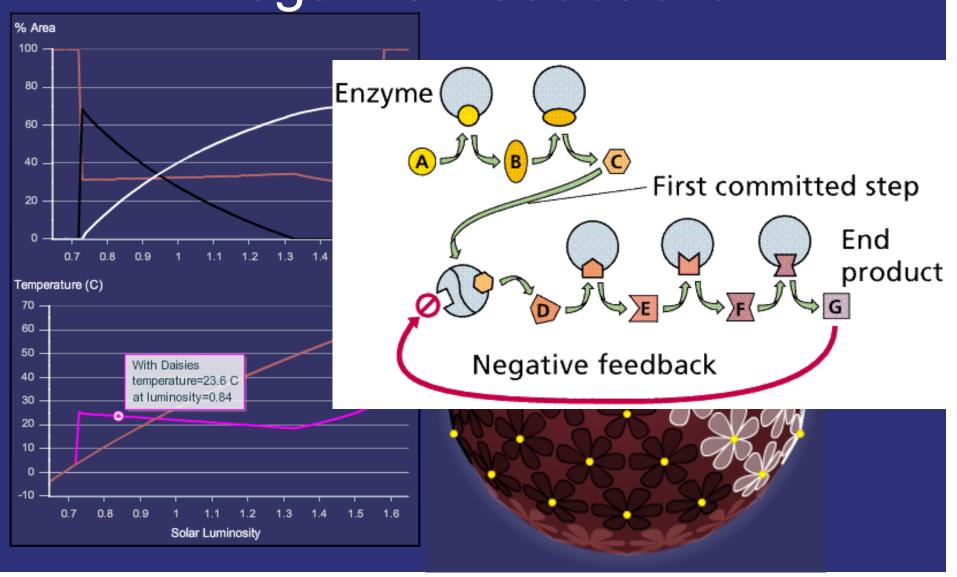
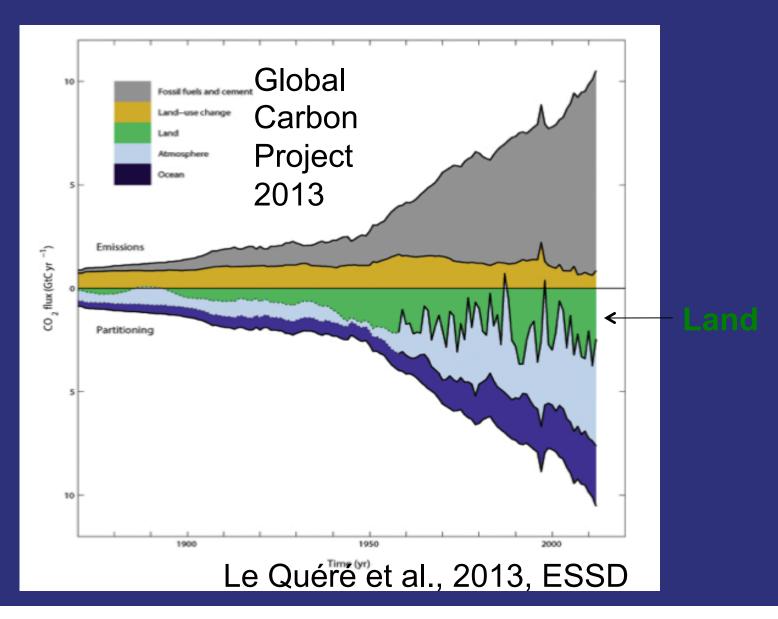


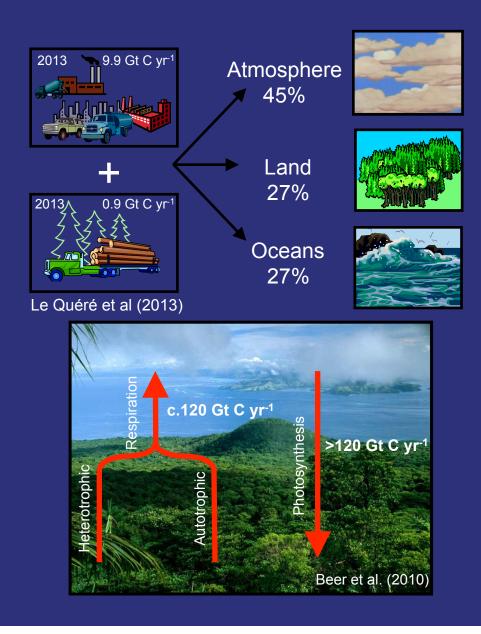
Negative Feedbacks

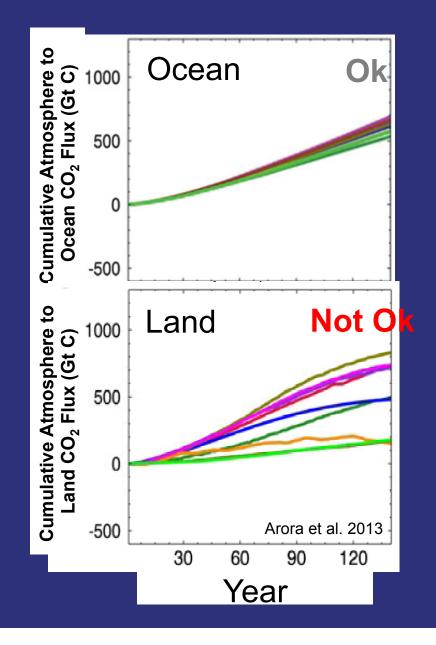


Terrestrial land sink is the largest source of variability in the atmospheric CO₂ growth rate

