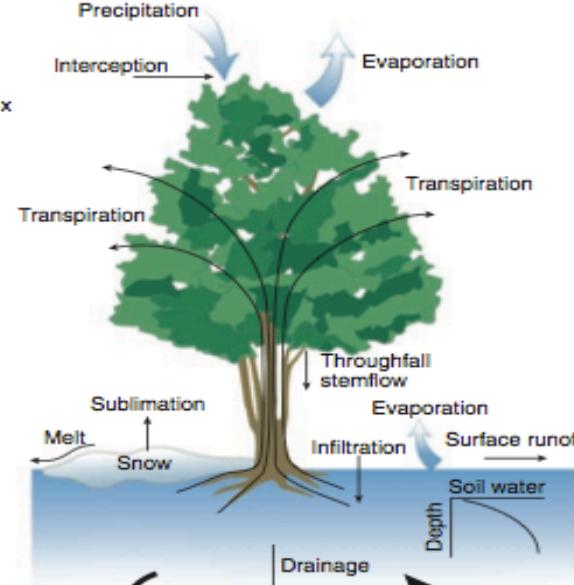
Fig. 2. Global distribution of NPP ($\times 10 \text{ gm C/m}^2/\text{yr}$) at the tracer model resolution.

Forests in Flux

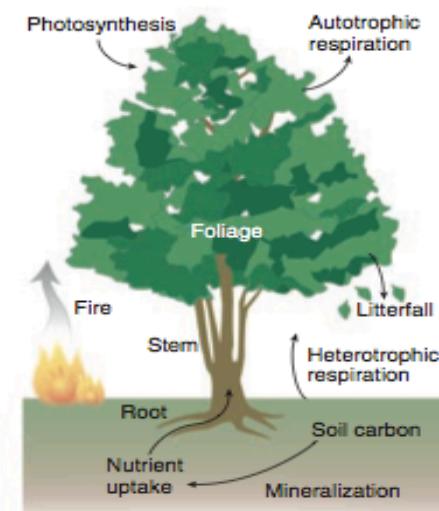
A Surface energy fluxes



B Hydrology



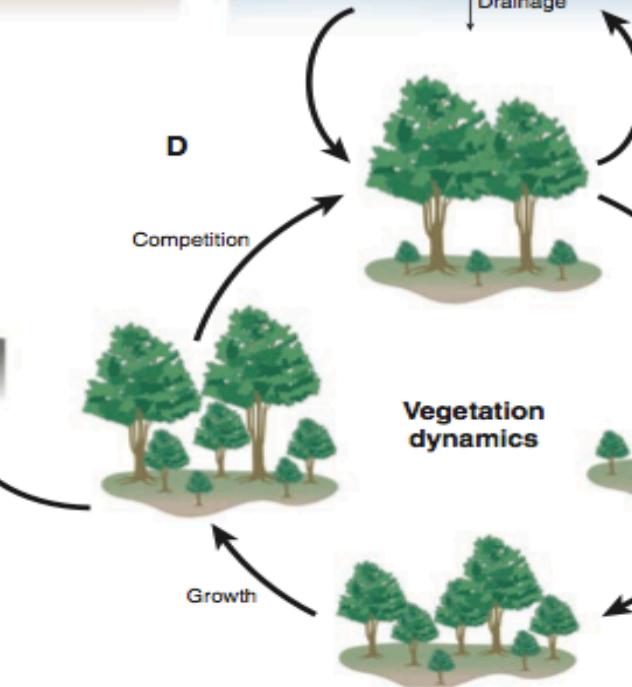
C Carbon Cycle



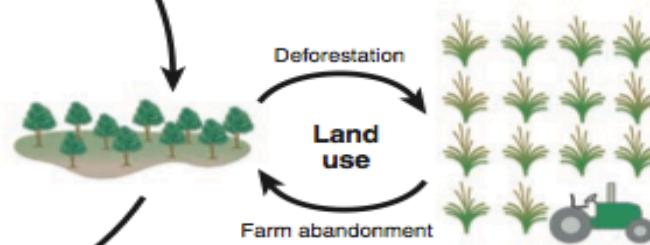
F Urbanization



D Vegetation dynamics



E Land use

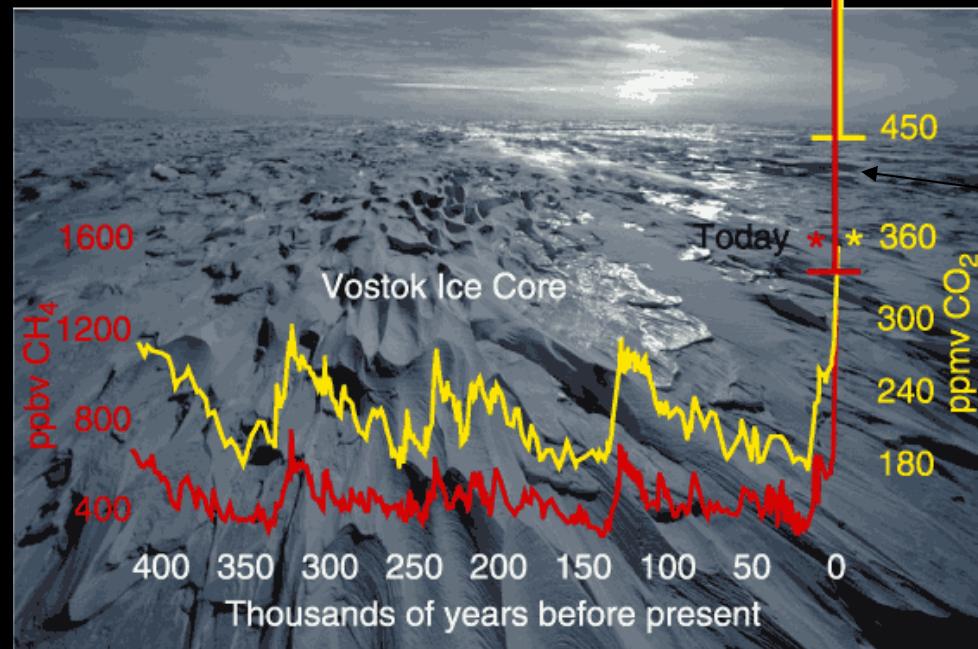


Bonan 2008

Why does it matter?



Atmospheric CO₂
has increased rapidly
to levels above
anything in Earth's
recent past

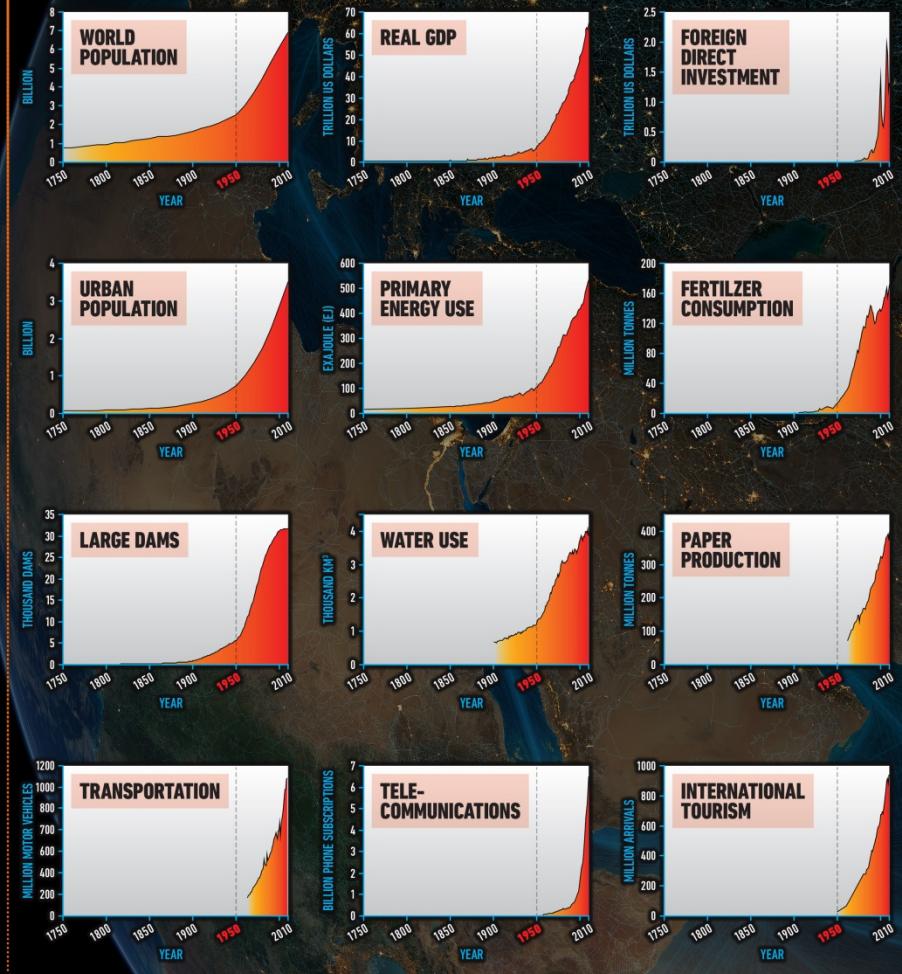


Today
400 ppm CO₂
2 ppm CH₄

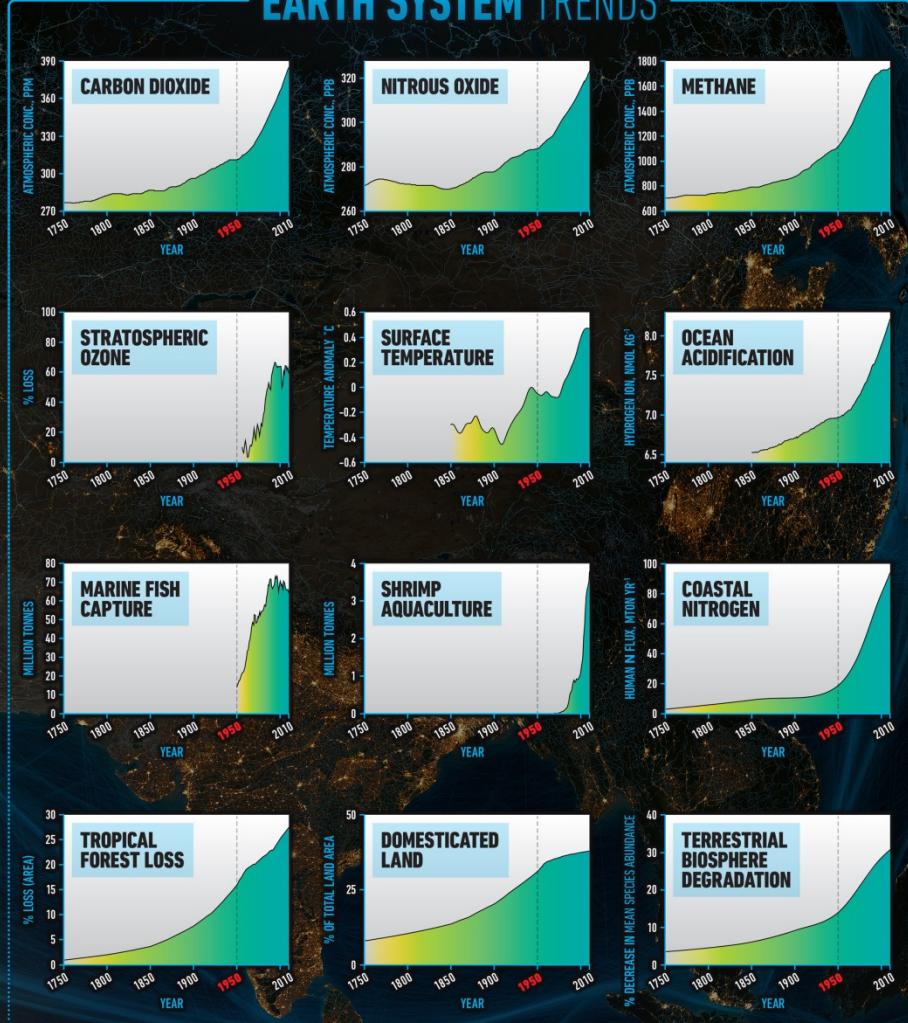
Sources: Petit et al
(1999) Nature
399:429-436 and
IPCC(2000)

THE GREAT ACCELERATION

SOCIO-ECONOMIC TRENDS



EARTH SYSTEM TRENDS



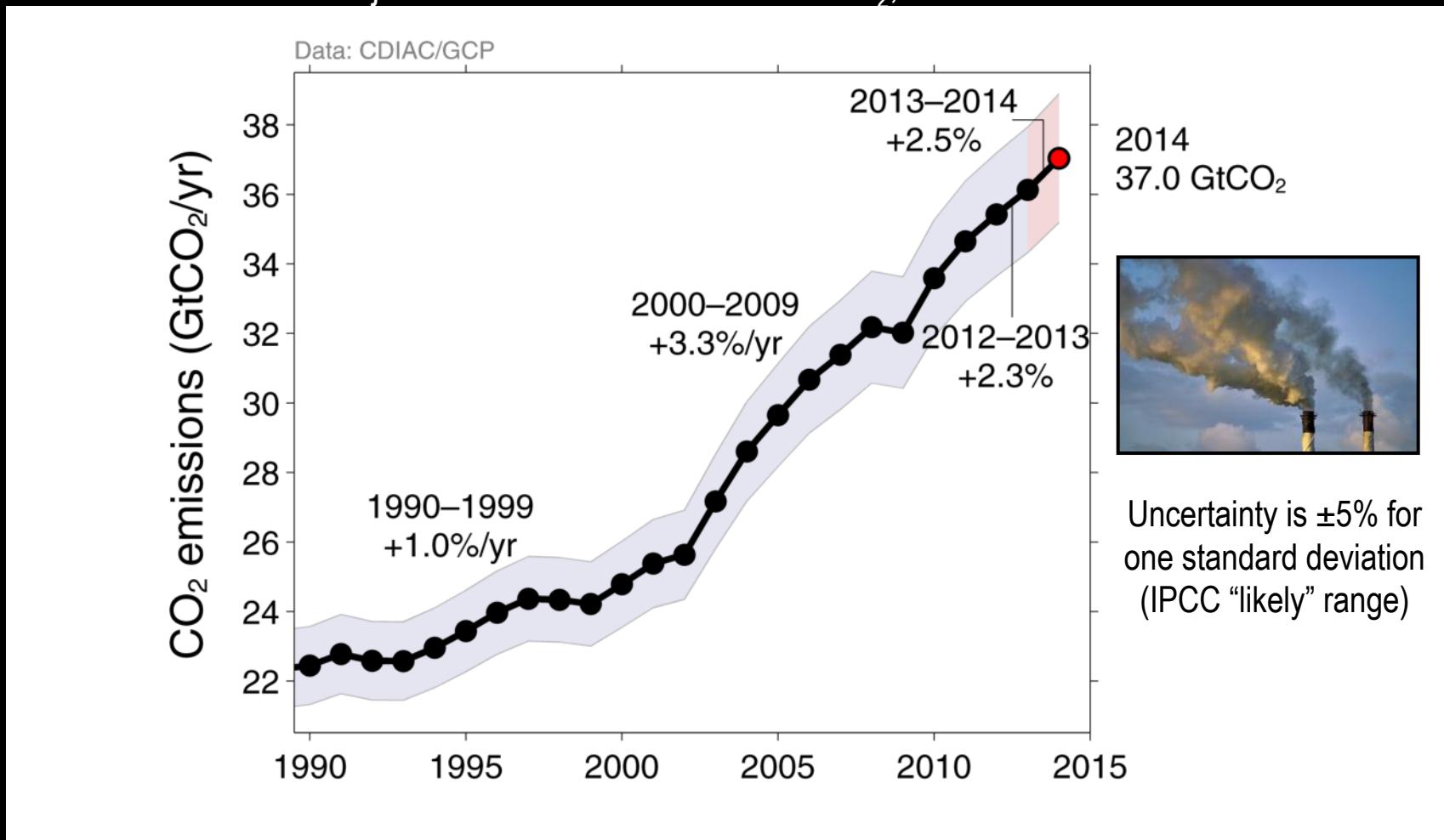
REFERENCE: Steffen, W., W. Broadgate, L. Deutsch, O. Gaffney and C. Ludwig (2015). The Trajectory of the Anthropocene: the Great Acceleration, Submitted to *The Anthropocene Review*.

MAP & DESIGN: Félix Pharand-Deschênes / Globaïa

Fossil Fuel and Cement Emissions

Global fossil fuel and cement emissions: $36.1 \pm 1.8 \text{ GtCO}_2$ in 2013, 61% over 1990

- Projection for 2014 : $37.0 \pm 1.9 \text{ GtCO}_2$, 65% over 1990



Estimates for 2011, 2012, and 2013 are preliminary
 Source: [CDIAC](#); [Le Quéré et al 2014](#); [Global Carbon Budget 2014](#)

Global Carbon Budget

The cumulative contributions to the Global Carbon Budget from 1870
Contributions are shown in parts per million (ppm)

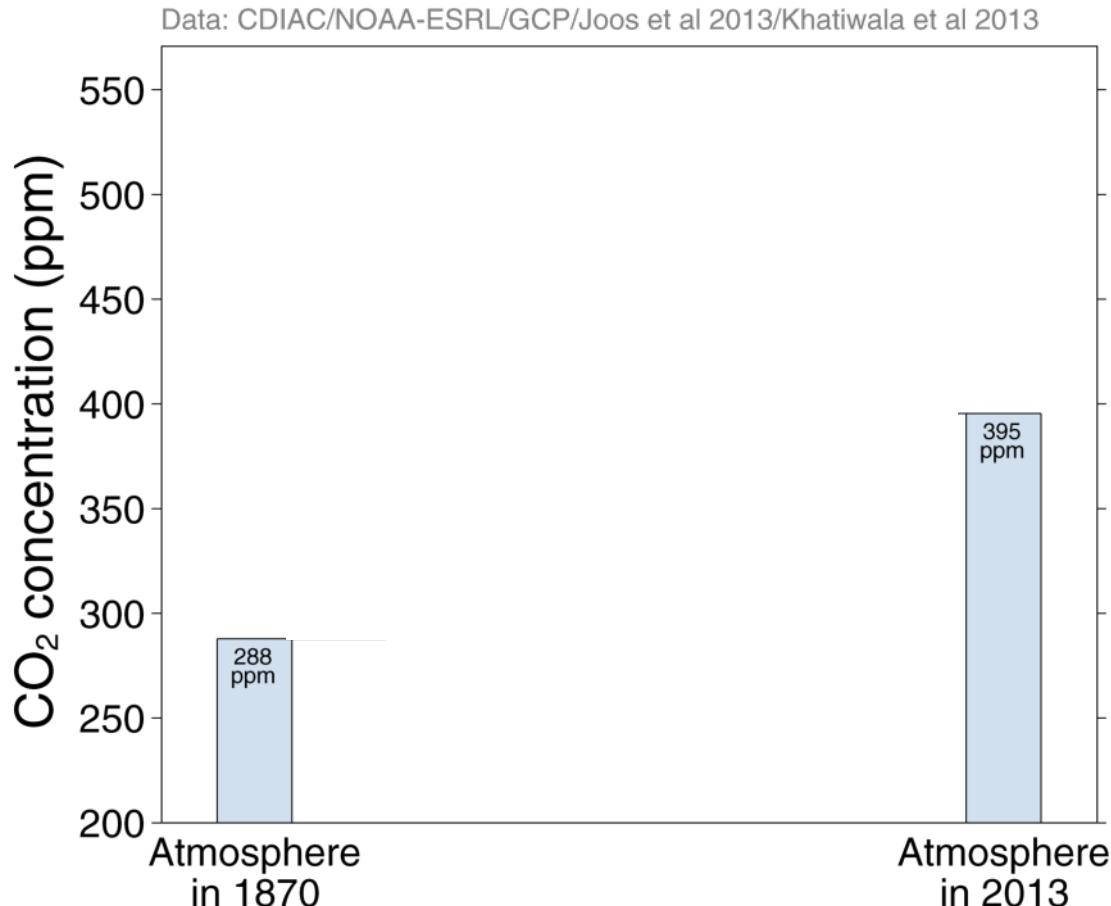
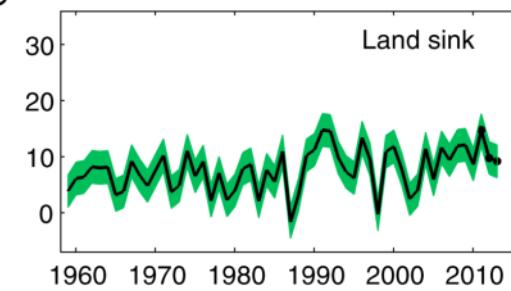
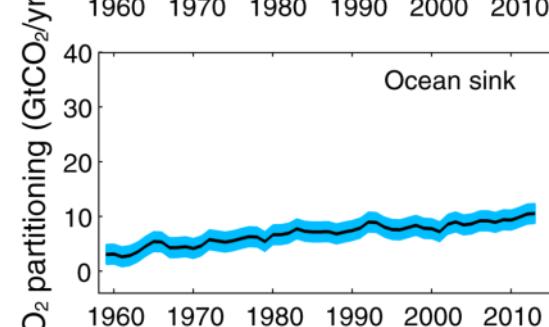
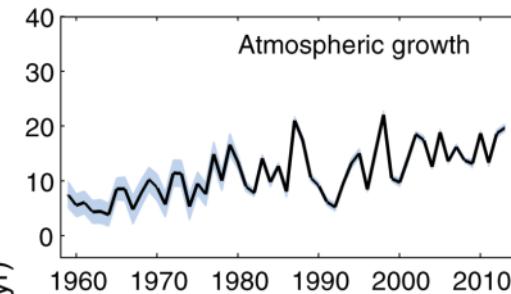
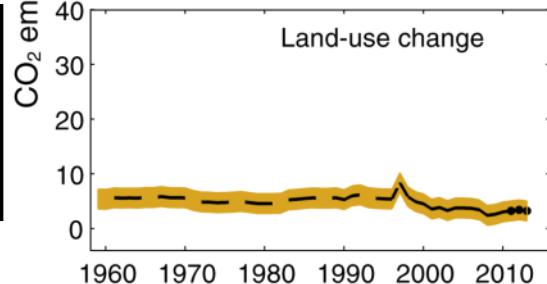
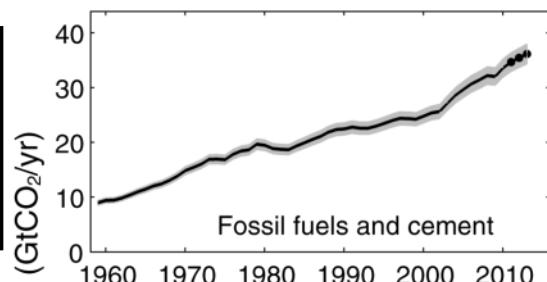


Figure concept from [Shrink That Footprint](#)