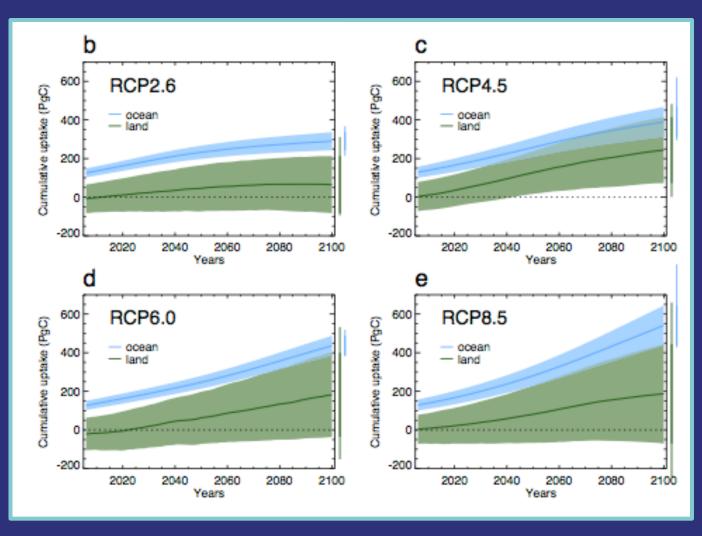


Terrestrial Biosphere CO₂ Flux Dominates Carbon Cycle Prediction Uncertainty





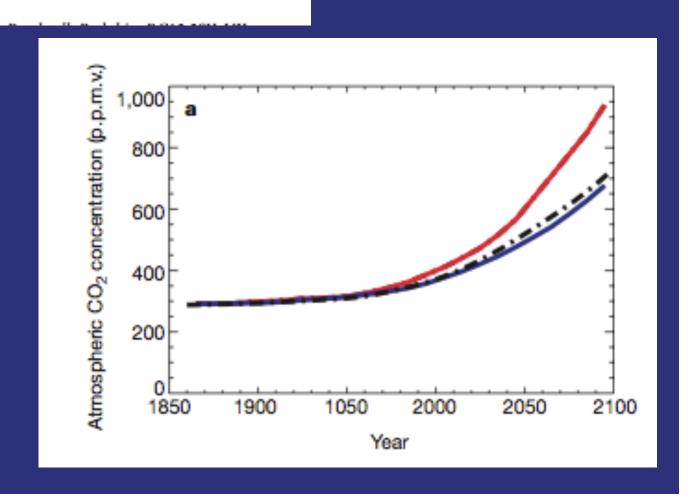
Terrestrial carbon cycle feedback is a leading order uncertainty for climate simulation



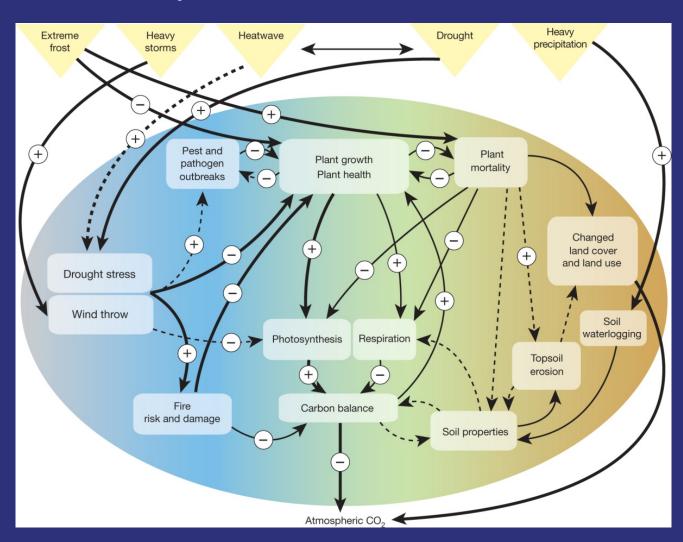
IPCC AR5 WG1 CH6 (draft)

Acceleration of global warming due to carbon-cycle feedbacks in a coupled climate model

Peter M. Cox*, Richard A. Betts*, Chris D. Jones*, Steven A. Spall* & Ian J. Totterdell†



Processes and feedbacks triggered by extreme climate events?



M Reichstein et al. Nature 500, 287-295 (2013) doi:10.1038/nature12350

No one trusts a model except the one who wrote it; everyone trusts an observation except the one who made it – Harlow Shapley (by way of Matt Disney)

Willow Creek - NetCam SC IR - Thu Sep 20 11:31:17 2012 Temperature: 36.0 °C internal, 9.0 °C outside RH: 0%, Pressure: 944.0 millibars Exposure: 400

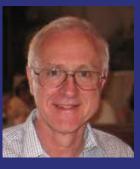
PROLOGUE



Eddy covariance is mature technology

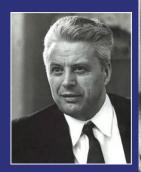


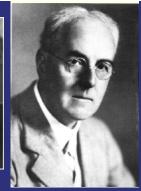






History







 1880-1920s Turublence theory (Reynolds, Prandtl, Richardson, Taylor)

• 1940s-1950s Surface-layer theory (Monin-Obhukov, Kolmogorov), development of fast sensors for anemometry





1970s CO₂ fluxes (Desjardins, Leuning)

• 1980s Infrared gas analyzers (Verma, Anderson, Valentini)

 1990s First long-term regional CO₂ flux networks (Wofsy, Baldocchi, Goulden, Law, Aubinet)

 2000s Global syntheses (FLUXNET, Falge, Papale, Reichstein)