Source: [CDIAC](#); [NOAA-ESRL](#); [Houghton et al 2012](#); [Giglio et al 2013](#); [Joos et al 2013](#); [Khatiwala et al 2013](#); [Le Quéré et al 2014](#); [Global Carbon Budget 2014](#)

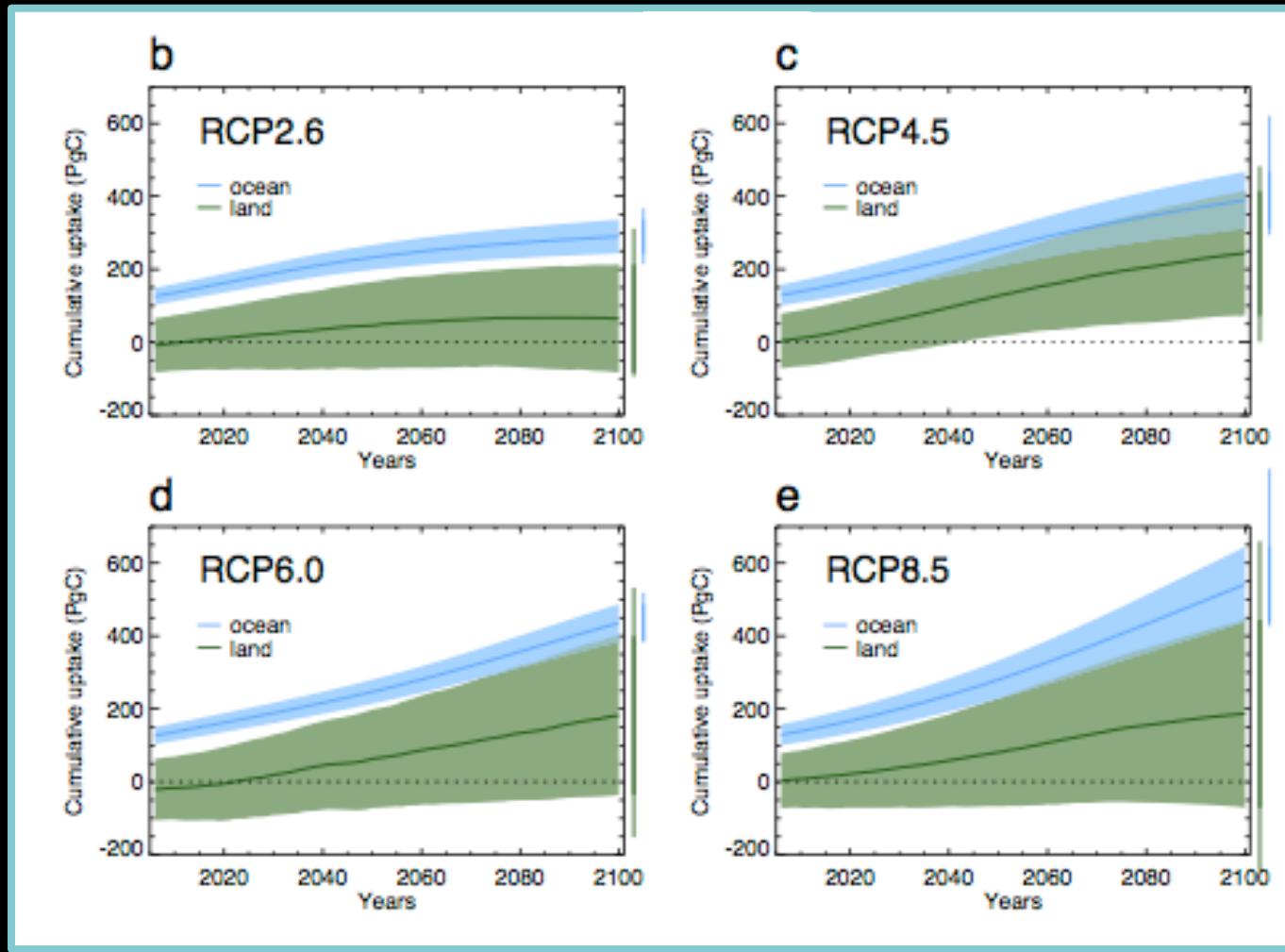
Changes in the Budget over Time

The sinks have continued to grow with increasing emissions, but climate change will affect carbon cycle processes in a way that will exacerbate the increase of CO₂ in the atmosphere

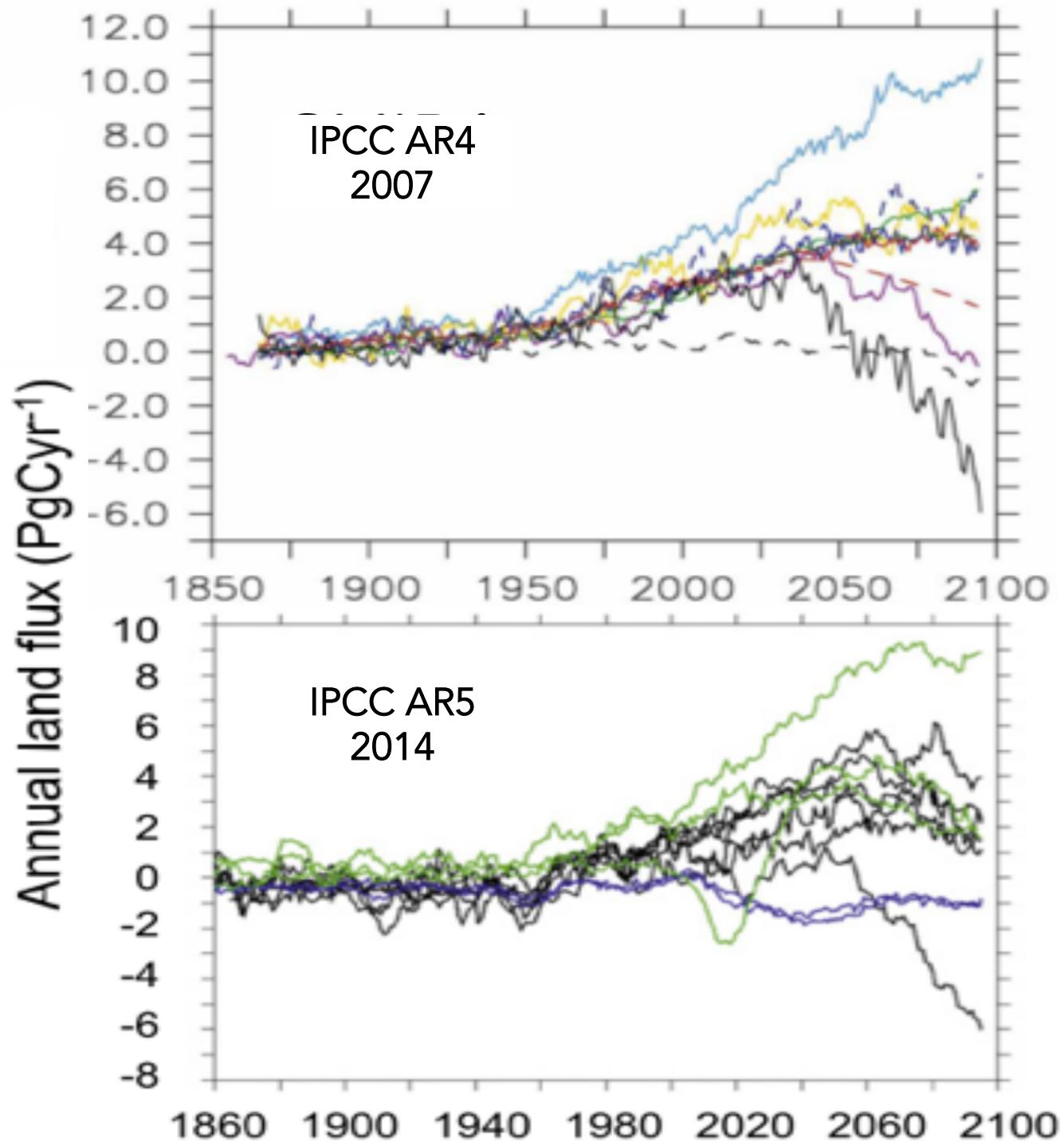


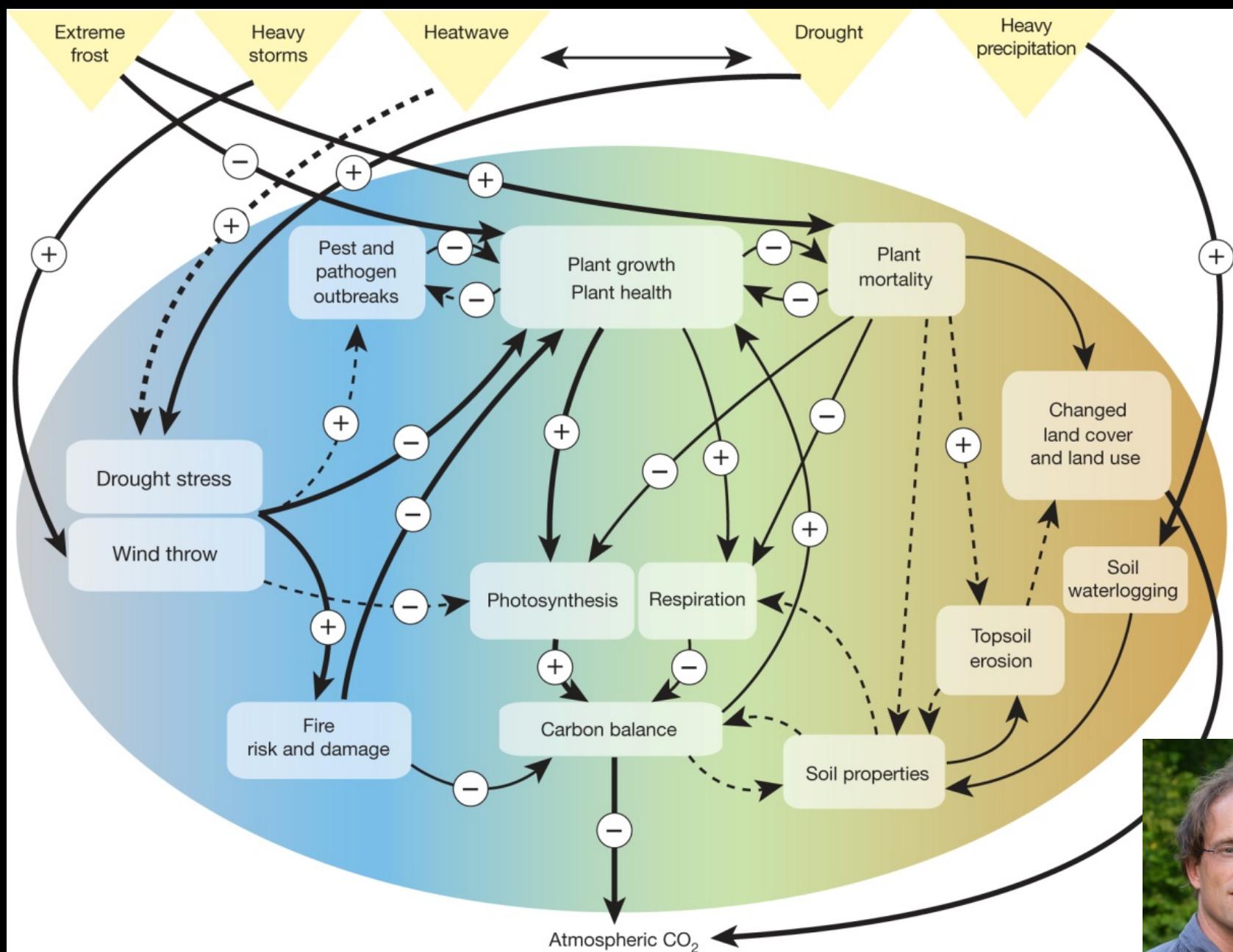
Source: [CDIAC](#); [NOAA-ESRL](#); [Houghton et al 2012](#); [Giglio et al 2013](#); [Le Quéré et al 2014](#); [Global Carbon Budget 2014](#)

Terrestrial carbon cycle feedback is a leading order uncertainty for climate simulation



IPCC AR5 WG1 CH6

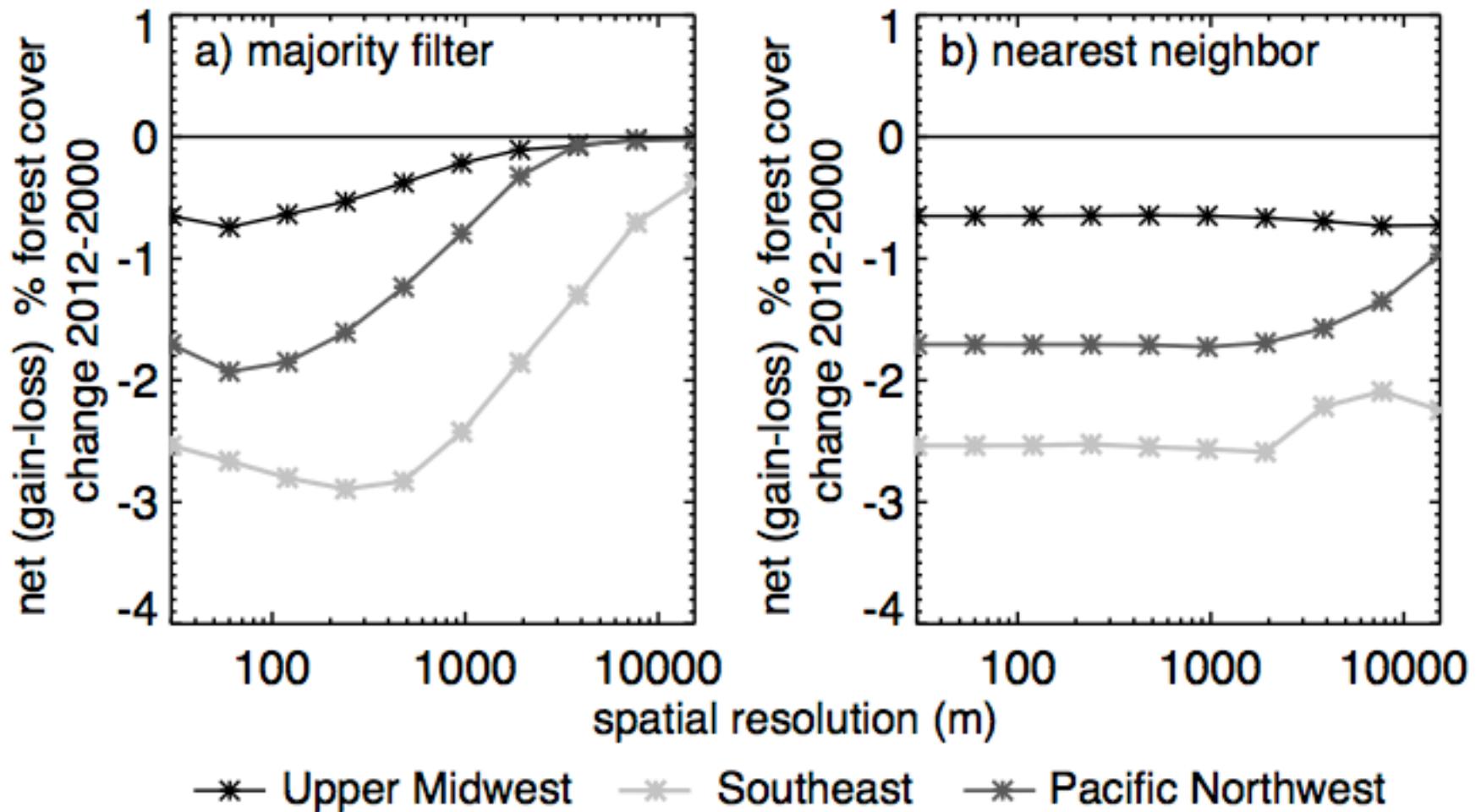




M Reichstein et al. *Nature* 500, 287-295 (2013) doi:10.1038/nature12350



The scale and method we monitor land use matters



Becknell et al., Bioscience, 2015

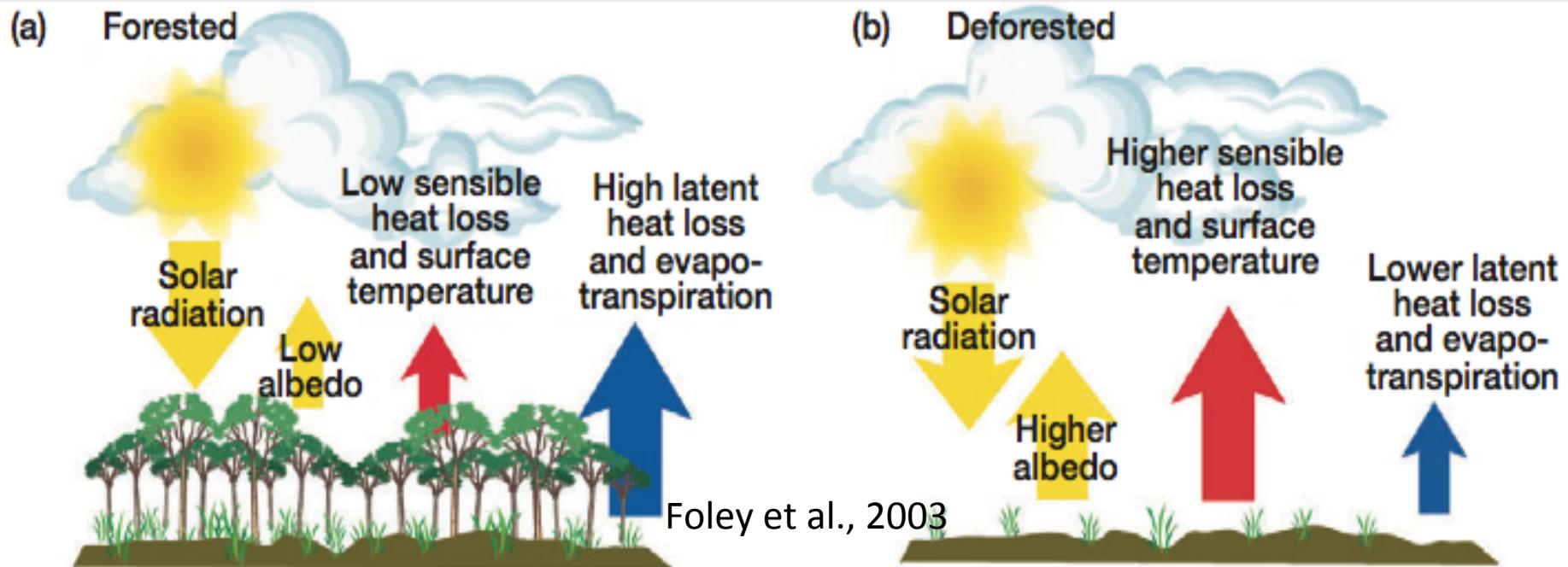
What do I (we) do?

<http://flux.aos.wisc.edu>

- Probe spatial heterogeneity in biologically-mediated surface-atmosphere exchanges from sites to regions (meters-1000s km)
 - Forests, wetlands, lakes, urban (temperate-boreal-tropical-Mediterranean-alpine, terrestrial-aquatic, management gradients)
 - Multiple greenhouse gases (methane), esp. with eddy covariance
 - Feedbacks from energy balance and a land surface variability on the atmospheric boundary layer and synoptic-PBL interactions in observations and models (LES, PBL, mesoscale, climate)
 - Up/down scaling across multiple measurements: eddy covariance, biometric, airborne budgets, inverse modeling, hyperspectral remote sensing (leaf to satellite)
 - Informing ecosystem and atmospheric models with diverse measurements across space (data assimilation, model informatics)
 - <http://pecanproject.org>