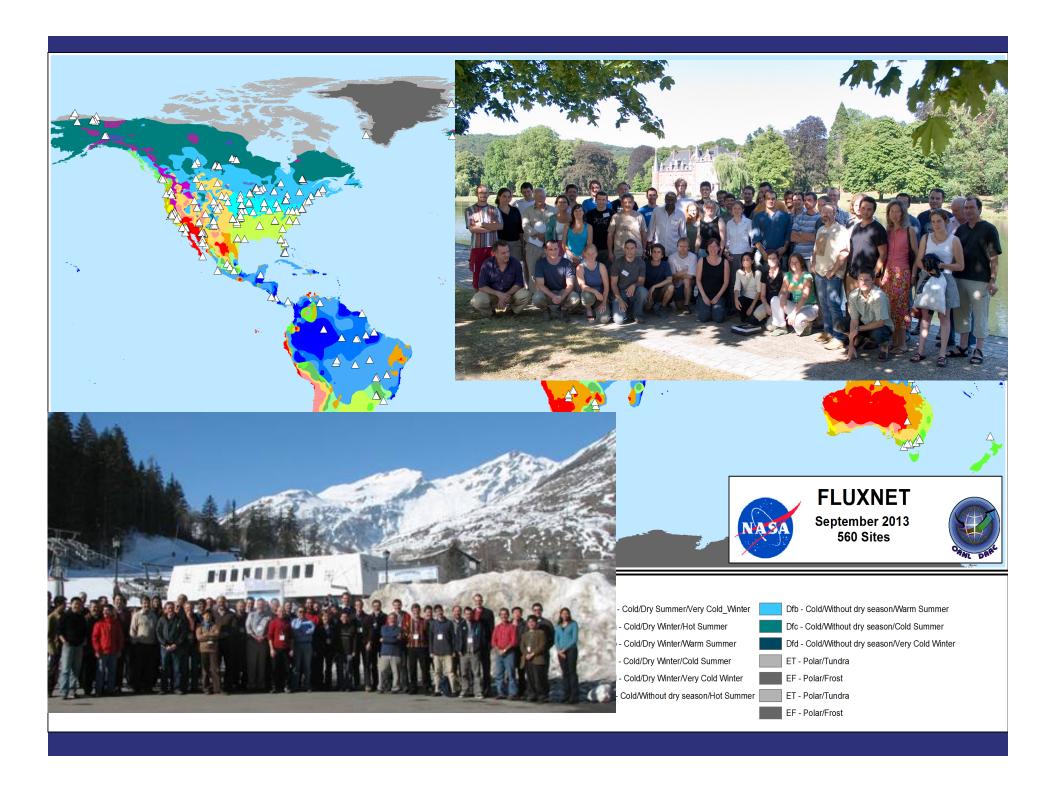
• 2010s Model-data integration, development of operational measurements (NEON, ICOS, you?)

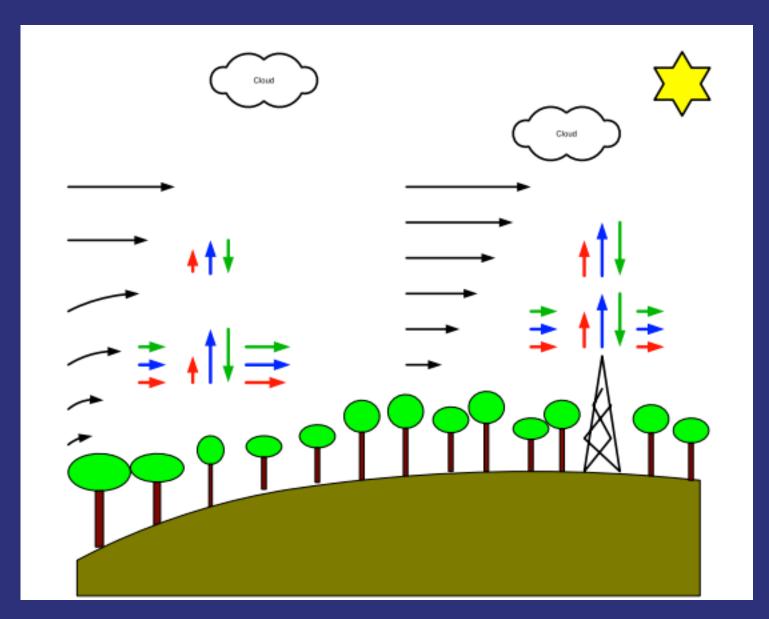












D. Baldocchi

$$\frac{\overline{D\widetilde{C}}}{Dt} = 0 \longrightarrow \frac{\overline{\partial \widetilde{C}}}{\partial t} + \widetilde{U}_{j} \frac{\partial \widetilde{C}}{\partial x_{j}} = 0$$

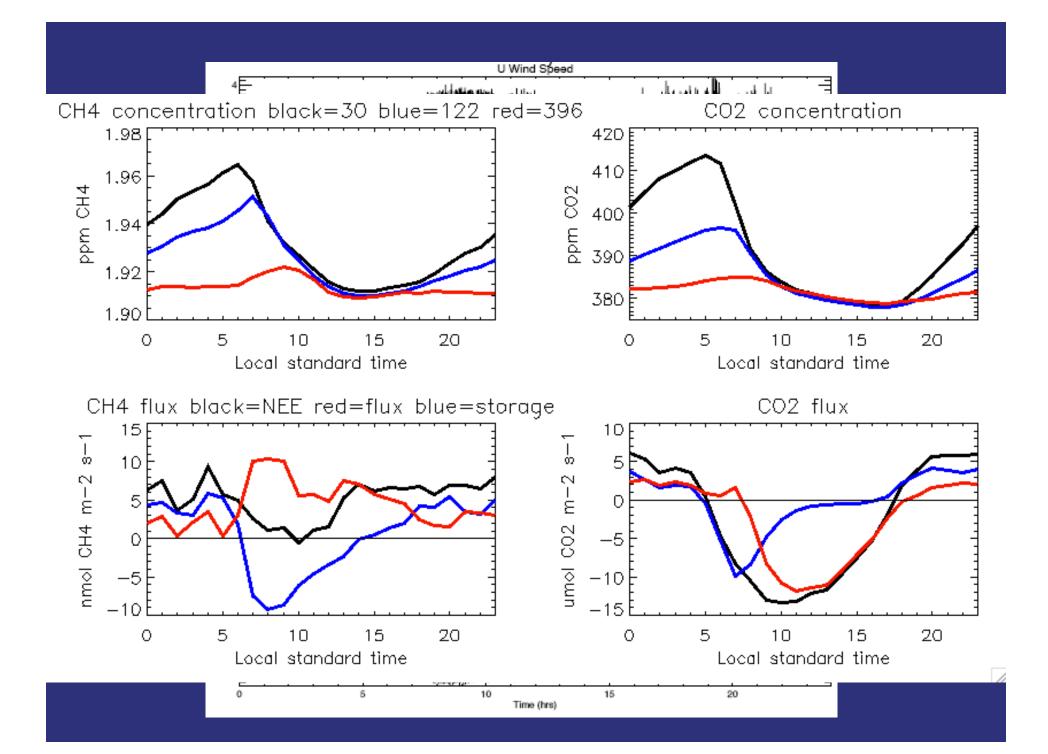
STORAGE TURBULENT-FLUX

NEE
$$\equiv \int_{0}^{z_{r}} s dz + (\overline{w'c'})|_{z=0}$$

$$= \int_{0}^{z_{r}} \frac{\partial \bar{c}}{\partial t} dz + (\overline{w'c'})_{r} + \bar{w}_{r}(\bar{c}_{r} - \langle \bar{c} \rangle)$$
(4)

$$\frac{\partial \overline{C}}{\partial t} + \frac{\partial \overline{C'}}{\partial t} + \overline{U_j} \frac{\partial \overline{C}}{\partial x_j} + \overline{U_j'} \frac{\partial \overline{C}}{\partial x_j} + \overline{U_j'} \frac{\partial \overline{C'}}{\partial x_j} + \overline{U_j'} \frac{\partial \overline{C'}}{\partial x_j} + \overline{U_j'} \frac{\partial \overline{C'}}{\partial x_j} = 0$$

$$\frac{\partial \overline{C}}{\partial t} + \overline{U_j} \frac{\partial \overline{C}}{\partial x_j} + \overline{u_j'} \frac{\partial \overline{C'}}{\partial x_j} = \frac{\partial \overline{C}}{\partial t} + \frac{\partial \overline{U_j} \overline{C}}{\partial x_j} + \frac{\partial \overline{u_j'} \overline{C'}}{\partial x_j} = 0$$



Who we are

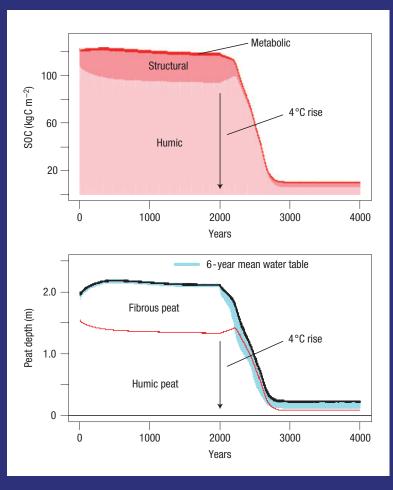




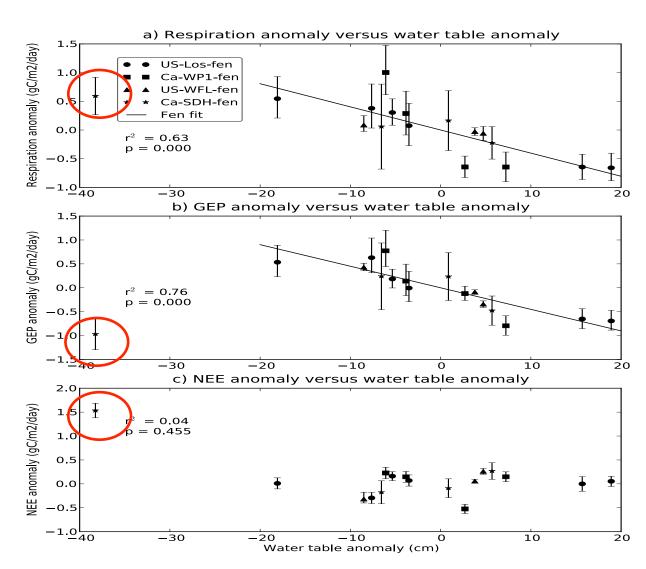


Peatland carbon is vulnerable to climate and hydrological change

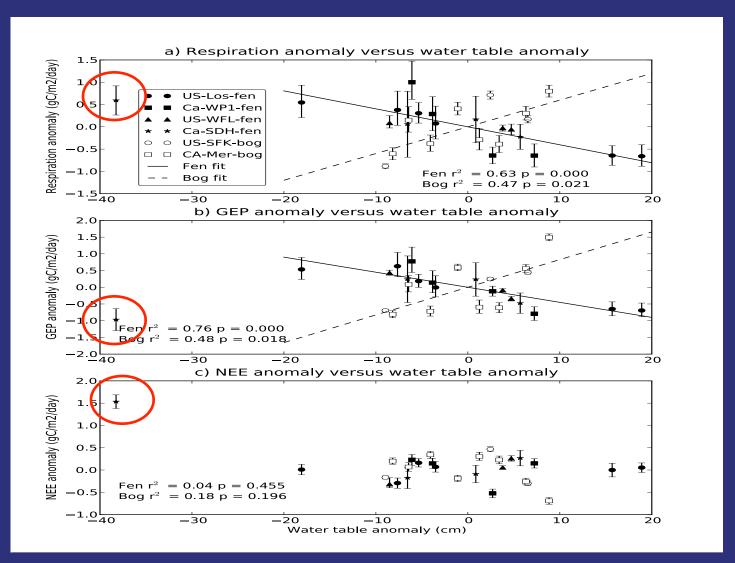
- Peat carbon is preserved by cool temperatures and flooded conditions
- Warming and drying can disrupt the process and lead to carbon loss