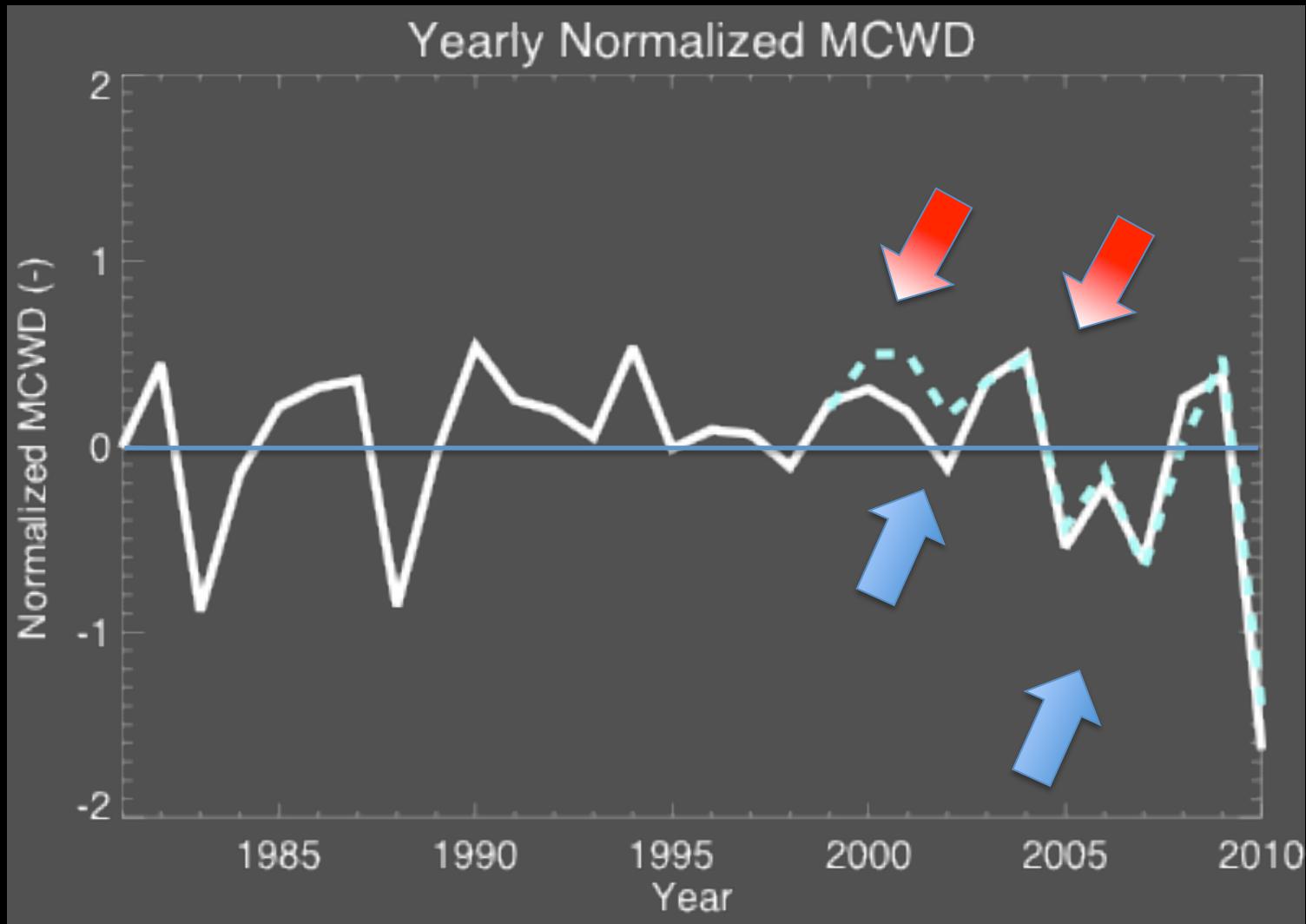


Bagley et al., 2014,
J. Clim

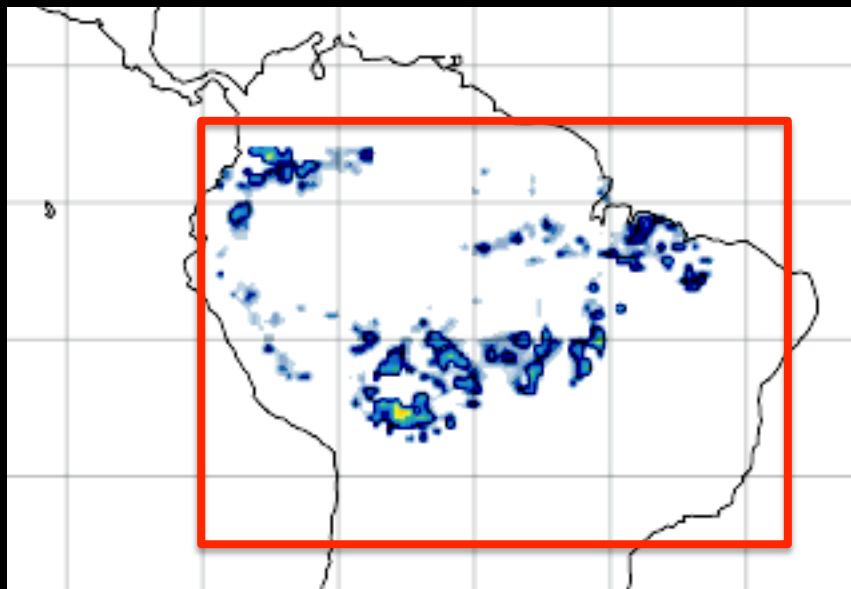


A history of drought and floods in the Amazon

Bagley et al., 2014, J. Clim

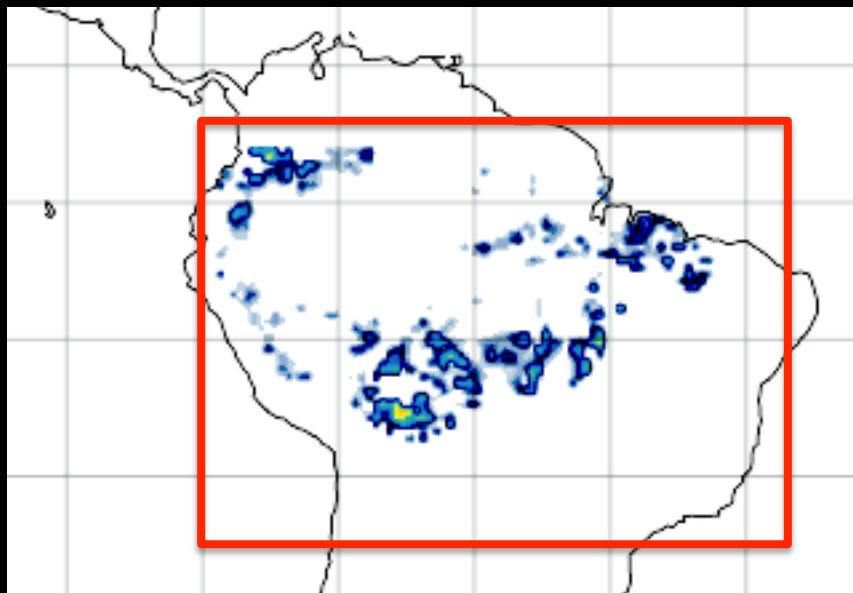


WRF-Noah Setup

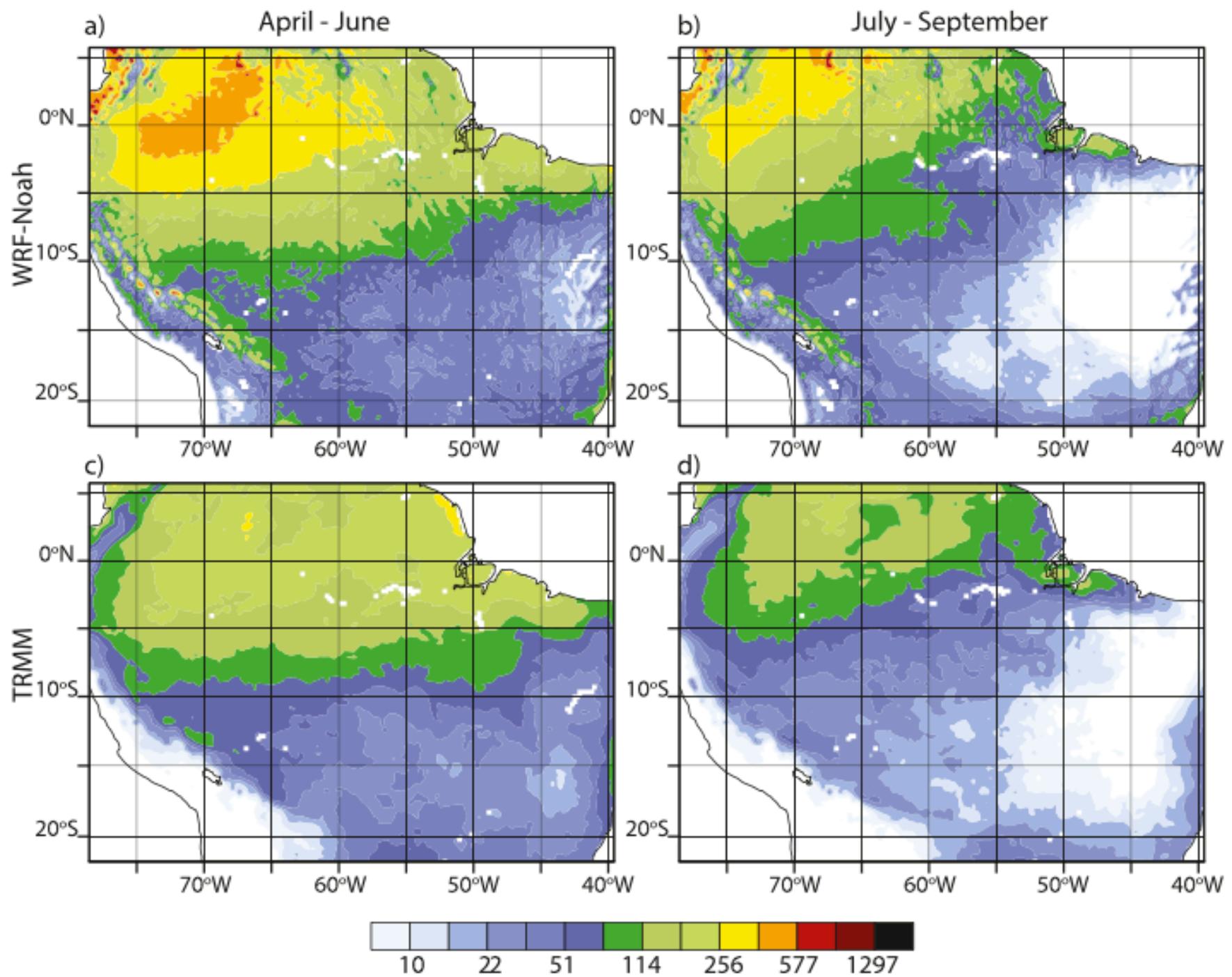


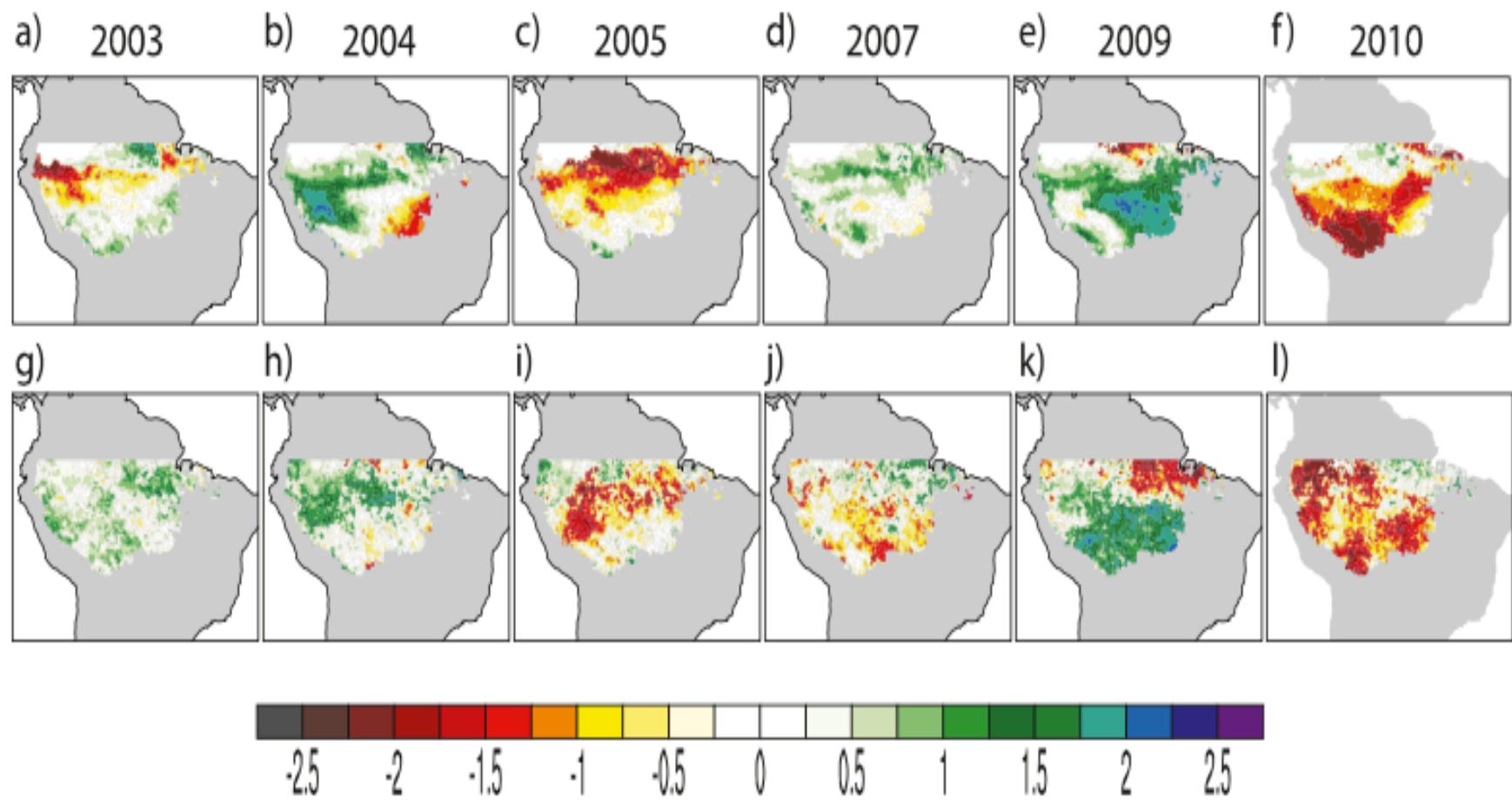
- Spatial Resolution: 20km x 20km
- Timestep: 60 seconds
- For 2003, 2004, 2005, 2007, 2009, and 2010 the model was run from March 15 – October 15 with and without deforestation
- Total of 12 seven-month simulations completed with **hourly** output

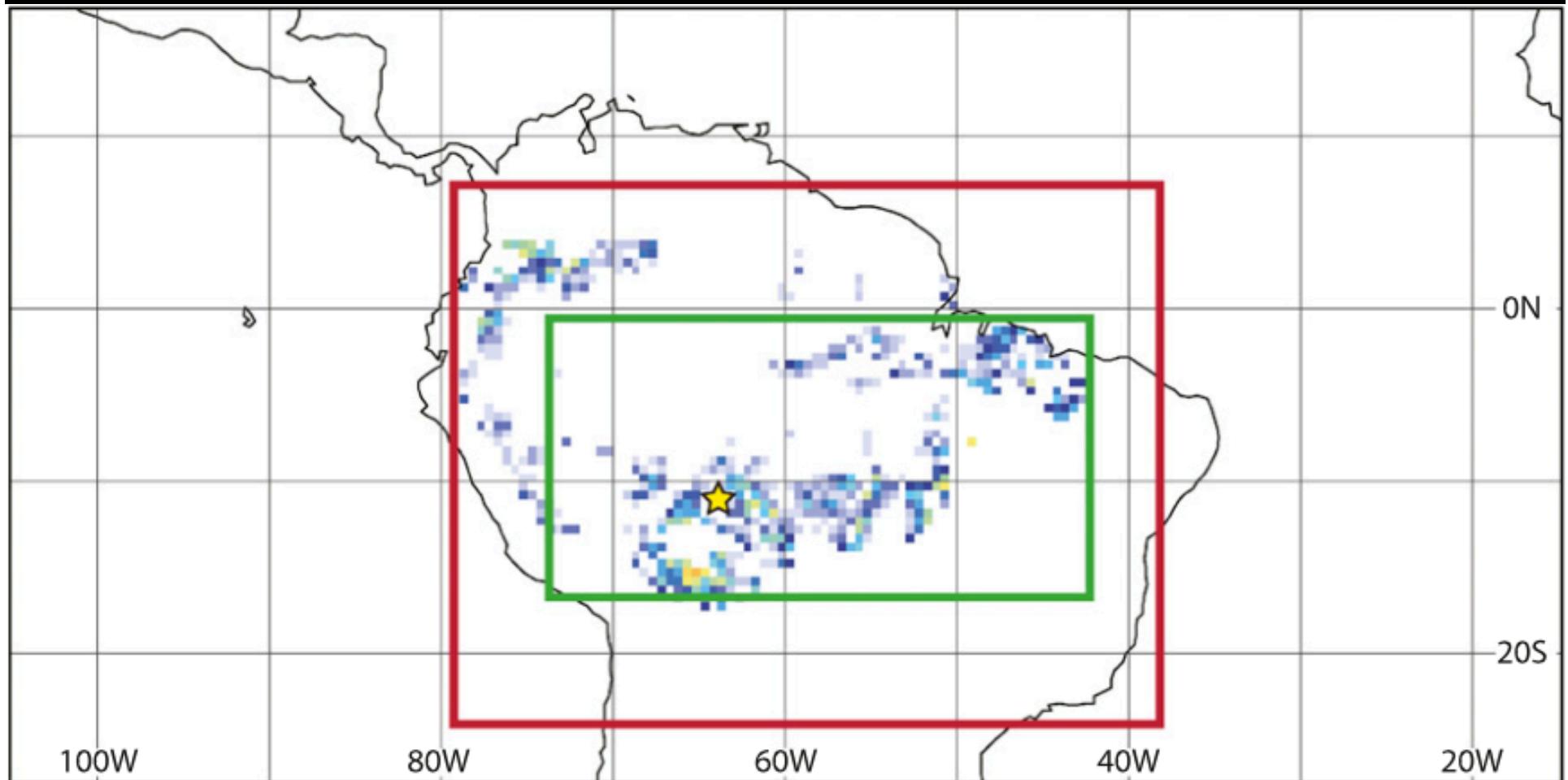
WRF-Noah Setup



- Gridpoints with land use > 50% converted to pasture
- Gridpoints with land use between 5%-50% converted to mixed forest and land use