Hydrology does not drive NEE in four fens



Same for bogs, but in a different way

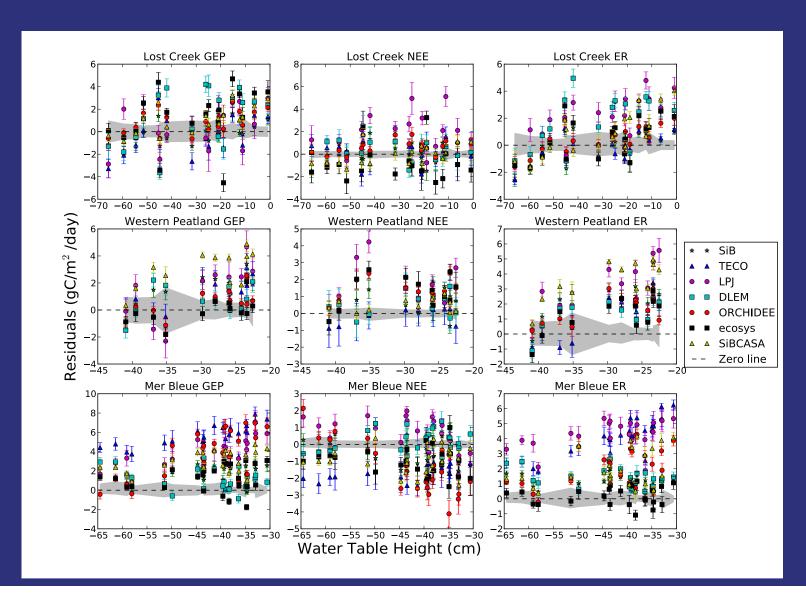


How well did models simulate peatland processes?

Model name	Temporal resolution	Soil layers	Soil C pools	N cycle	Max soil moisture
DLEM	Daily	2	3	Yes	Saturation
Ecosys	Hourly	8	9	Yes	Saturation (with water table)
LPJ	Daily	2	2	No	Field capacity
ORCHIDEE	30-min	2	8	No	Field capacity
SiB	30-min	10	None	No	Saturation
SiBCASA	30-min	25	9	No	Saturation
TECO	30-min	10	5	No	Saturation

Sulman et al., JGR-G, 2011

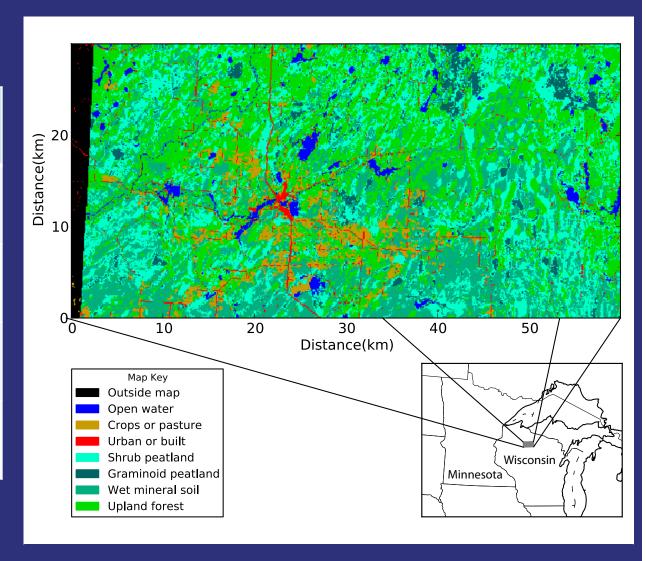
Monthly residuals were correlated with observed water table



Maybe longer term?

Ecoregion	Active area fraction
Upland	38%
Mineral wetland	27%
Shrub peat	29%
Graminoid peat	5%

LANDIS-II model



Sulman et al., Ecosystems, 2013

Water table effects on carbon Peatlands: balance

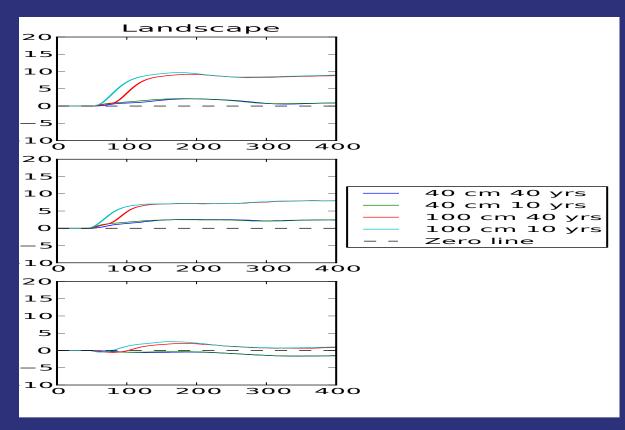
- 100 cm declines:
 - Short term: C gain
 - Long term: C loss
- 40 cm declines
 - Short term: C neutral
 - Long term: C loss

Mineral wetlands:

C gain for both

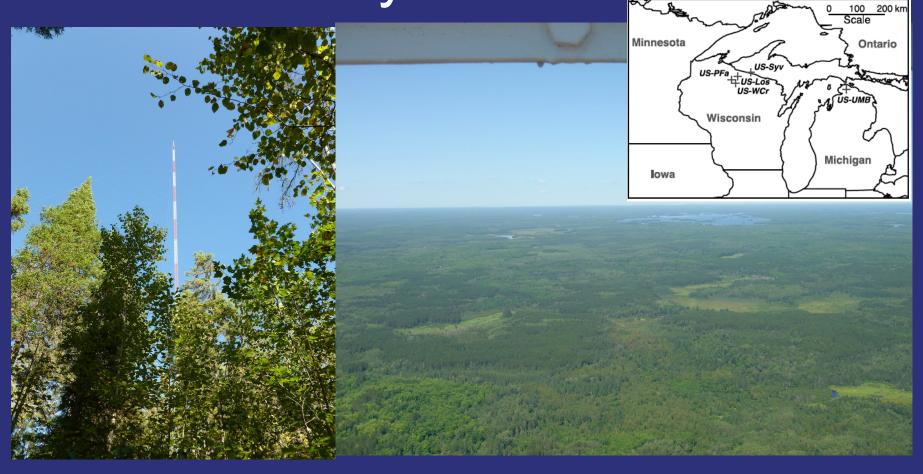
Whole landscape

- Short-term: C increase
- Long-term: C steady
- Time scale of decline made little difference