ABRACOS/WRF/
PEGASUS



WRF-Noah



ABRACOS
Observations

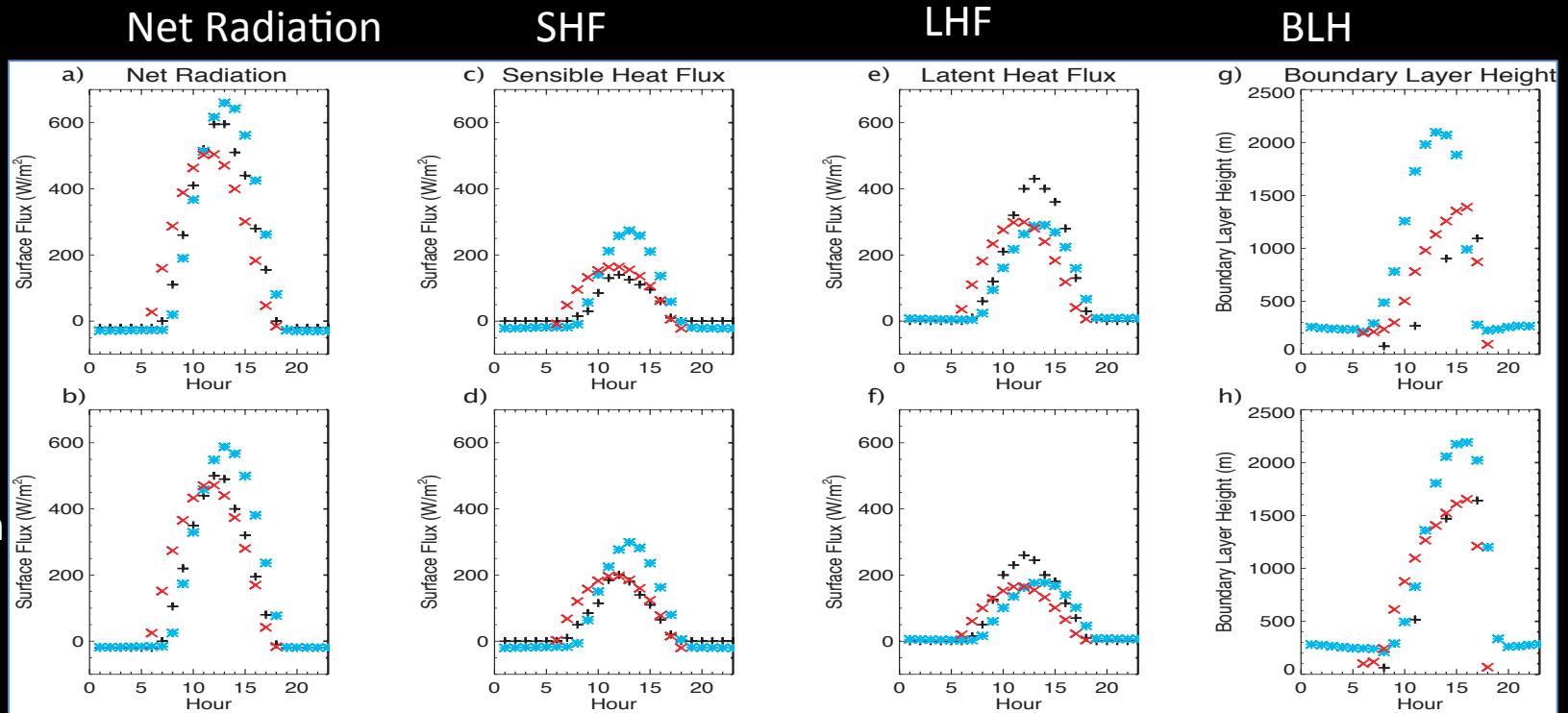


PEGASUS

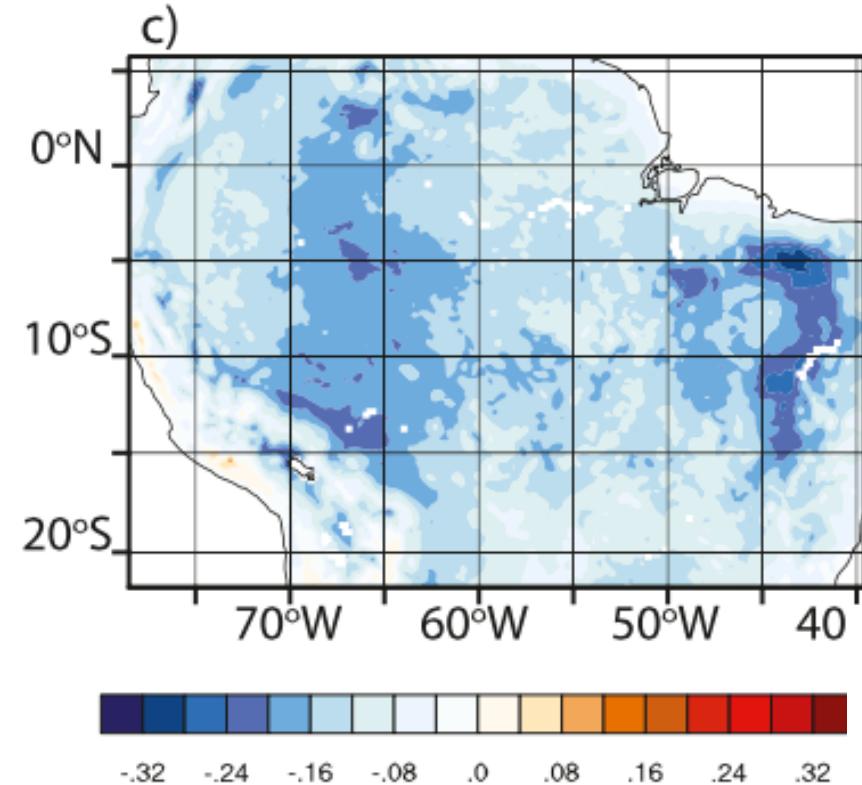
Potential
Vegetation

Deforestation

Impact

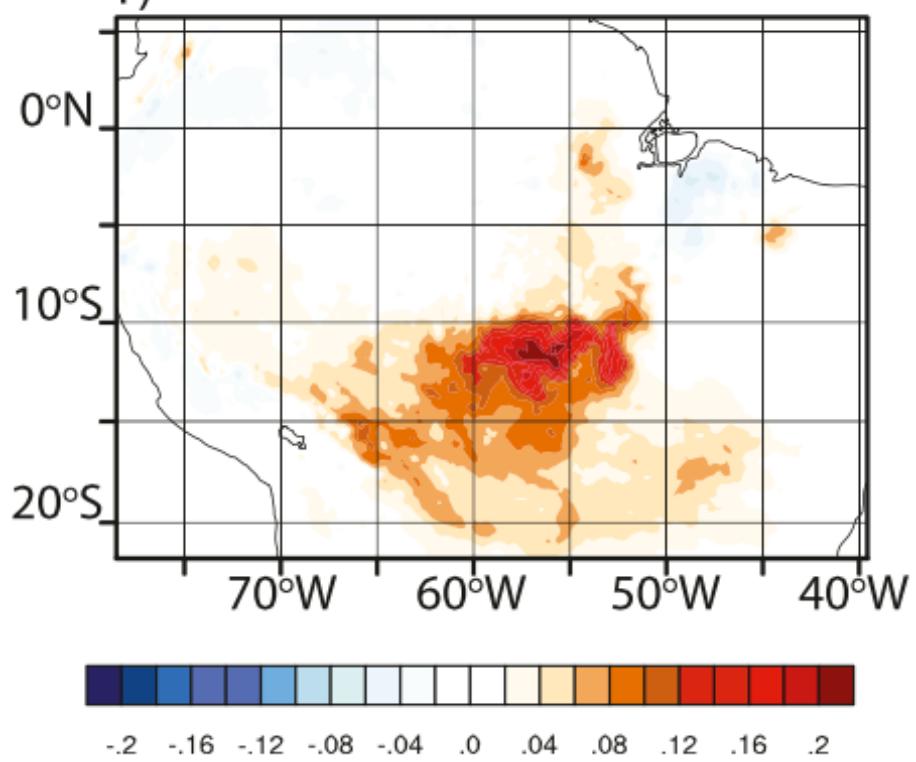


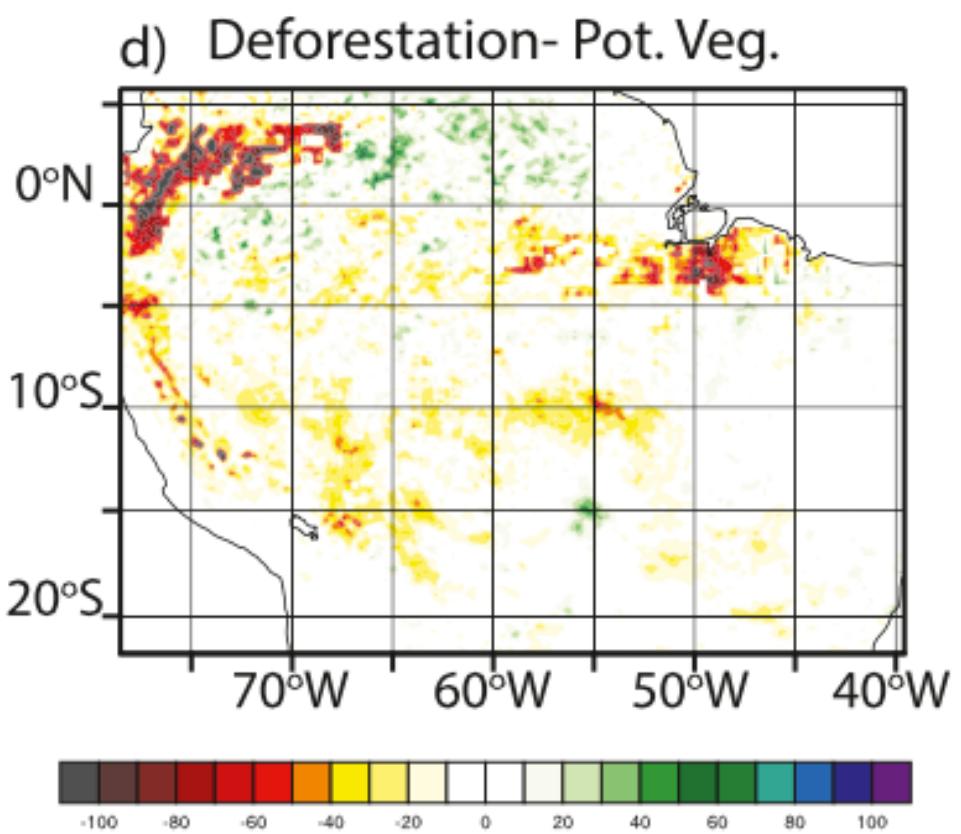
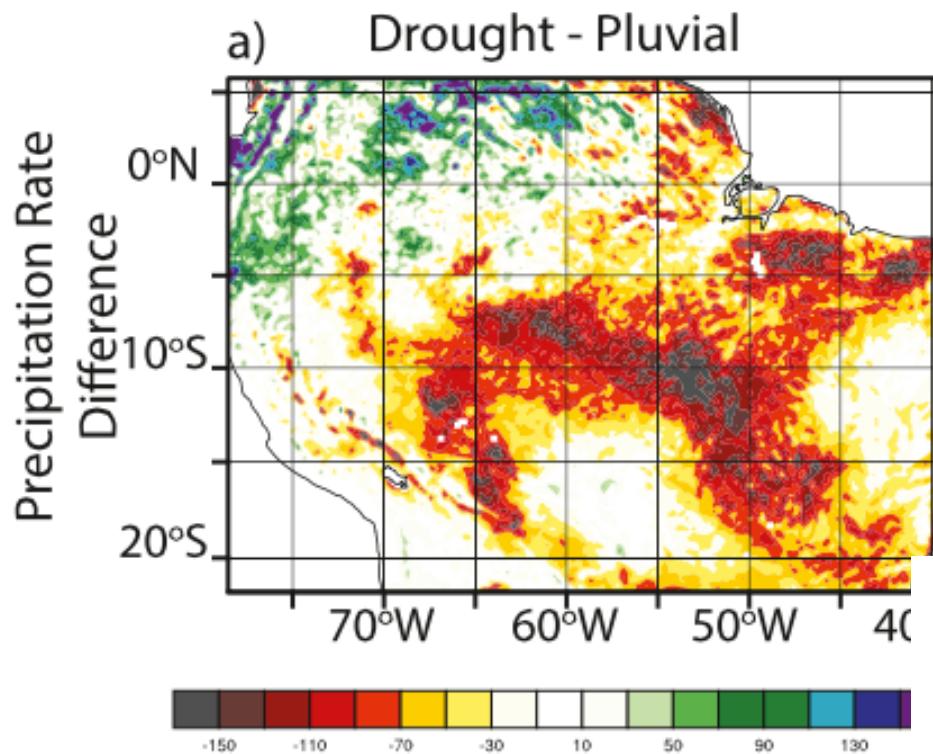
Temperature Difference



Drought-Pluvial

f)



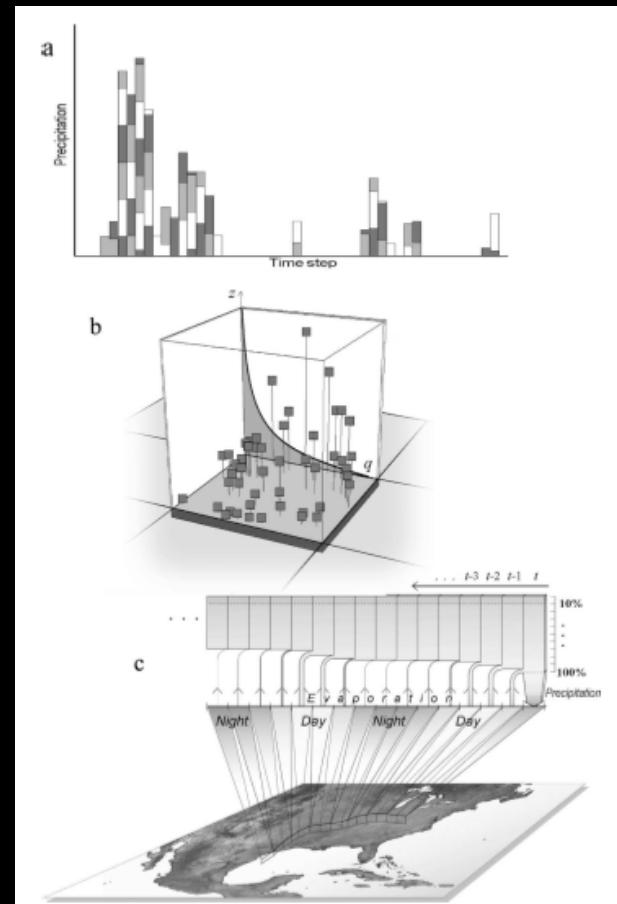


	Full Amazon rain forest region		Region converted to pasture	
	April–June	July–September	April–June	July–September
Precipitation rate (mm month^{-1})	-9.36 (-3.16%)	-9.47 (-5.45%)	-32.51 (-16.27%)	-25.72 (-18.53%)
Sensible heat flux (W m^{-2})	-0.758 (-2.75%)	1.04 (+2.30%)	-0.14 (-0.65%)	6.26 (+11.44%)
ET (mm month^{-1})	-2.68 (-1.94%)	-5.95 (-4.61%)	-15.89 (-13.42%)	-28.37 (-29.55%)
Net surface radiation (W m^{-2})	-4.39 (-2.45%)	-4.82 (-2.50%)	-4.02 (-2.48%)	-10.11 (-6.04%)
Boundary layer height (m)	-6.21 (-1.27%)	3.90 (+.62%)	-5.95 (-1.28%)	28.29 (+4.25%)
2-m temperature (K)	-0.10	+0.036	-0.033	+0.32
2-m specific humidity (kg kg^{-1})	-5.74E-5 (-0.41%)	-1.50E-4 (-1.04%)	-1.84E-4 (-1.30%)	-4.67E-4 (-4.10%)
Lifting condensation level (m)	-4.49 (-0.56%)	31.31 (+2.31%)	7.08 (+.58%)	105.38 (+5.66%)

Back Trajectory Analysis Description

1. Identify precipitation event
2. Initialize 100 parcels at grid cell of precipitation at pseudo-random heights
3. Generally following isentropic lines follow parcels 14 days backward in time or until the parcel intersects the surface
4. As it passes over adjacent gridpoints assume the a portion of its moisture is given to it by the evapotranspiration occurring at that point
5. Aggregate parcels to get evaporative source of precipitation event

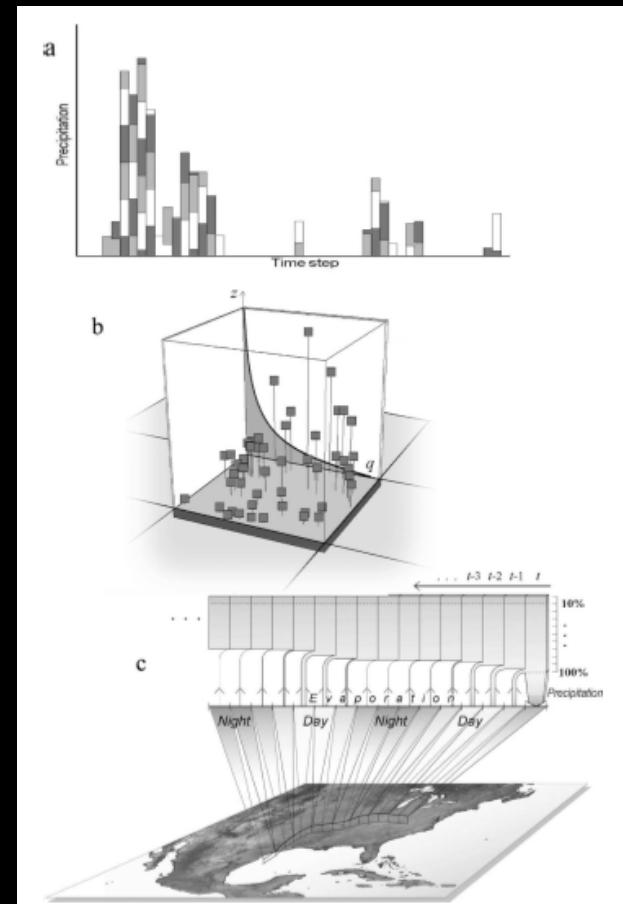
Precipitation, Recycling, and Land Memory: An Integrated Analysis
(Dirmeyer 2009)



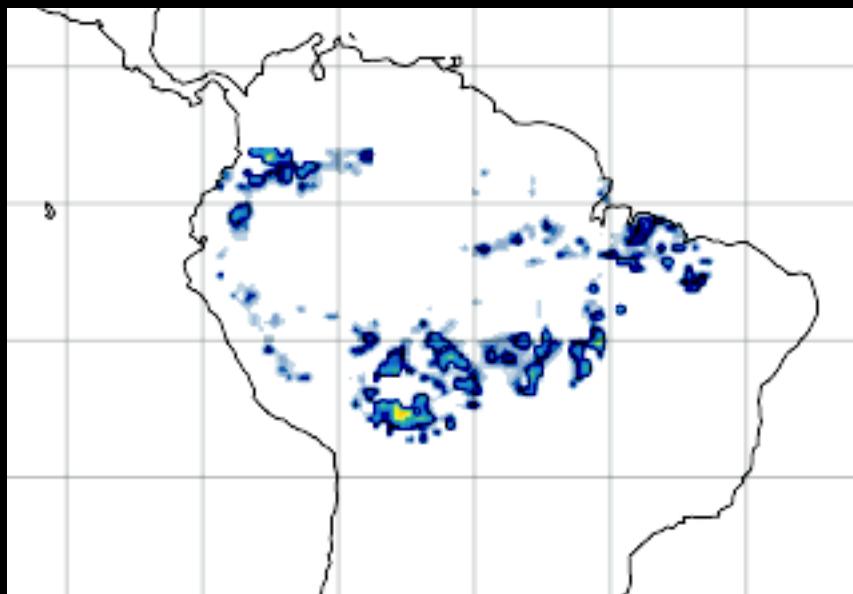
Back Trajectory Analysis Description

By compiling this information across all the precipitation events, we can invert the backtrajectories to determine where moisture evaporated from a given point tends to rain out of the atmosphere

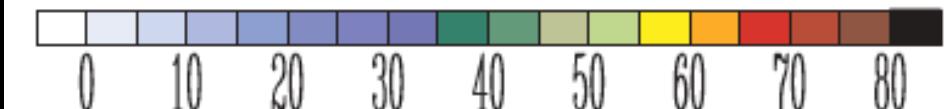
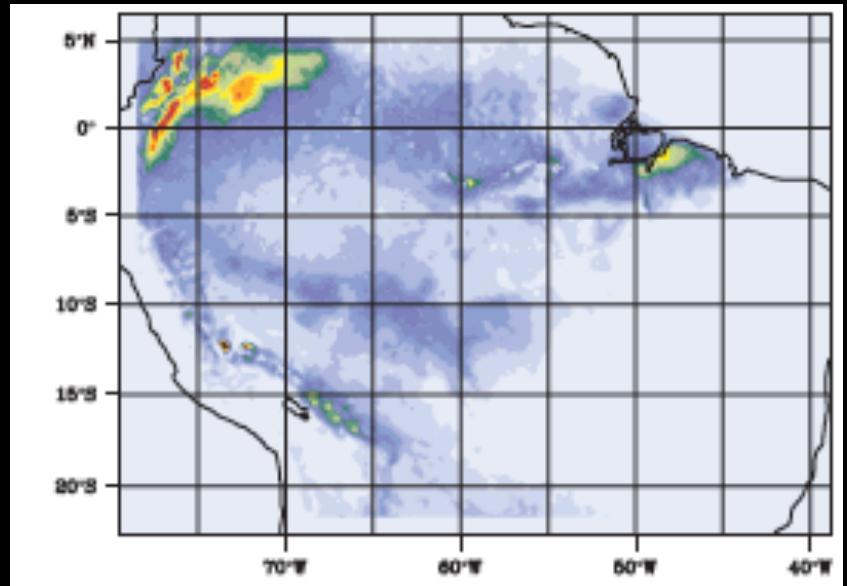
Precipitation, Recycling, and Land Memory: An Integrated Analysis
(Dirmeyer 2009)



Moisture Trajectory Analysis

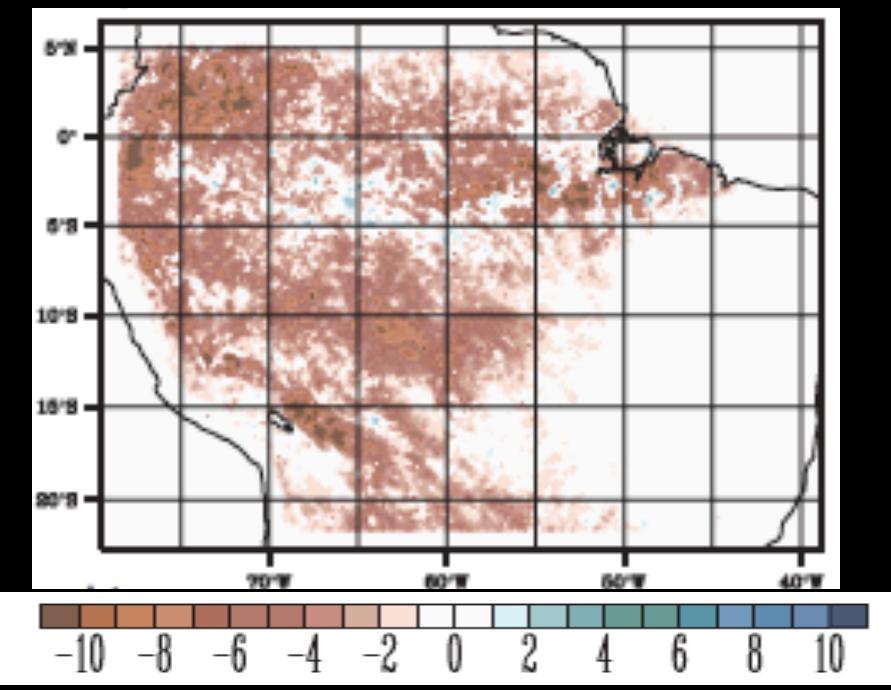


Deforested Regions

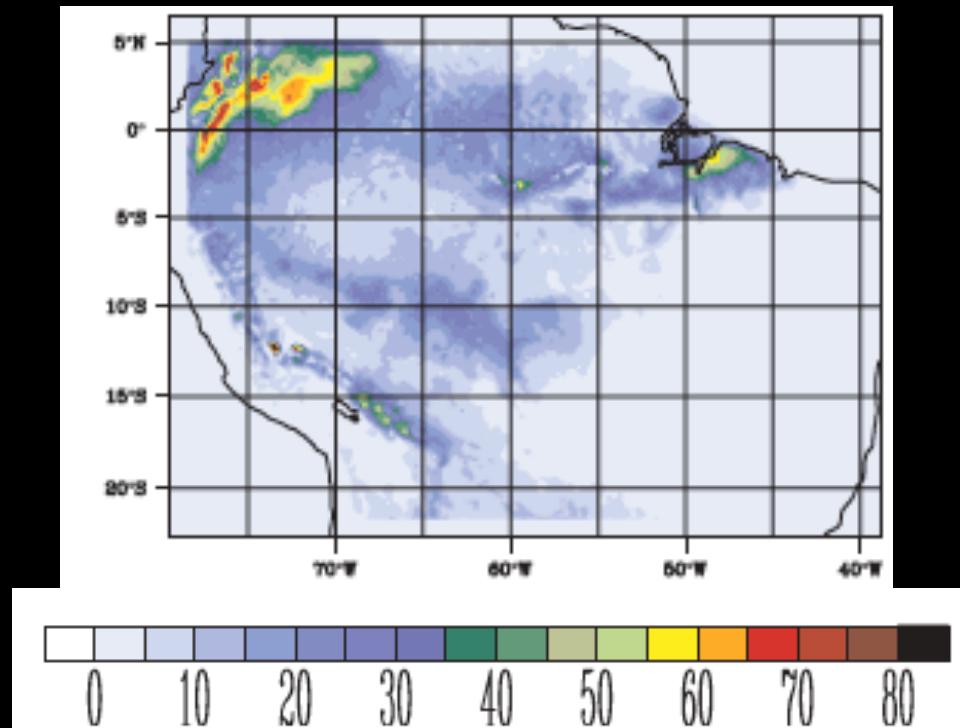


Mean forward trajectory
precipitation rate from deforested
points

Moisture Trajectory Analysis

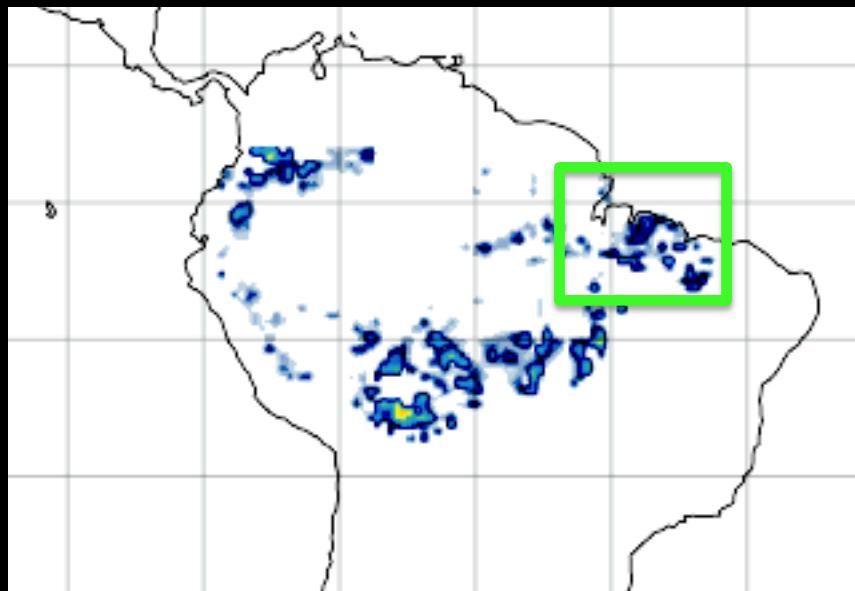


Impact of deforestation on precipitation rate from deforested points

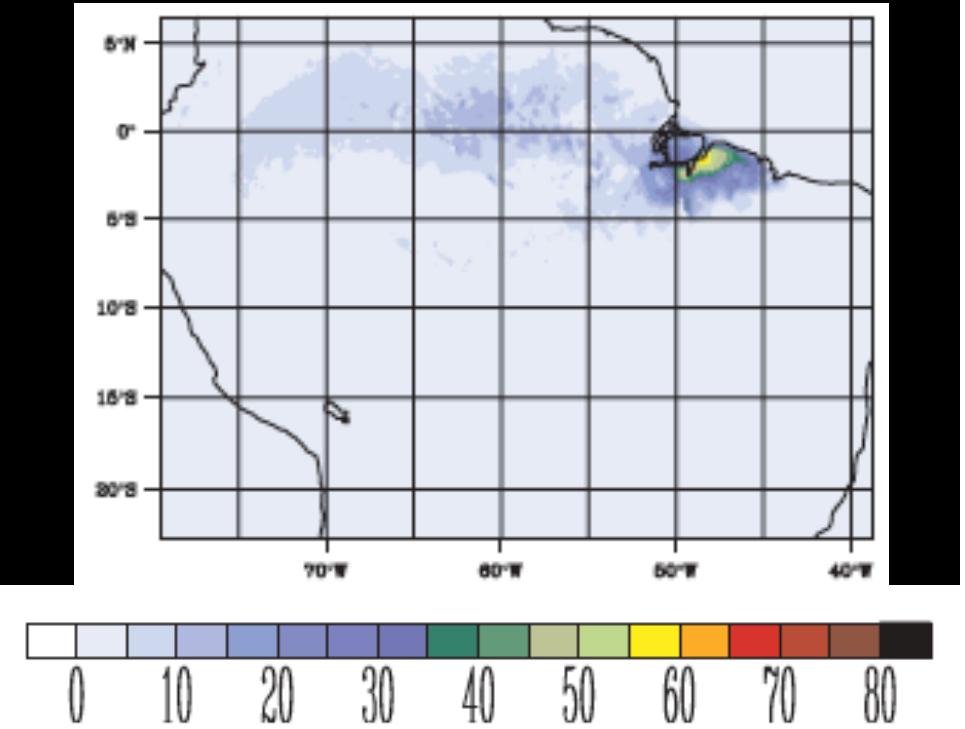


Mean forward trajectory precipitation rate from deforested points

Moisture Trajectory Analysis

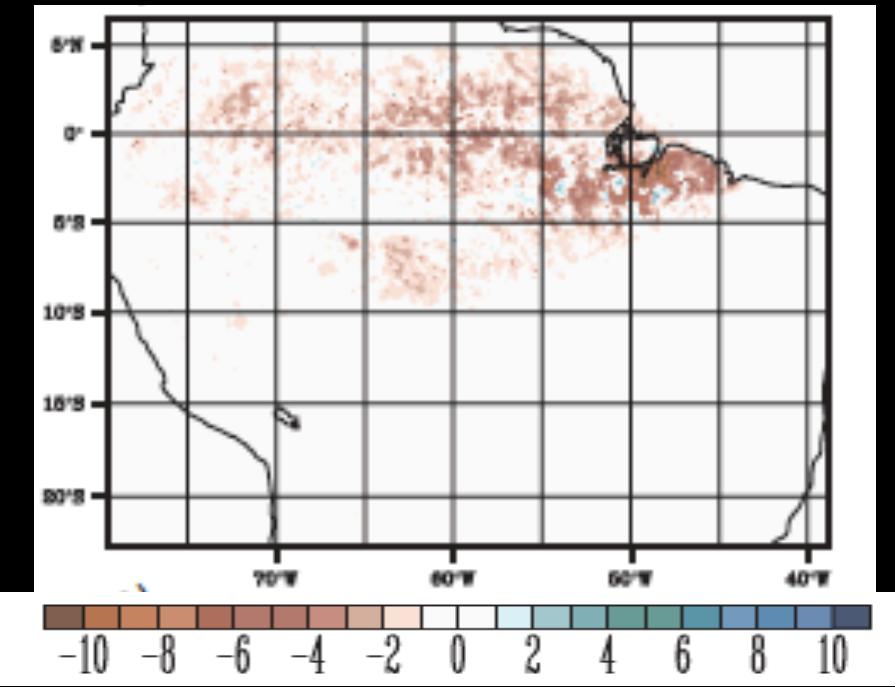


Source Region

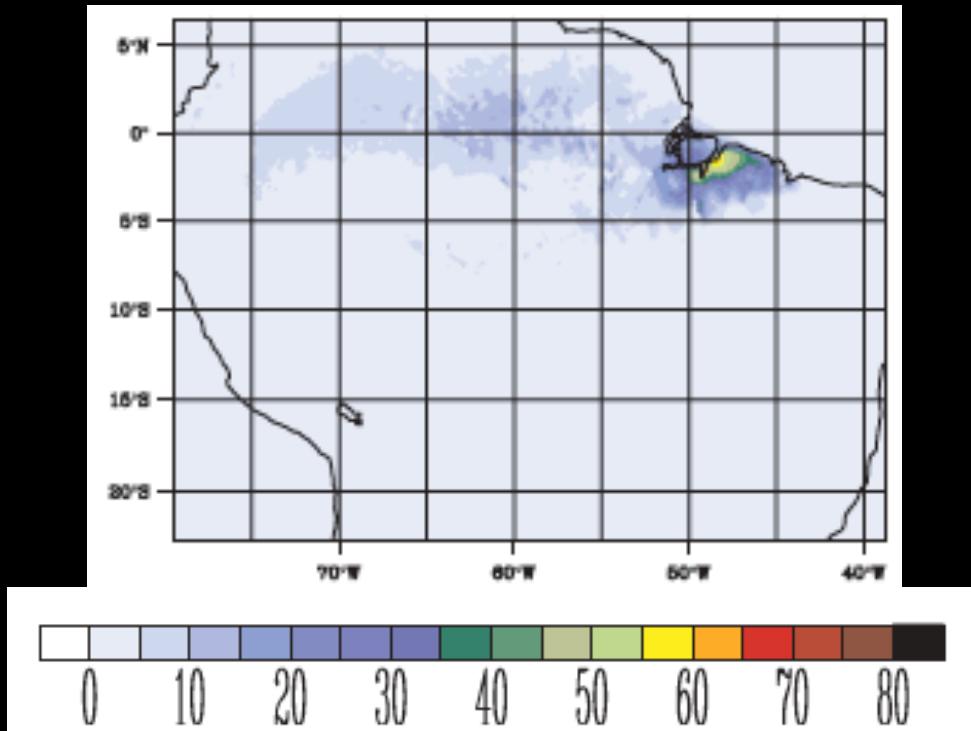


Mean forward trajectory
precipitation rate from deforested
points

Moisture Trajectory Analysis

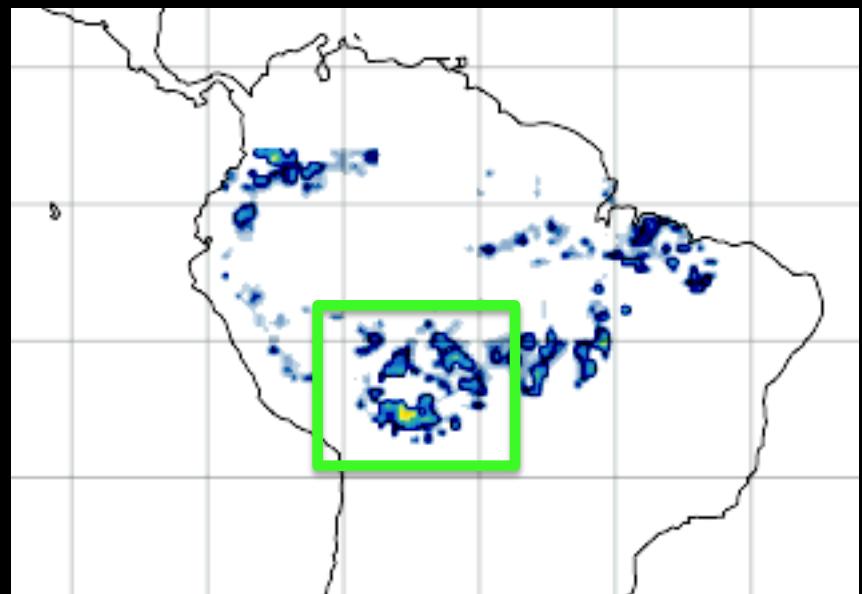


Impact of deforestation on
precipitation rate from deforested
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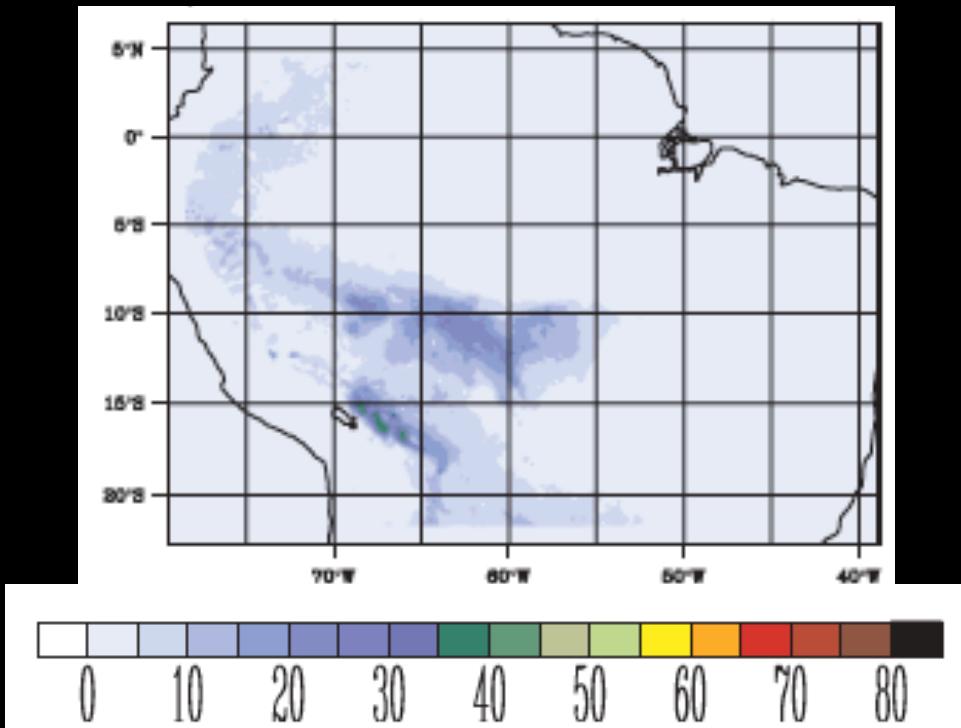


Mean forward trajectory
precipitation rate from deforested
points

Moisture Trajectory Analysis

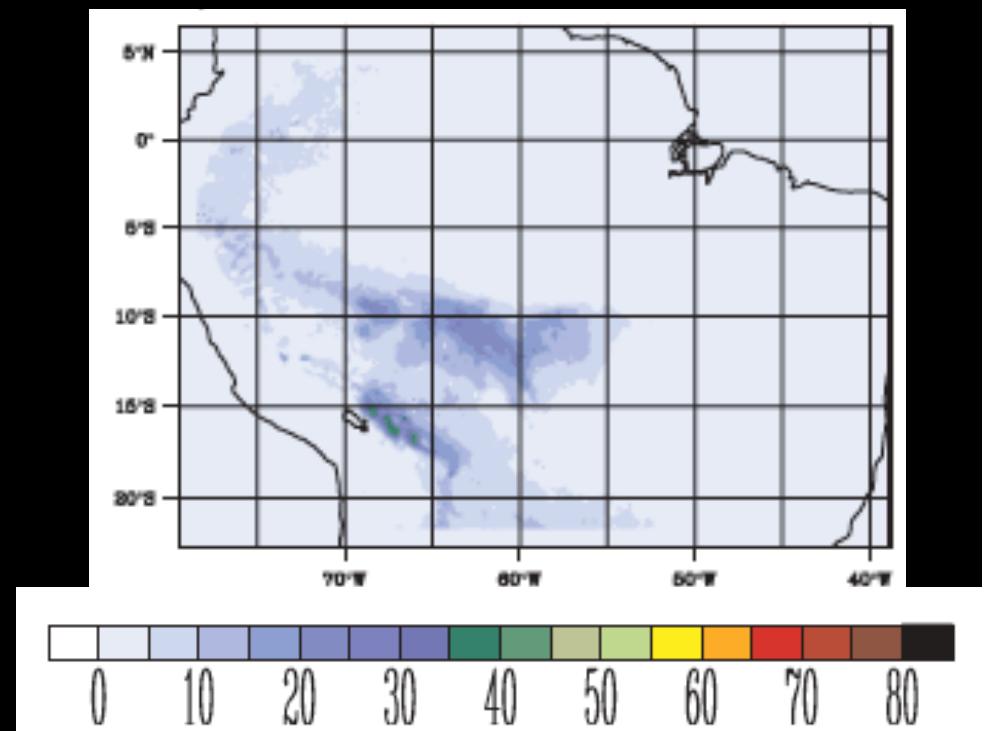
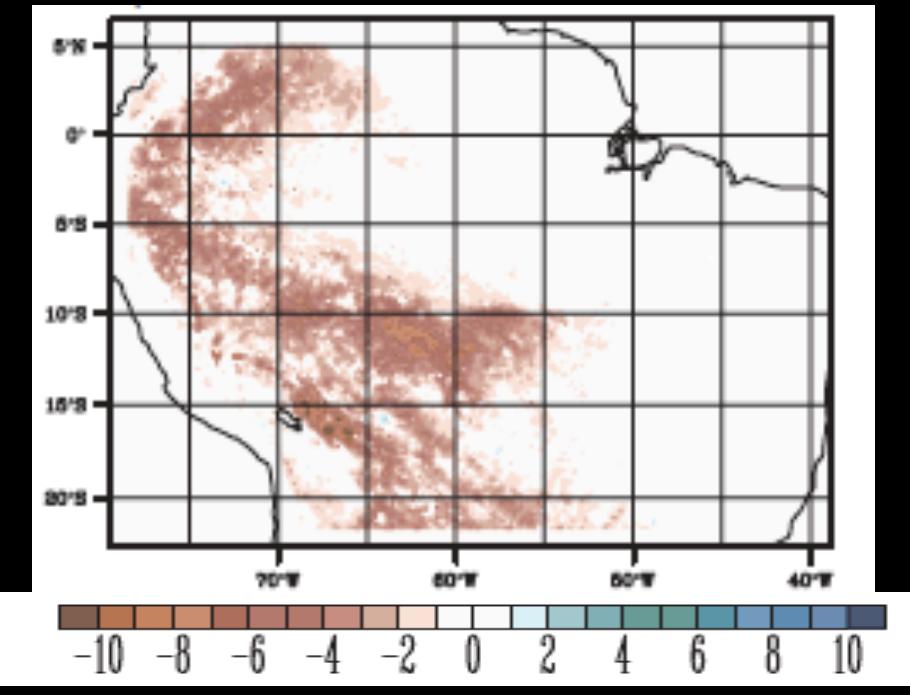


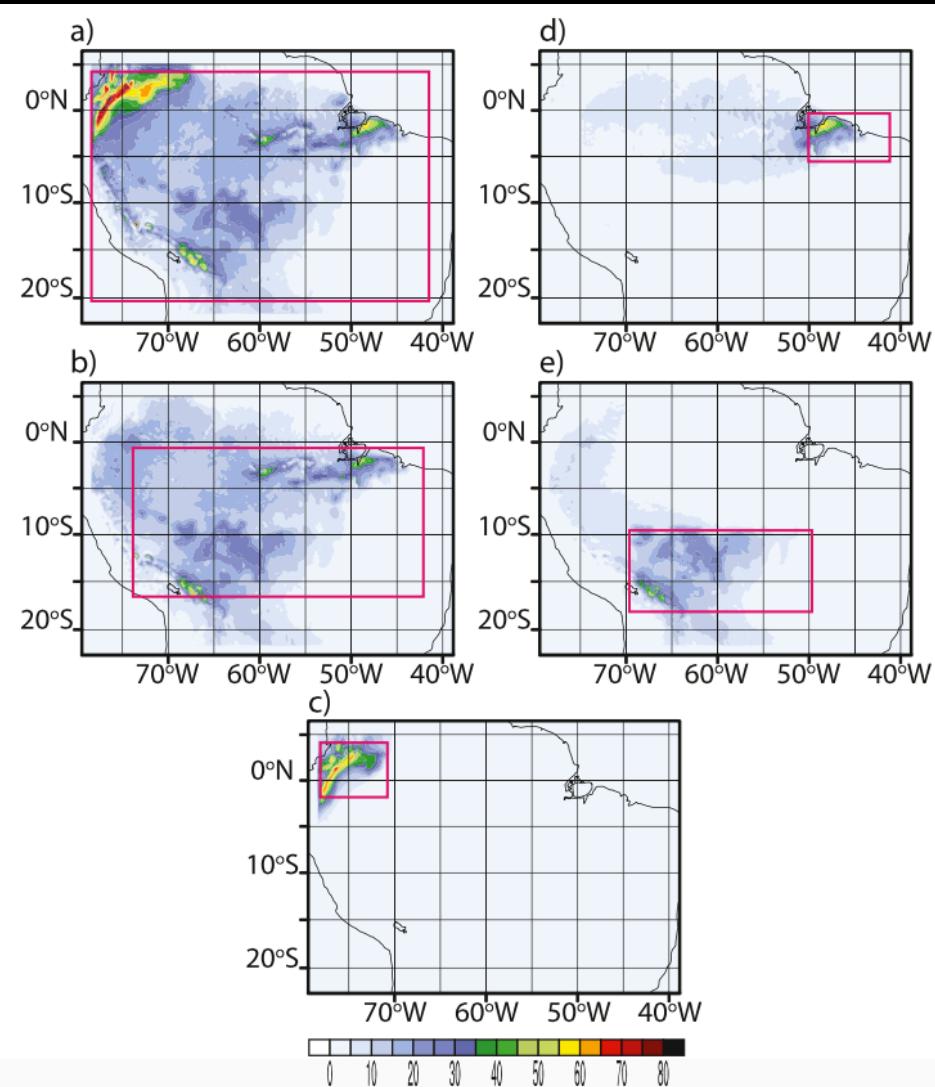
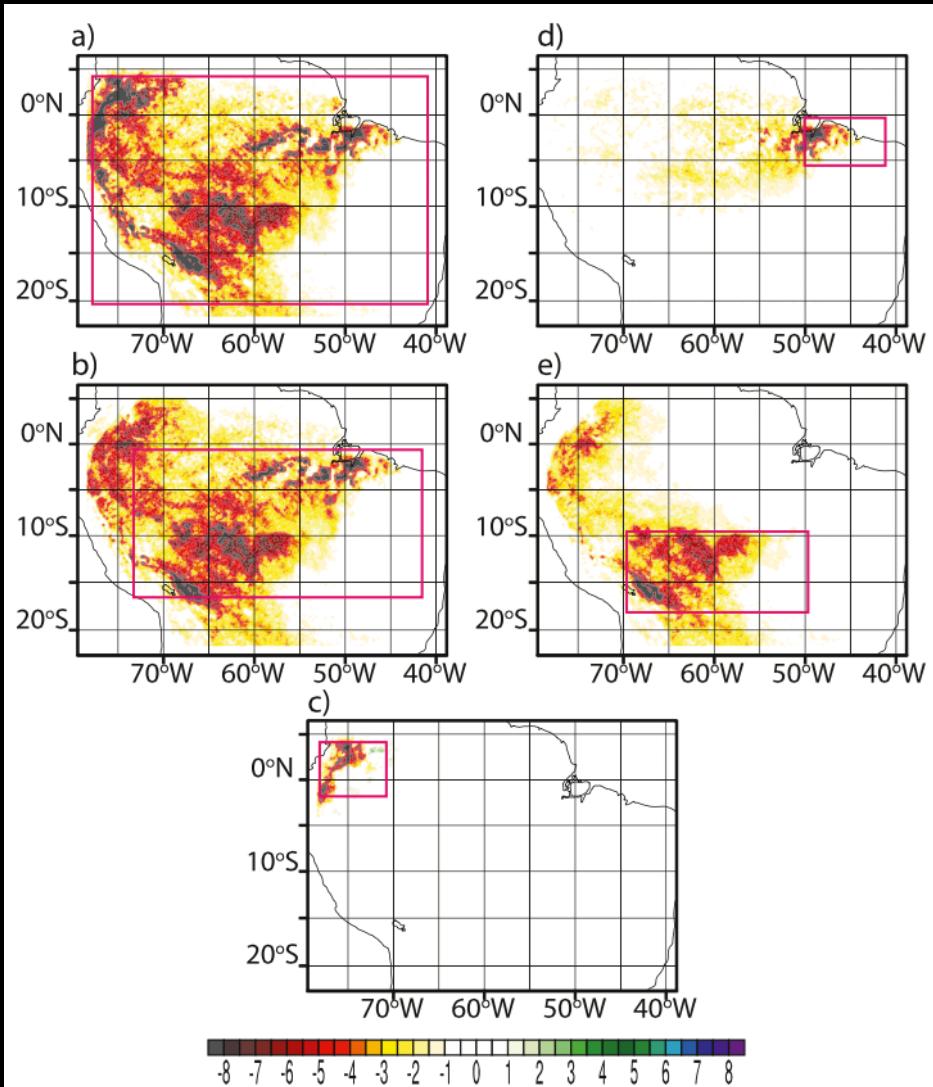
Source Region



Mean forward trajectory
precipitation rate from deforested
points

Moisture Trajectory Analysis





Recycling Ratio Description

Recycling ratio is the fraction of precipitation in a given region that last evapotranspired from the region itself.

Using backtrajectory analysis this is trivially calculated.