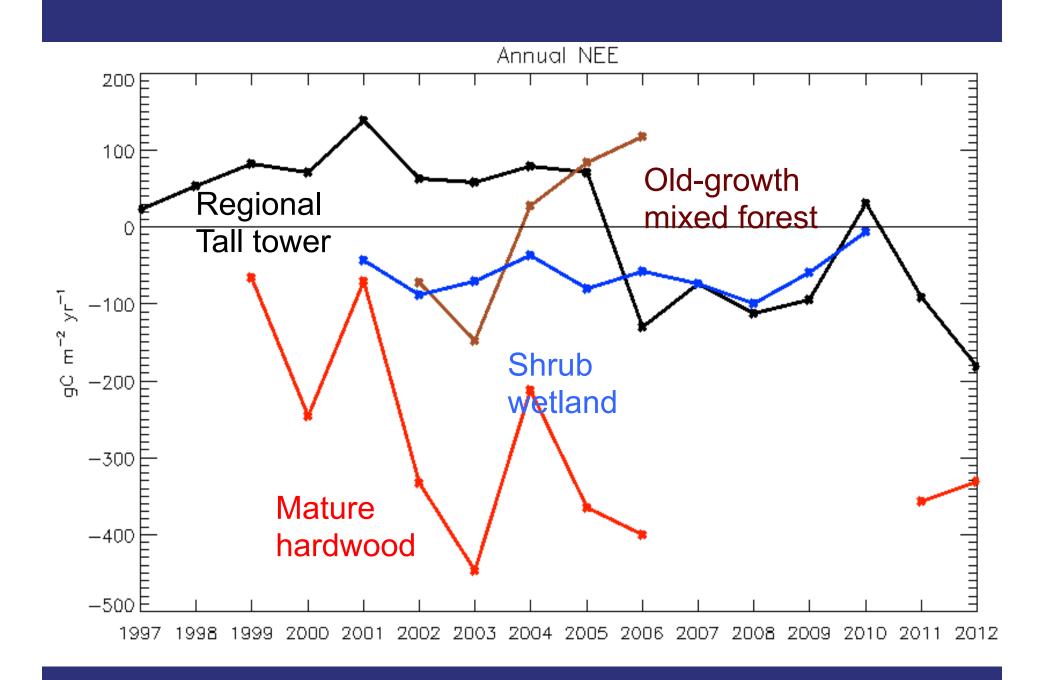


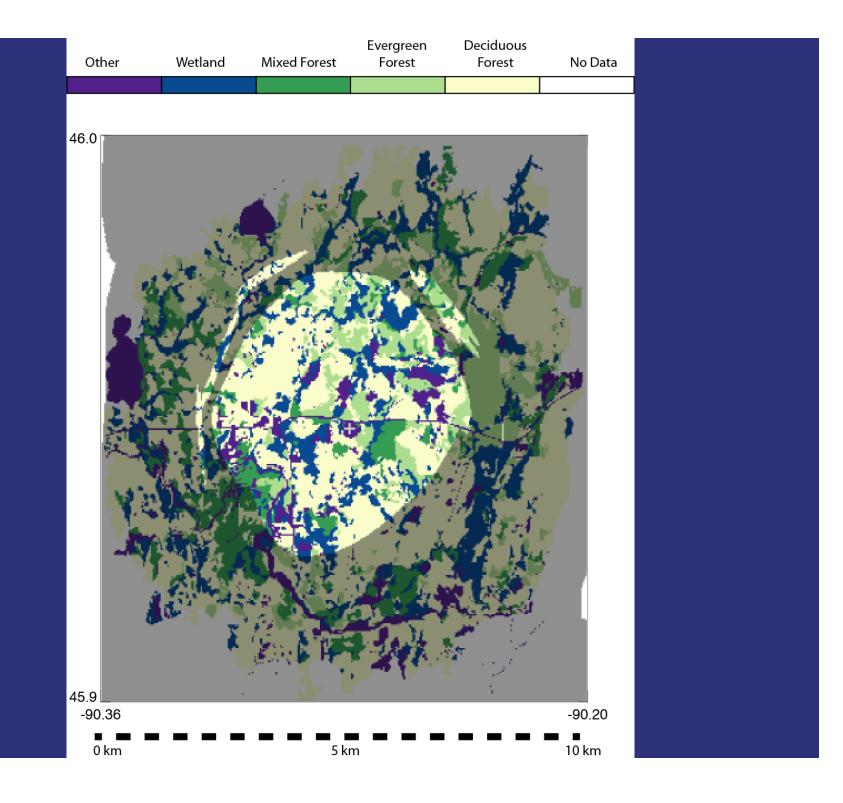
Net change from control run for shallow peat simulations: Different water table scenarios

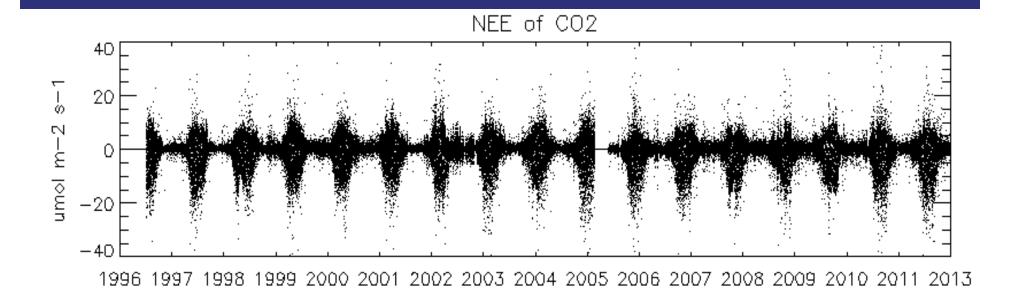
A very tall tower!

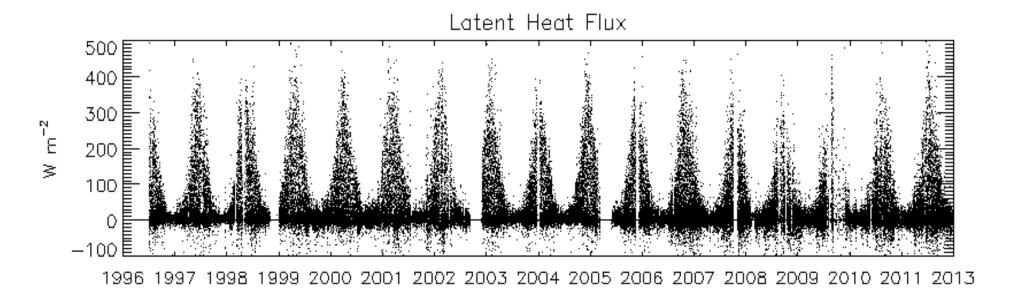


Desai, A.R., 2014. Influence and predictive capacity of climate anomalies on daily to decadal extremes in canopy photosynthesis. Photosynthesis Research, 119, 31-47, doi:10.1007/s11120-013-9925-z.



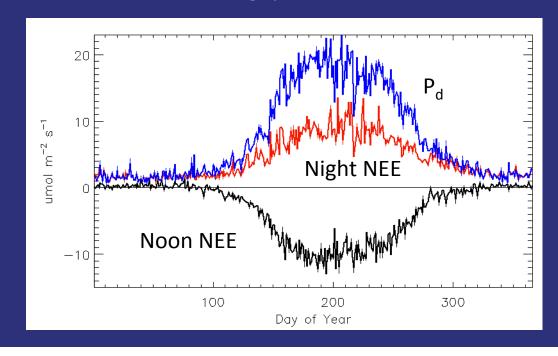






From NEE to Productivity

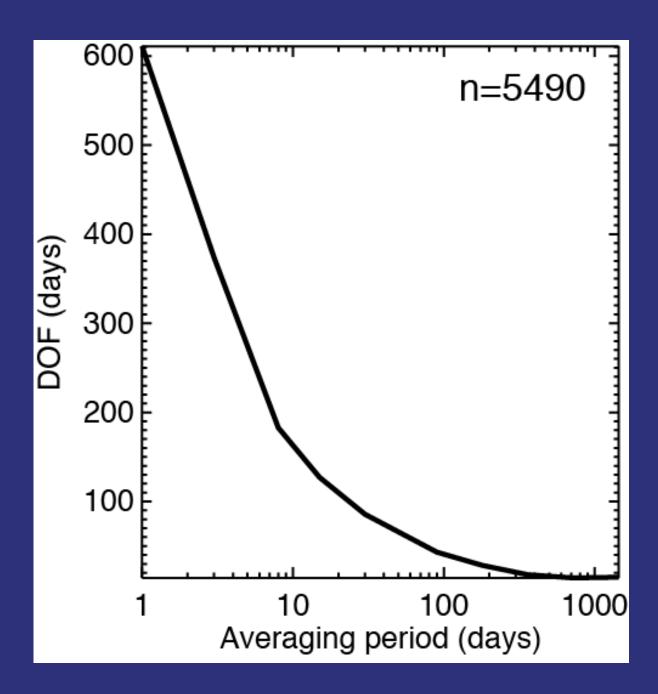
- Flux tower derived "GPP" is sensitive to model selection and gaps (Desai et al., 2008)
- INSTEAD: Use a data-based approach
 - P_d = Max nighttime observed NEE Mean noon (10-14)
 NEE
 - Reject noon NEE is > 50% gap-filled



Problem

- Every flux tower based correlation is significant when you have thousands to tens of thousands of datapoints
 - Effect sizes may be small, though
- Account for autocorrelation using "reduced degrees of freedom" metric!

$$N_* = \frac{N}{\sum_{t=N/2}^{N} \left[\left(1 - \frac{t}{N} \right) \rho_t^X \rho_t^Y \right]}$$

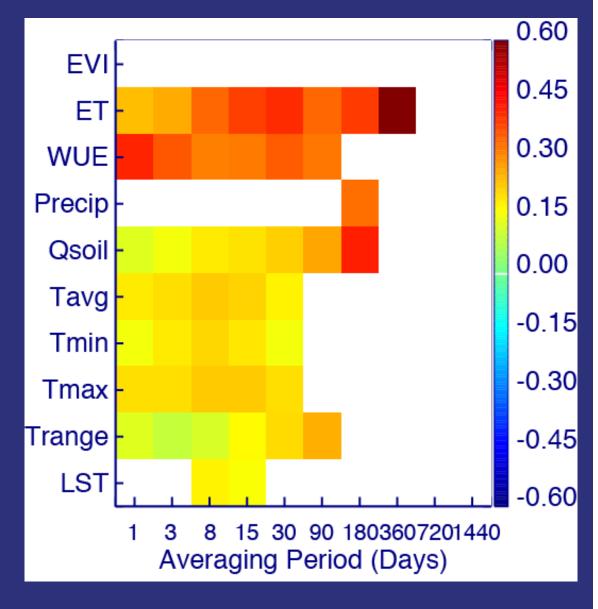


What to test?

Abbreviation	Description	Source
P_d	Photosynthetic drawdown	Flux tower
EVI	Enhanced Vegetation Index, 8-day average	MODIS TERRA/AQUA
ET	Evapotranspiration	Flux tower
WUE	Water Use Efficiency (P_d/ET)	Flux tower
Precip	Daily precpitation	NCDC + NARR
		Reanalysis
Q_{soil}	10 cm soil moisture	NARR Reanalysis
T_{mean}	Daily temperature	Flux tower + NCDC
T_{min}	Minimum daily temperature	Flux tower + NCDC
T_{max}	Maximum daily temperature	Flux tower + NCDC
T_{range}	Daily temperature range (max - min)	Flux tower + NCDC
LST	Land Surface Temperature, 8-day day/night	MODIS TERRA/AQUA
	average	

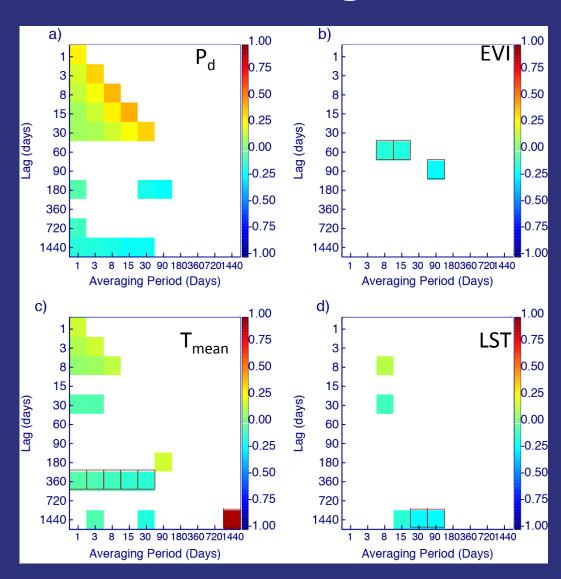
What do you get?