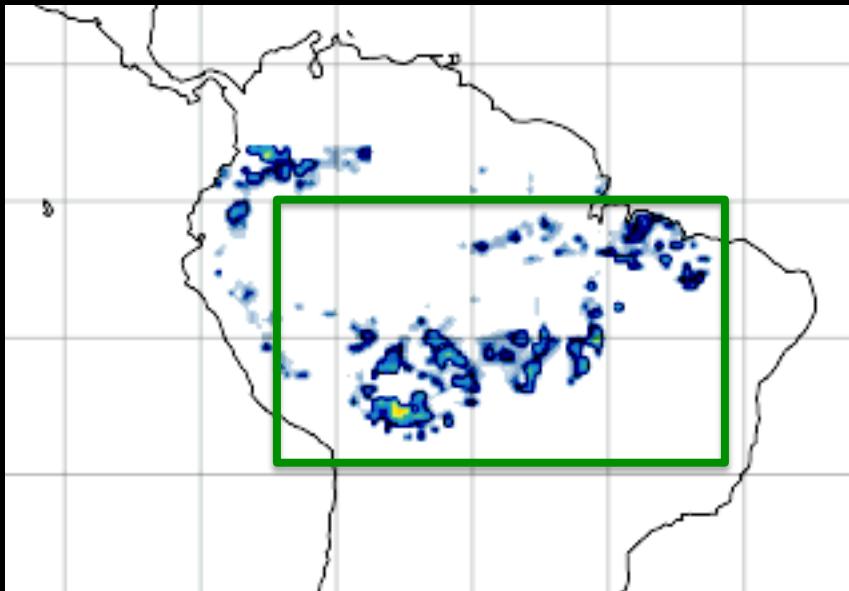
The recycling ratio is an important estimate of land atmosphere coupling.

$$R = \frac{P_{rec}}{P_{tot}}$$

July Regional Recycling

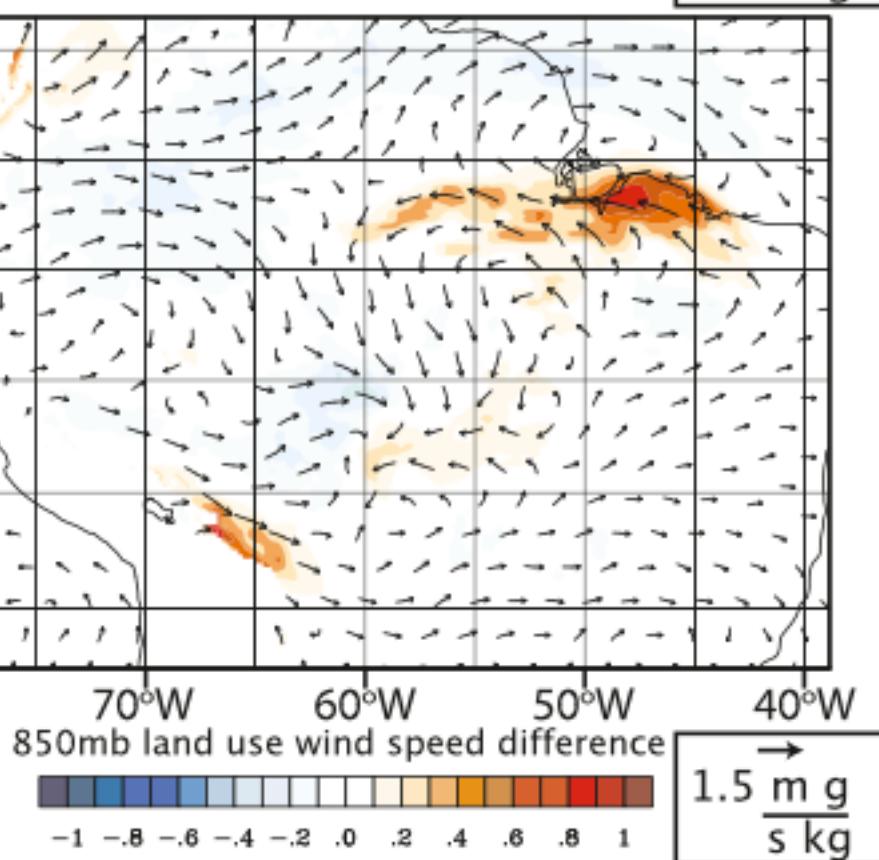
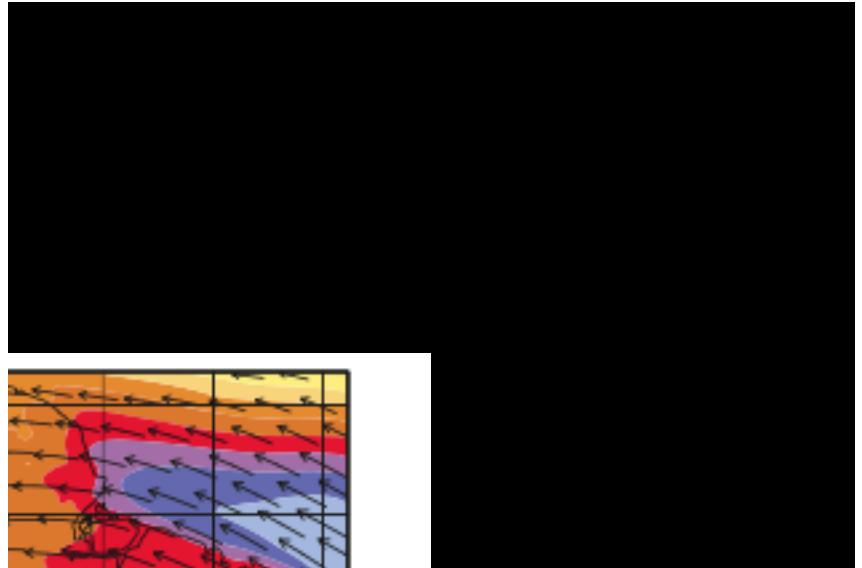
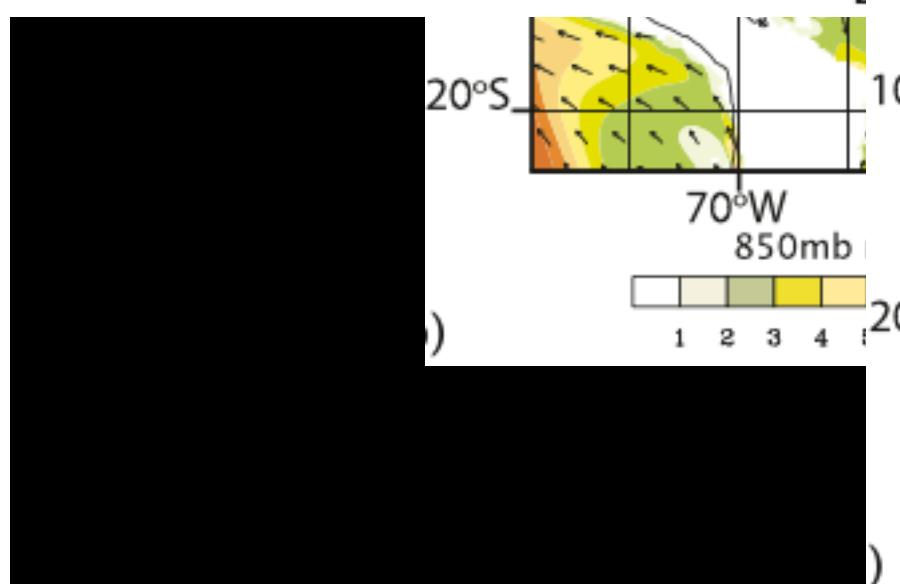
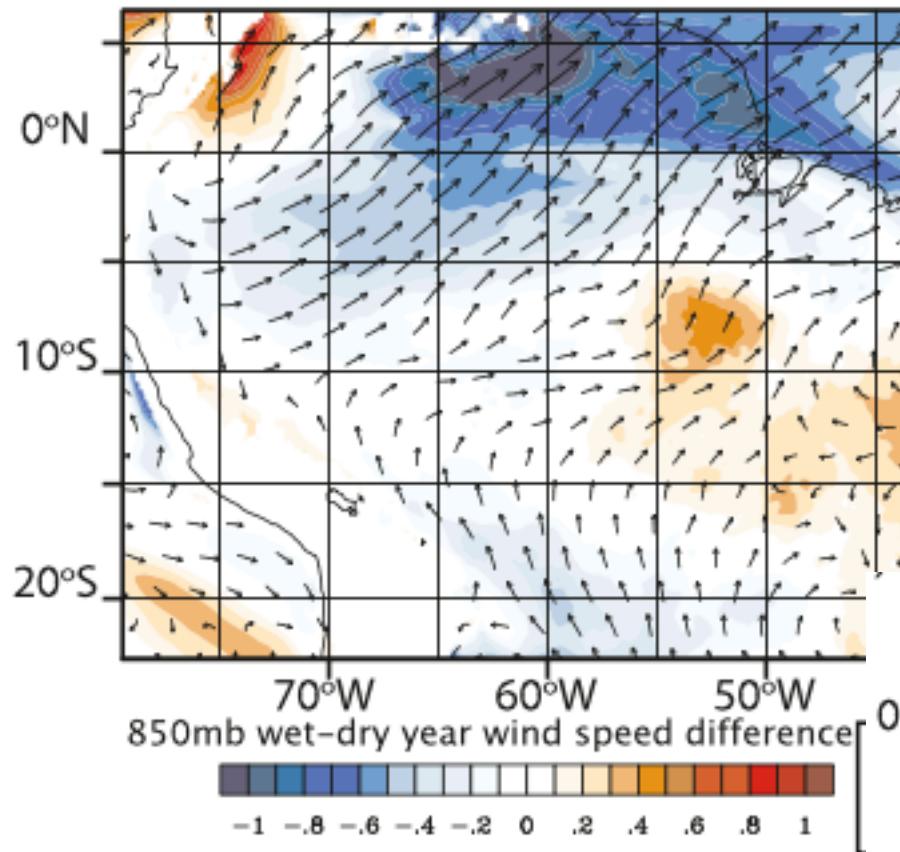
Pluvial Year
Recycling Ratio

.694



Drought Year
Recycling Ratio

.850



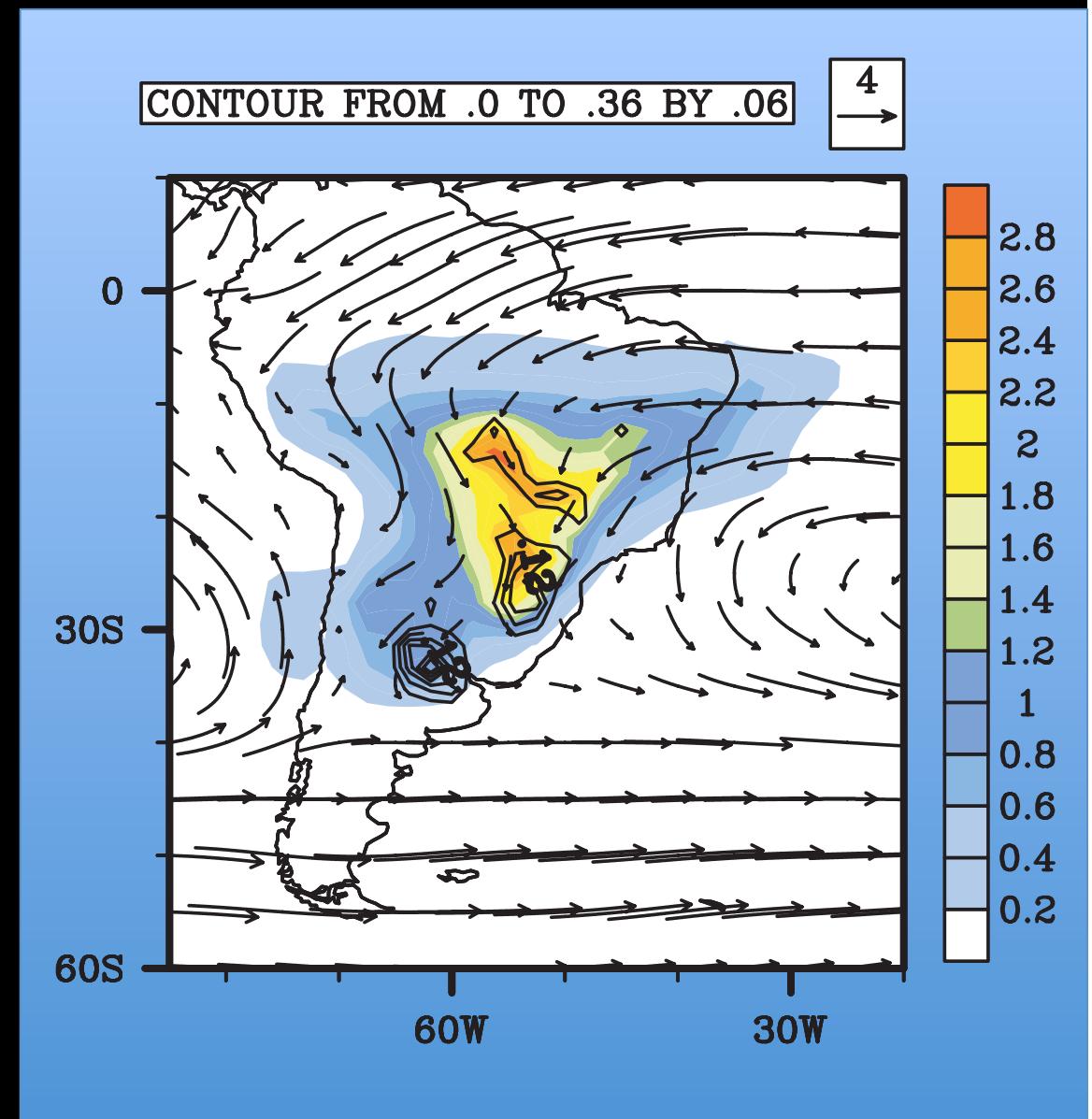
Amazon Rainforest Percent Changes with Deforestation

In nearly every
measure the
impact of
deforestation is
greater during
drought years

% Δ Precipitation Rate
% Δ Sensible Heat Flux
% Δ Latent Heat Flux
% Δ Net Surface Radiation
% Δ Boundary Layer Height
% Δ Rel. Soil Moisture Top Layer
% Δ Rel. Soil Moisture Bot. Layer
% Δ 2m Specific Humidity
% Δ Level of free convection
% Δ Lifting condensation level

		July - September
	Pluvial Years	Drought Years
% Δ Precipitation Rate	-4.99%	-5.93%
% Δ Sensible Heat Flux	+.48%	+4.28%
% Δ Latent Heat Flux	-3.63%	-5.57%
% Δ Net Surface Radiation	-2.41%	-2.70%
% Δ Boundary Layer Height	-.11%	+1.36%
% Δ Rel. Soil Moisture Top Layer	-3.00%	-4.38%
% Δ Rel. Soil Moisture Bot. Layer	+3.50%	+5.09%
% Δ 2m Specific Humidity	-.77%	-1.31%
% Δ Level of free convection	+2.62%	.52%
% Δ Lifting condensation level	+1.29%	+3.94%

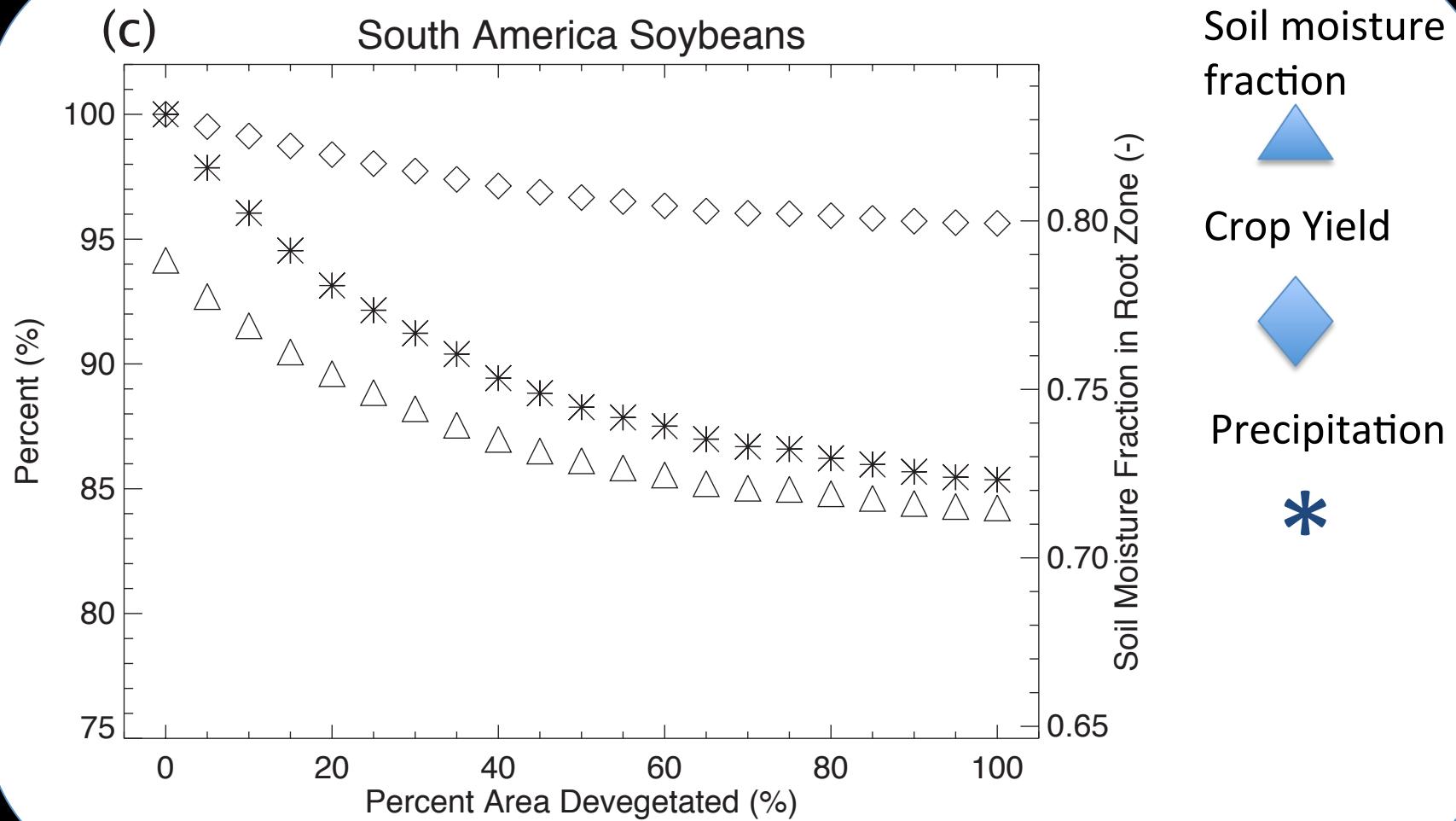
South American Soybeans Evaporative Source



Bagley et al., ERL, 2012

Potential Impact of Land Cover Change on Crop Yield

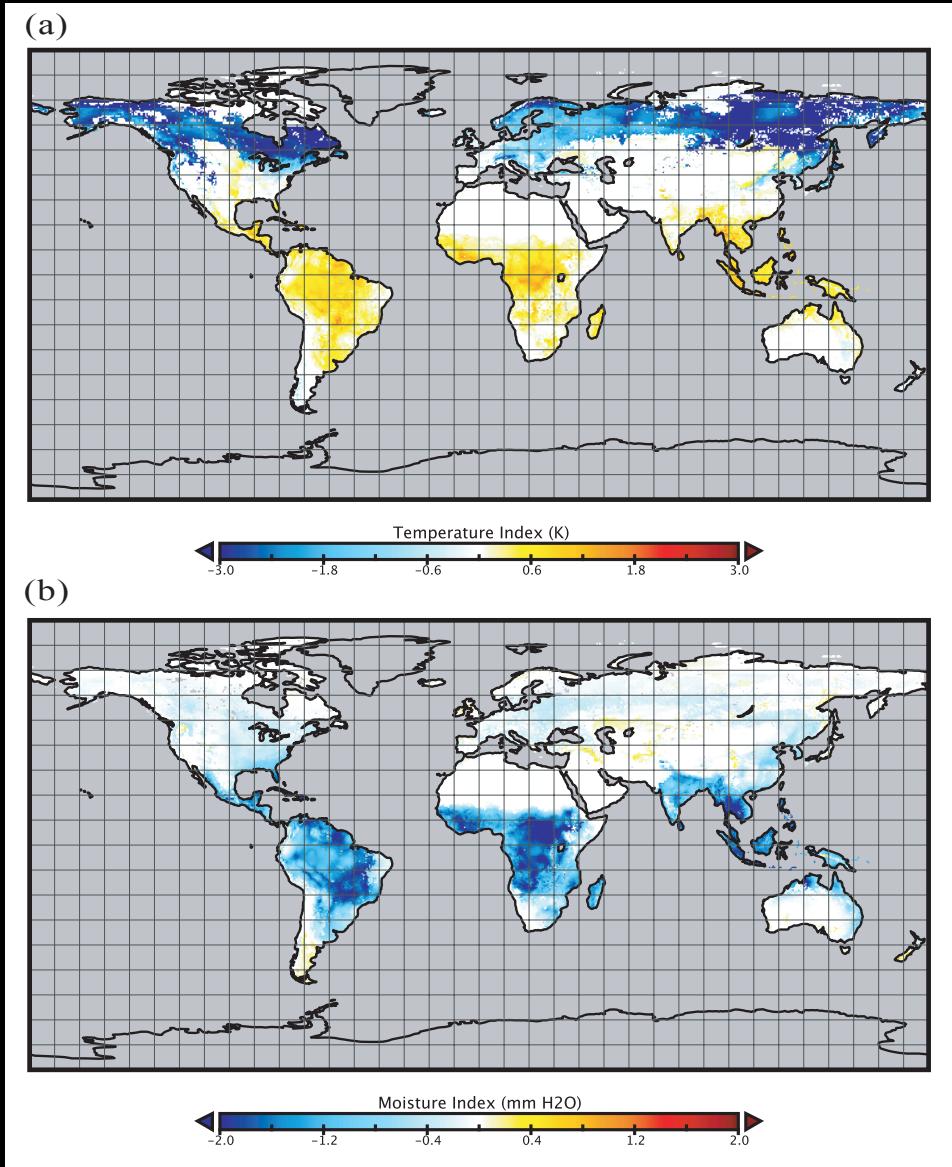
S. America Soybeans



Local regulation of surface climate by vegetation

$$H_reg_index = \Delta H \frac{|\Delta H|}{|\Delta H| + |H_{adv}|}$$

$$Q_reg_index = \Delta Q \frac{|\Delta Q|}{|\Delta Q| + |Q_{adv}|}$$

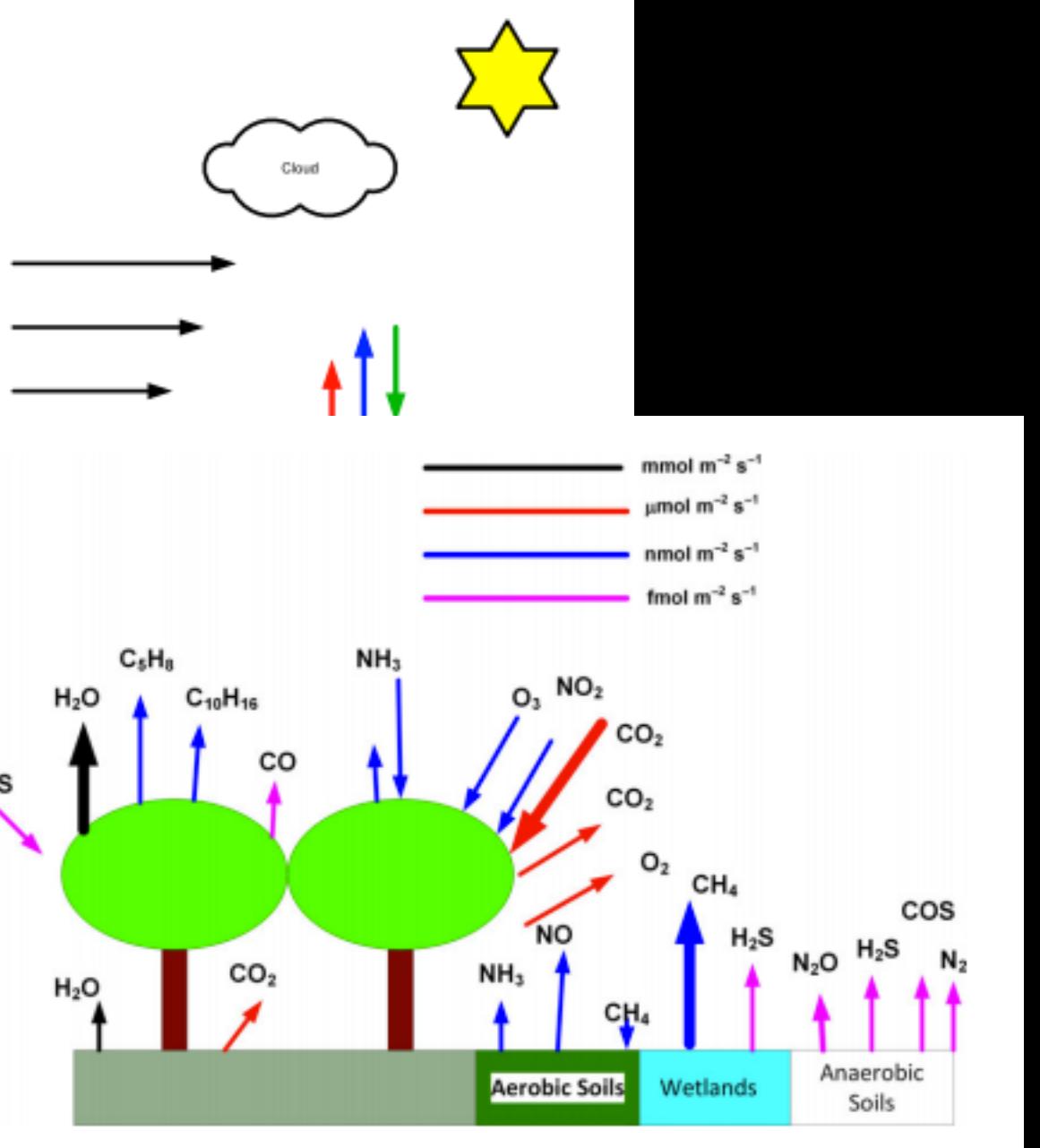
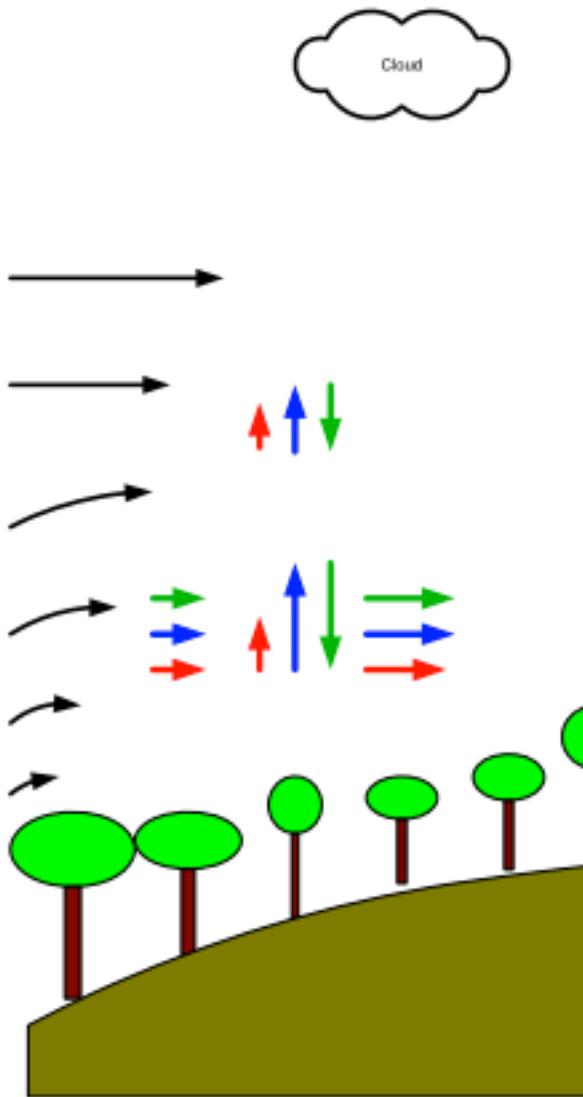


Let's get a little smaller



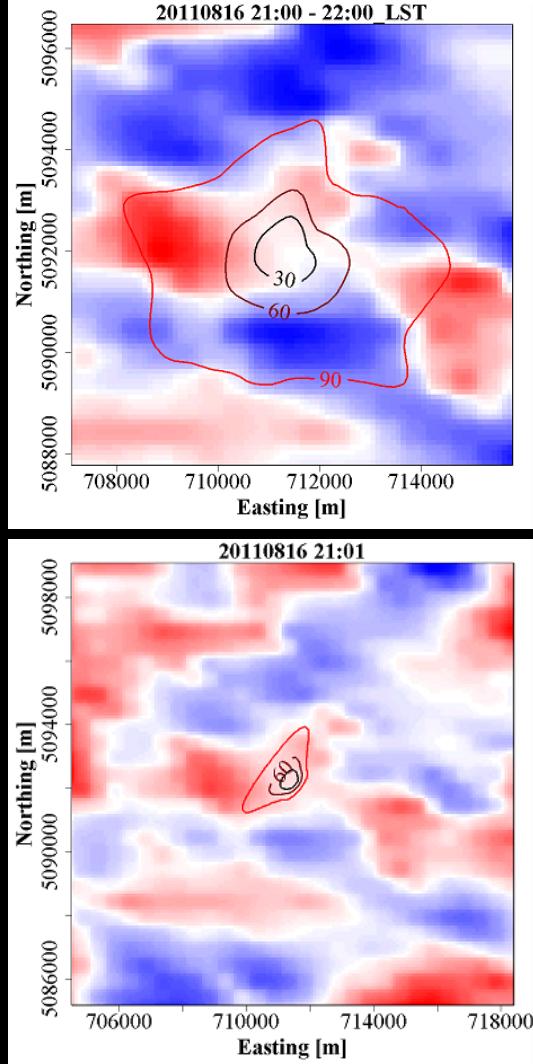
Maybe a few kilometers





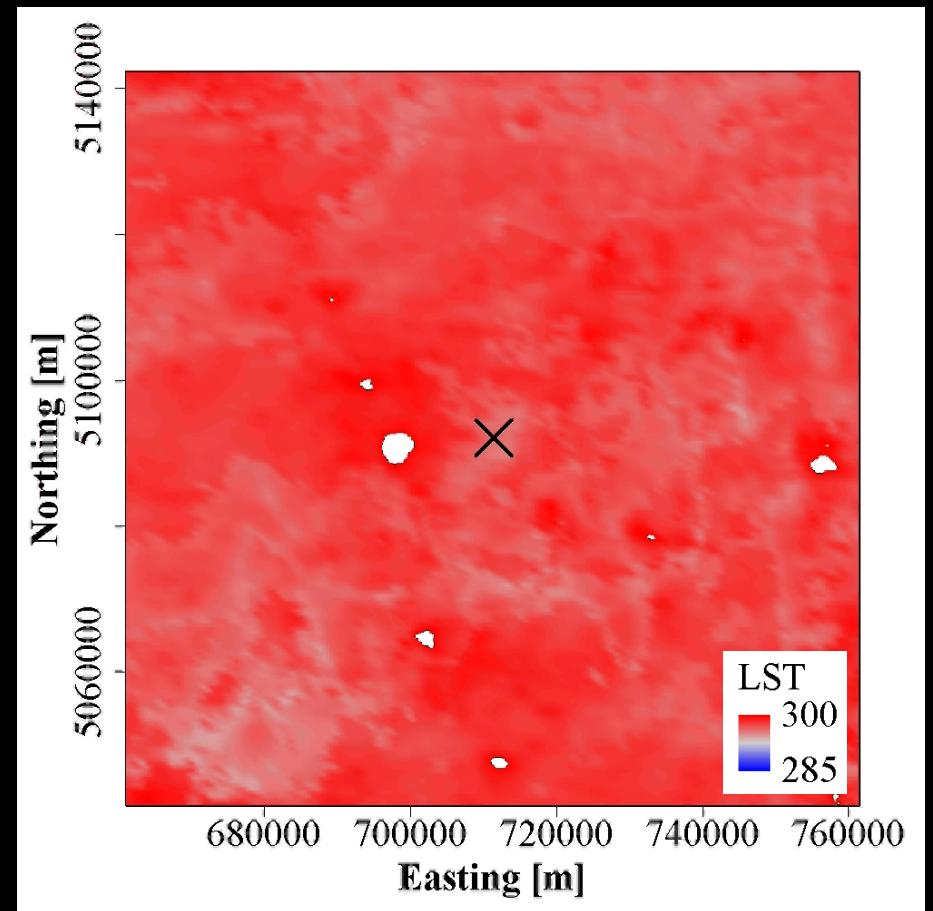
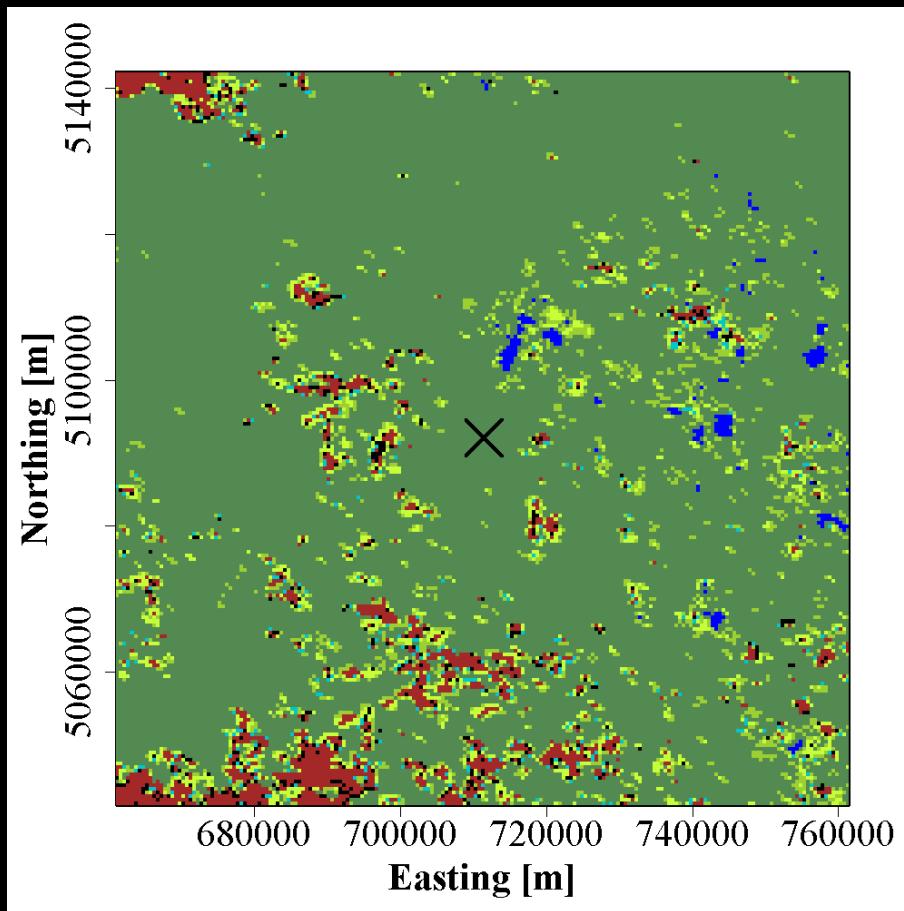
D. Baldocchi

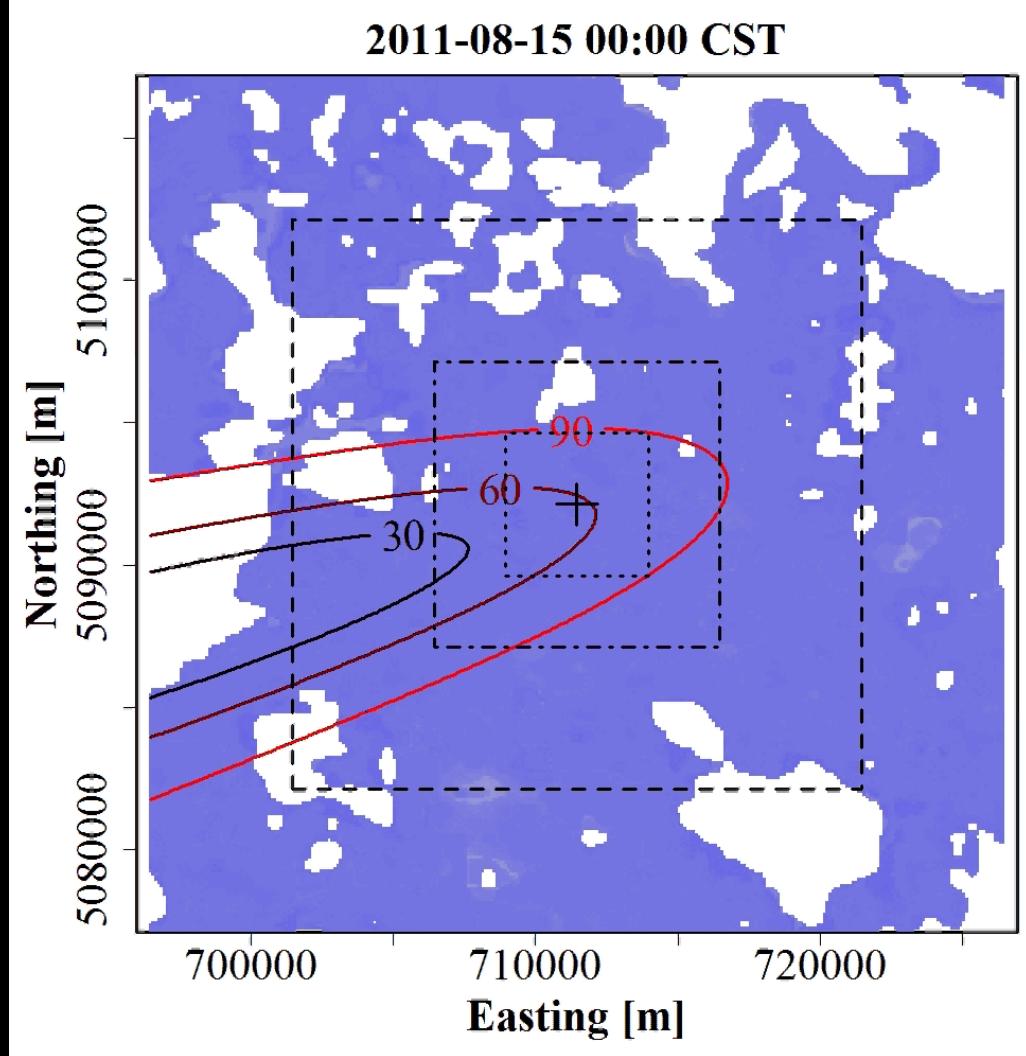
What does the tower flux measurement “see”?



Park Falls/Chequamegon National Forest region

- Relative homogenous...
 - But biophysical properties transient in space and time!





Sensible heat flux [W m⁻²]

2011-08-15 00:00 CST

230
-90

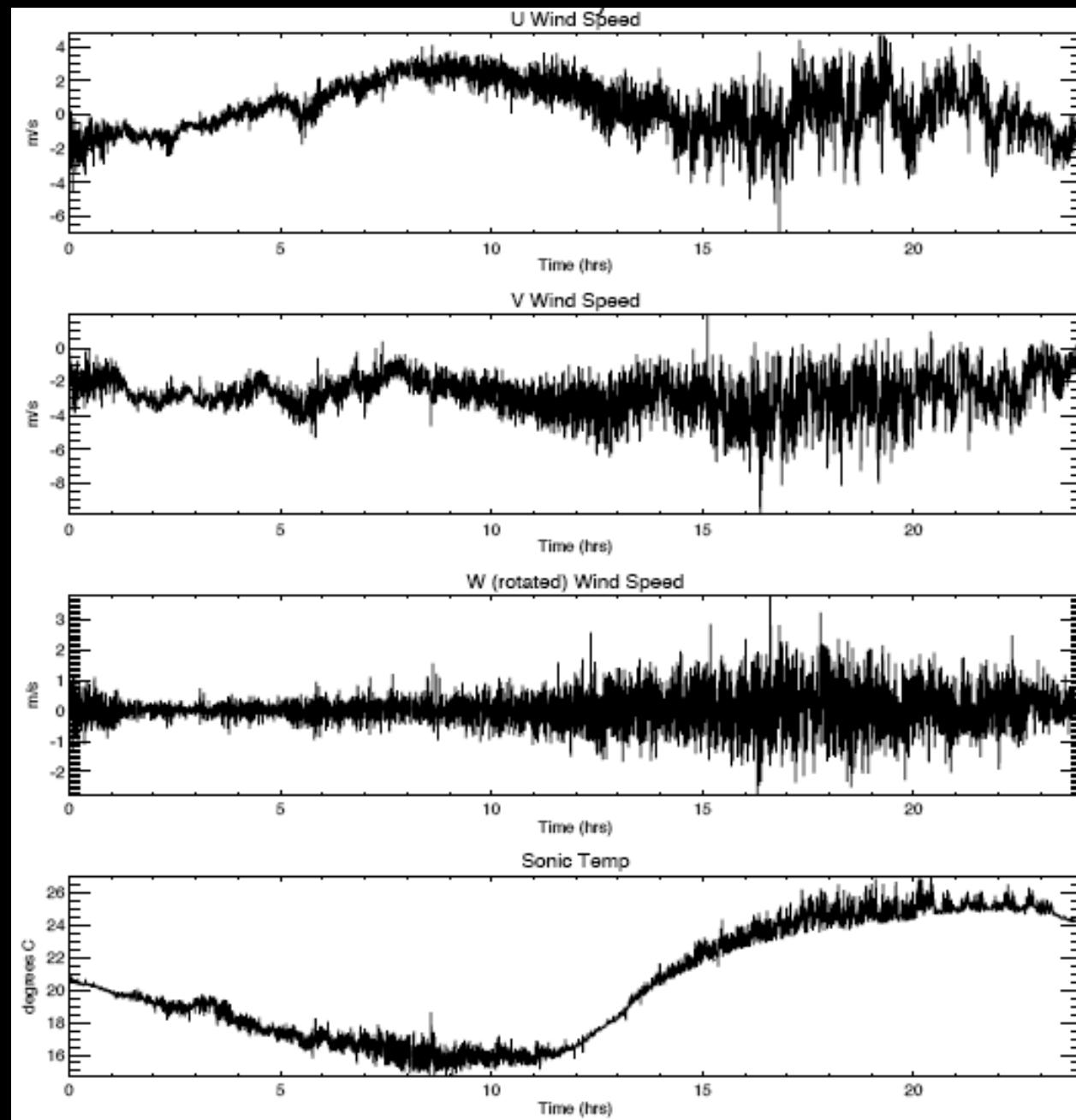
Flux footprint varies in space,
projected fluxes varies in time

Tower represents different surfaces
at different times

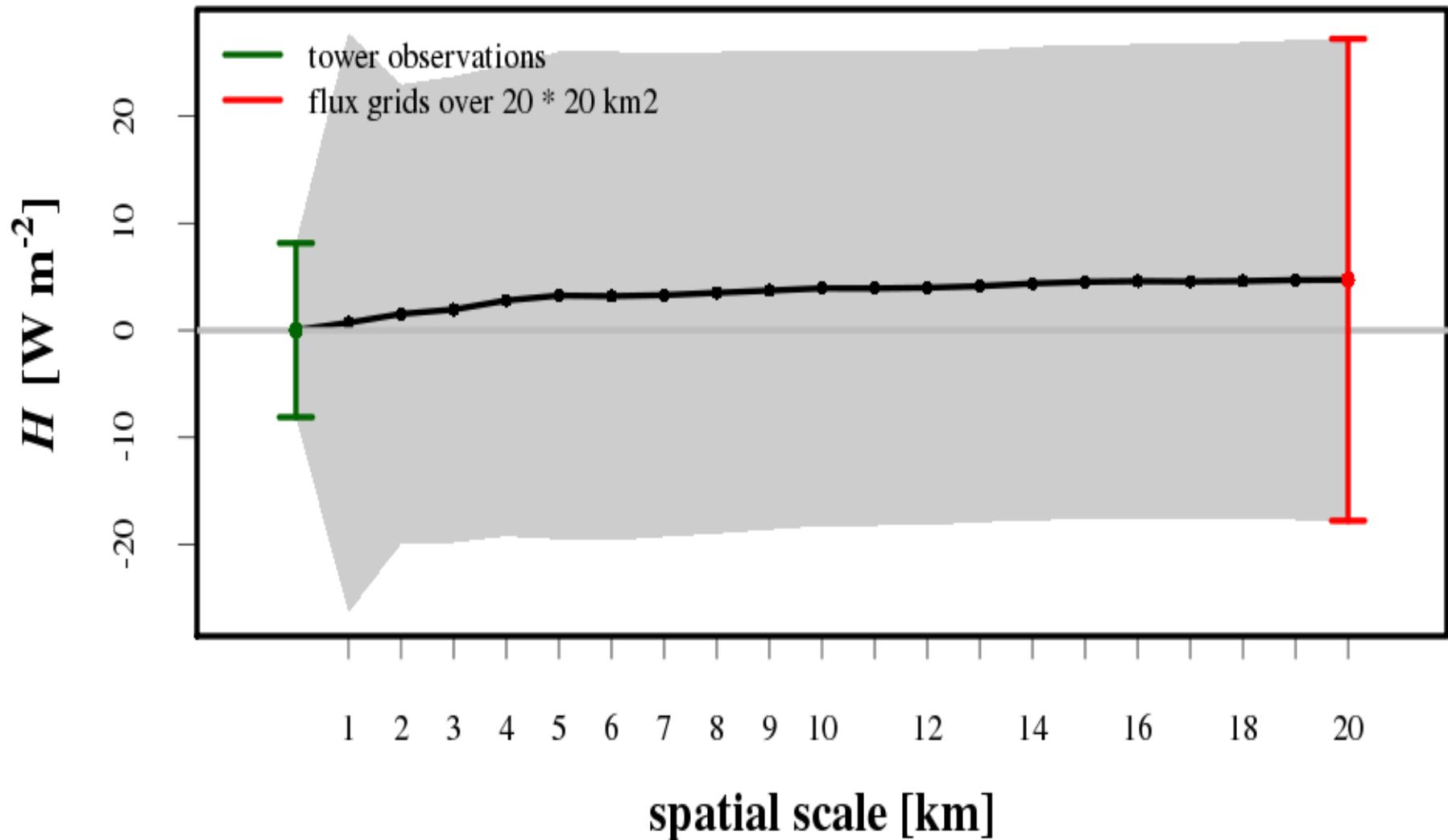
Temporally transient location bias
=“location drift”



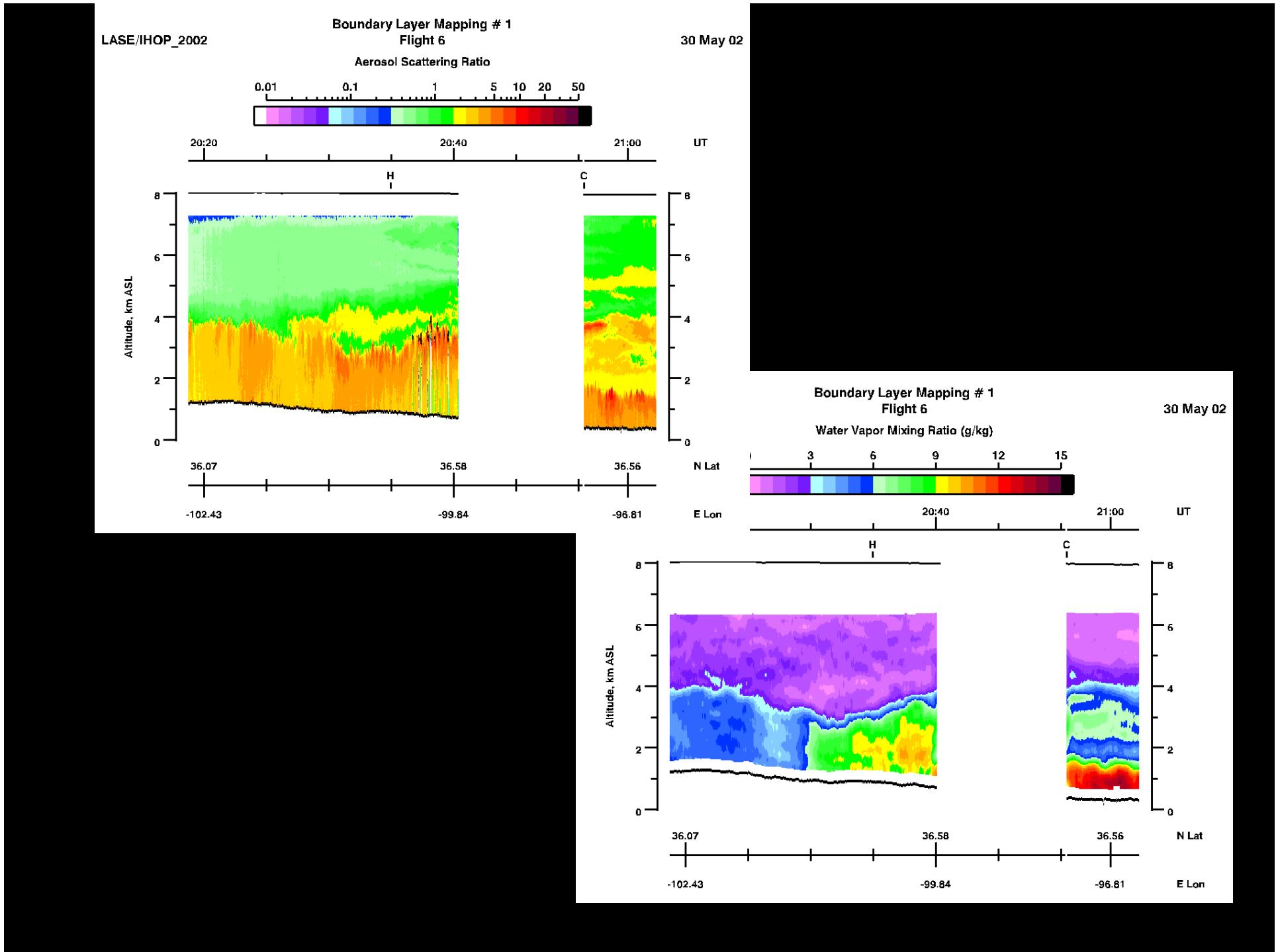
B.D. Cook

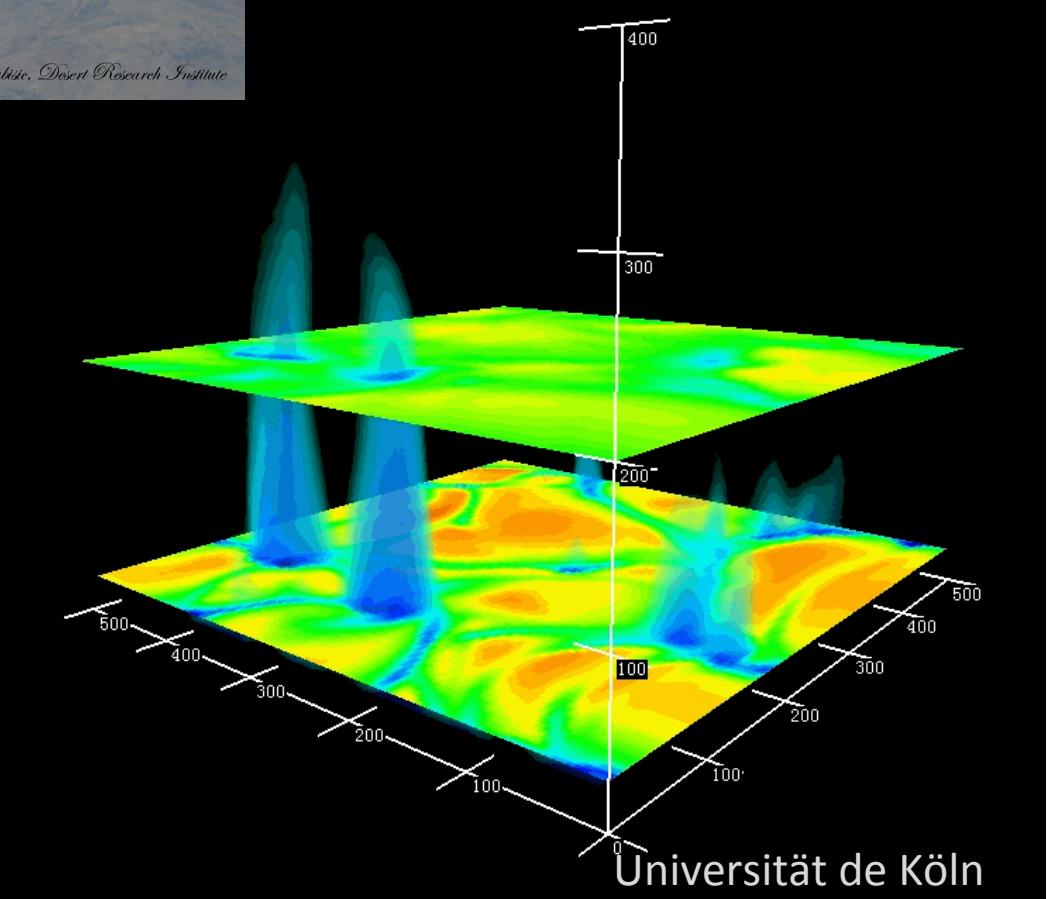


Mean and temporal-spatial variation of flux grids









Large eddy simulation (LES)

- A form of spatial filtering to the full turbulent conservation equations of momentum, mass, heat, and moisture - resolve and subgrid fluxes
- Works because of dissipative and scale-free nature of small-scale shear turbulence in the turbulent atmospheric boundary layer
- Unlike traditional “closure” ensemble-average solutions, resolves energy carrying turbulent motions
- Requires high spatial resolution (meters), and consequently, high temporal resolution (seconds)
- But: Good for testing effect of small scale spatial boundary conditions on atmosphere!

Energy Cascade

- Big whorls have little whorls
- That feed on their velocity,
- And little whorls have lesser whorls
- And so on to viscosity
- (in the molecular sense)
 - -- Lewis F. Richardson, 1922, cf. J Swift

Energy Cascade

$$TKE = \frac{1}{2} \overline{u_i'^2} = \int_0^{\infty} E(k) dk$$

- Can visualize energy spectrum at wavenumber
 $k = \text{eddies of size } 2\pi / k$

Energy Cascade

Garratt

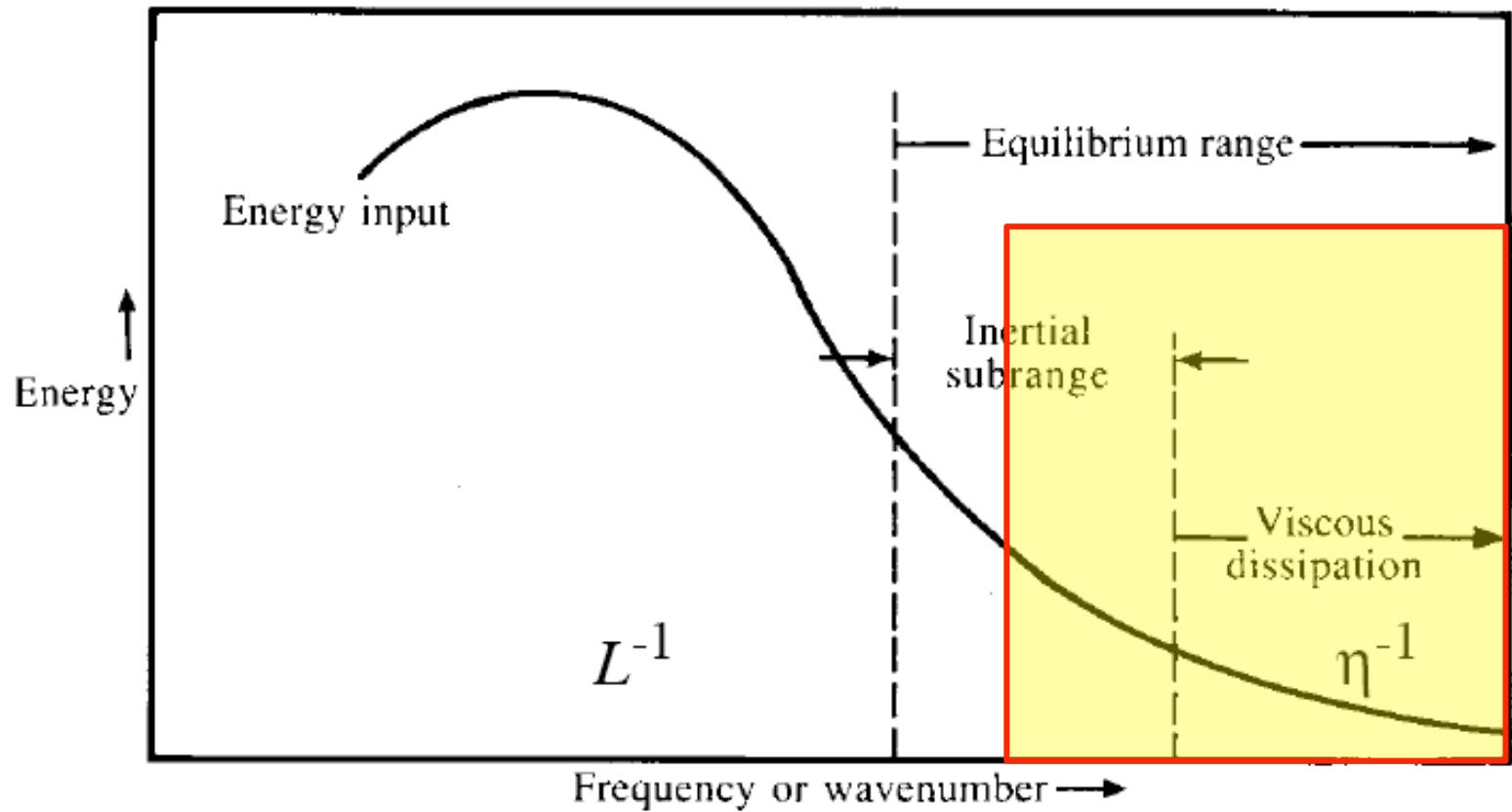


Fig. 2.1 Schematic representation of the energy spectrum of turbulence.

Tower data at 30 – 122 – 396 m to evaluate the simulations

Boundary layer characteristics

$$L = -1.4 \cdot 10^2 \text{ m}$$

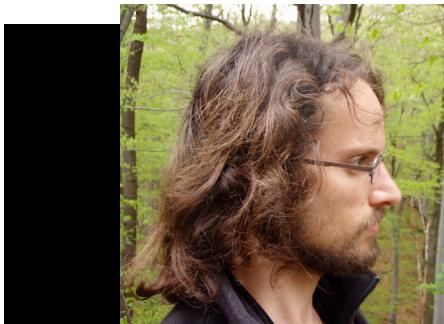
$$z_i = 1.3 \cdot 10^3 \text{ m}$$

$$u_* = 8.2 \cdot 10^{-1} \text{ m/s}$$

Simulation design

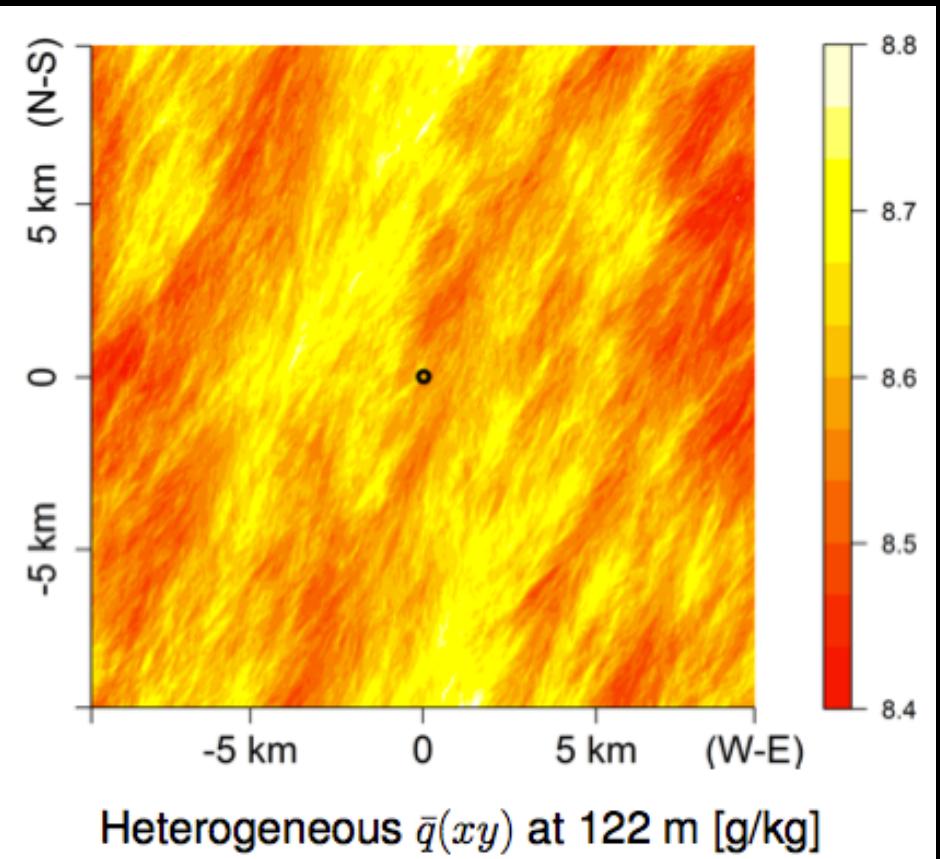
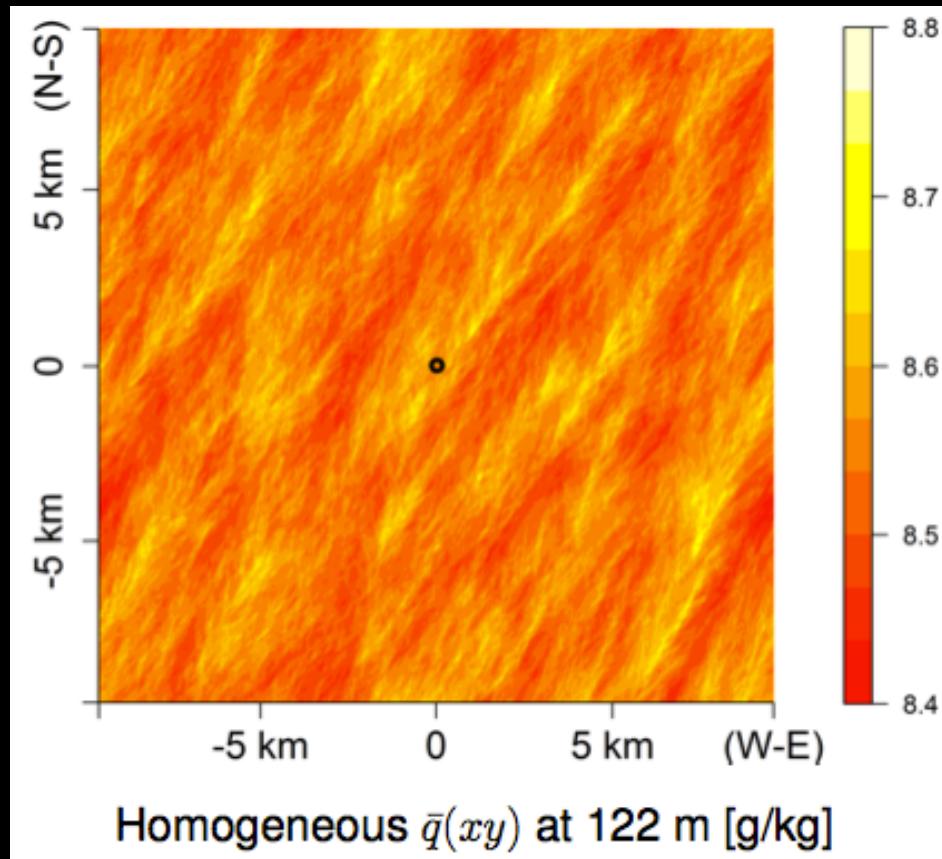
Timestep	0.5 – 1 s
Horizontal grid resolution	10 – 20 m
Gridpoints	$O(10^3 \times 10^3 \times 10^2)$
Vertical grid resolution	5 – 10 m
Horizontal area	100 – 400 km ²

$\sigma_{xy}(\cdot)_{het} - \sigma_{xy}(\cdot)_{hom}$	30 m	122 m	396 m
$T \text{ [K]}$	$+8.7 \cdot 10^{-3}$	$+9.6 \cdot 10^{-3}$	$+1.1 \cdot 10^{-2}$
$q \text{ [g/kg]}$	$+2.2 \cdot 10^{-2}$	$+2.3 \cdot 10^{-2}$	$+2.3 \cdot 10^{-2}$
$w \text{ [m/s]}$	$-5.6 \cdot 10^{-3}$	$-2.2 \cdot 10^{-2}$	$-3.8 \cdot 10^{-2}$

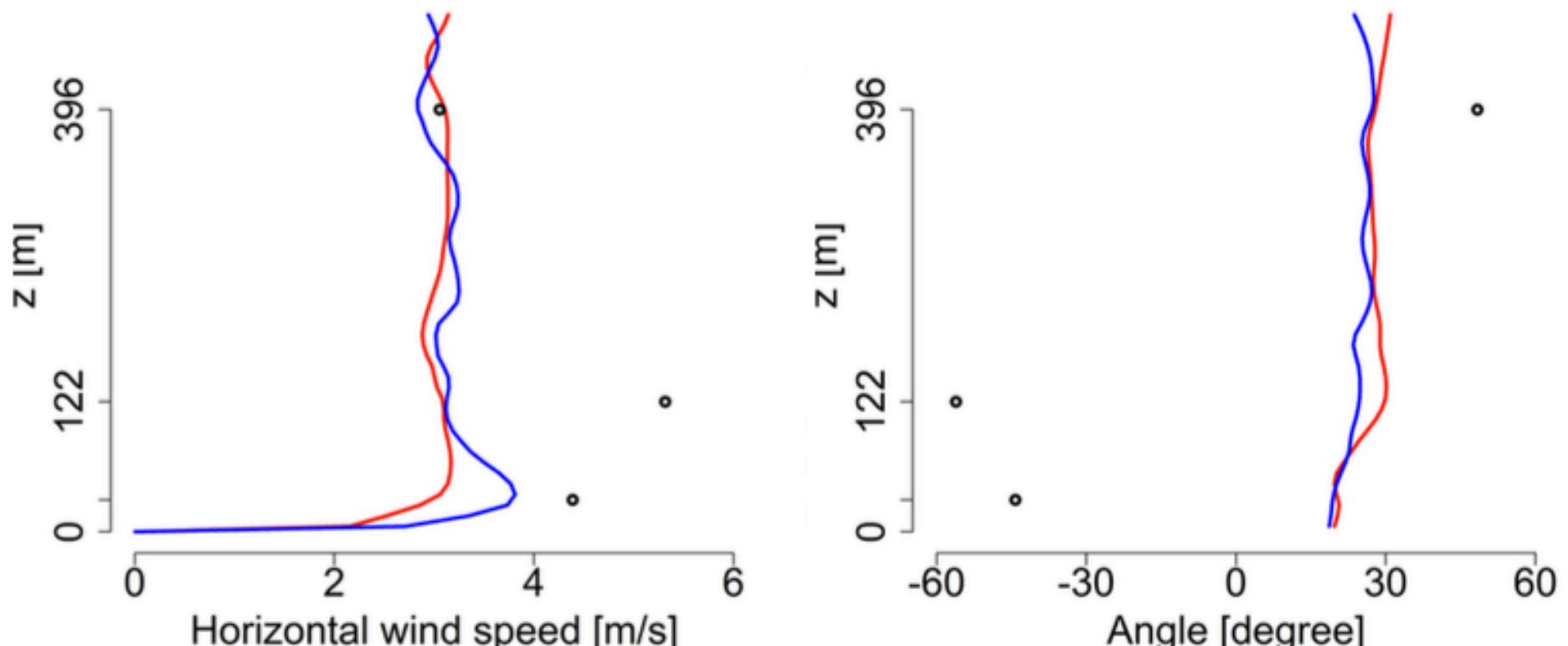


Frederick deRoo (KIT IMK-IFU), TERRENO

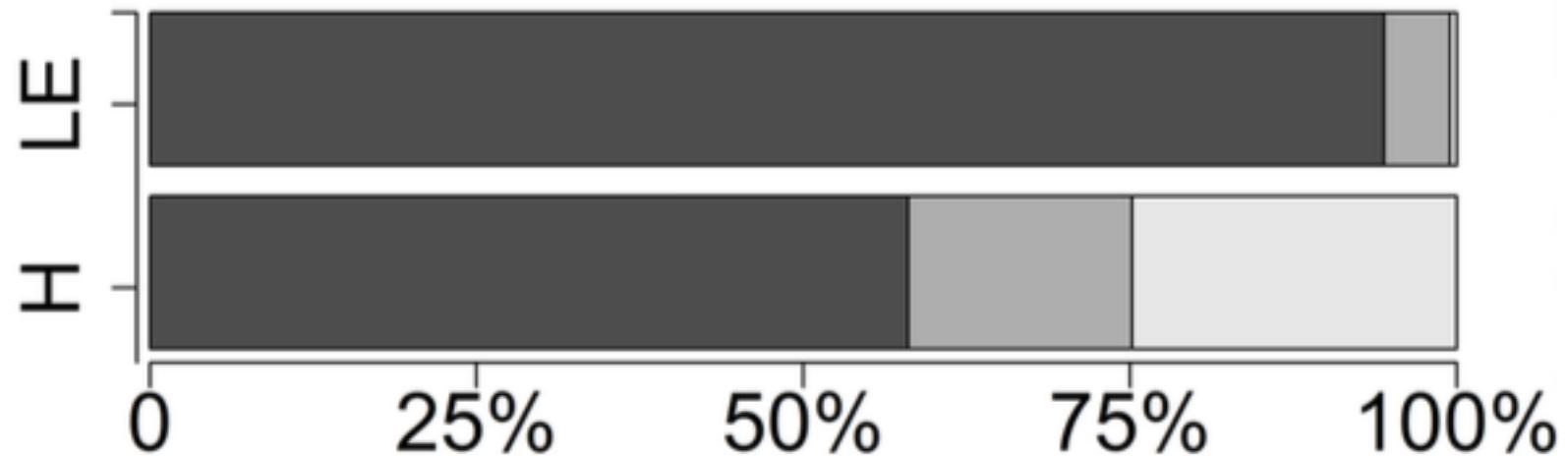
LES simulations around the tall tower show shifts in organized structures with heterogeneity of surface forcing



BUT: A problem...



Eddy fluxes from the homogeneous LES correspond better to the tower data



Virtual EC fluxes as fraction of the tower measurement at 12:00-13:00, 30 m
Darkgray: heterogeneous; Medium-gray: homogeneous

Thank you!



- I hope my examples convinced you that scale is fundamental to understanding ecosystem-atmosphere interactions
- I hope some of the innovations I presented actually solve some of our problems of scale
- None of this can be done without my lab, collaborators, funders, and the opportunity to discuss these with you!