- Only significant correlations
 shown
- Moisture and temperature anomalies positively correlate with P_d at subannual scales



Lags are interesting

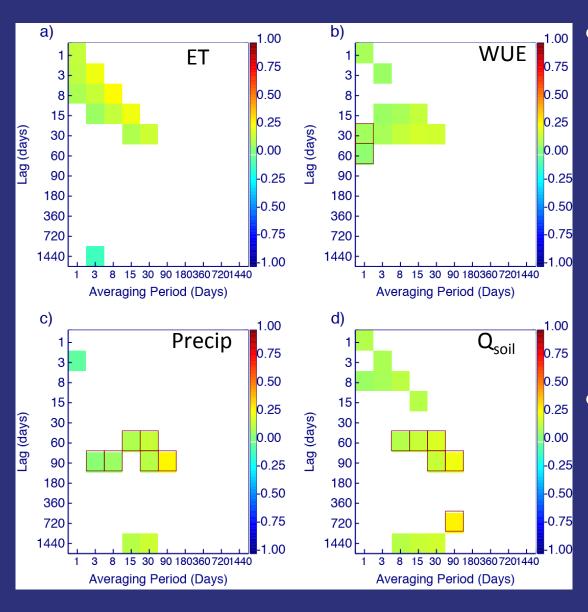
- Red squares = correlations > autocorrelation
- Remotely sensed variables (EVI,LST) have limited ability to predict P_d
- Previous year
 weekly-monthly
 temperature has a
 weak negative
 relationship to P_d



Important points 1

- Highly significant autocorrelations at daily to seasonal scales up to one month lag imply a strong biological feedback that can damp response to extremes
- Weak negative autocorrelations at multi-year scales also highlight slow press processes and oscillations
- Remotely sensed anomalies have little correlation to carbon flux even though mean seasonal variation correlates highly

Moisture lags are even more interesting

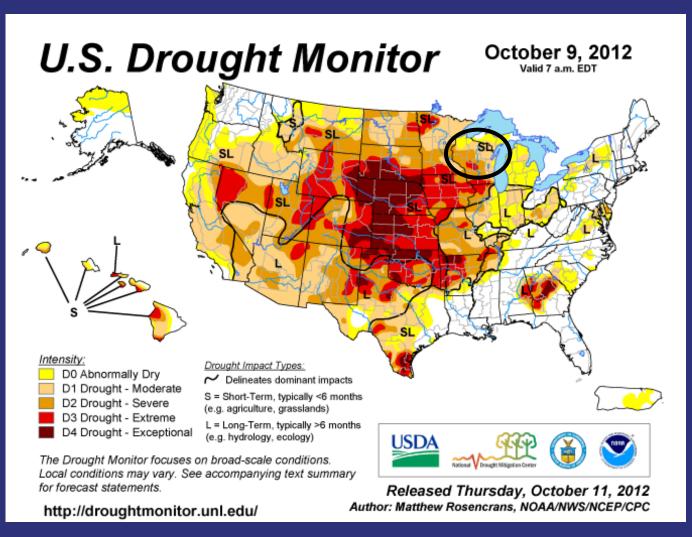


- Earlier season (2-3 month) weekly-seasonal precipitation/soil moisture has strongest predictive effect on P_d
- Beyond that, P_d autocorrelation dominates

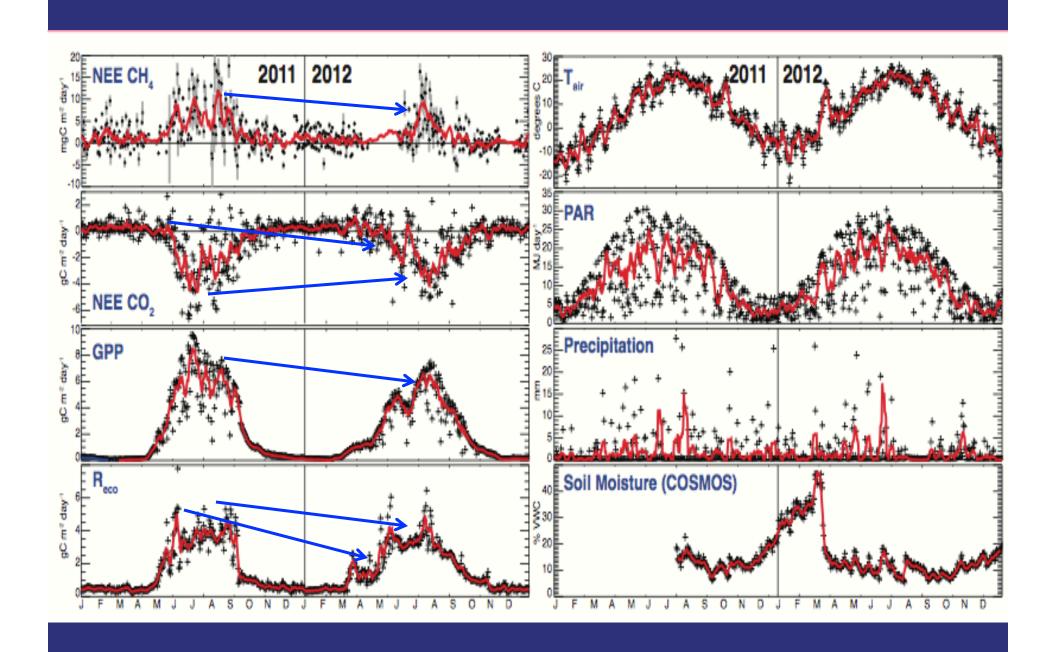
Important points 2

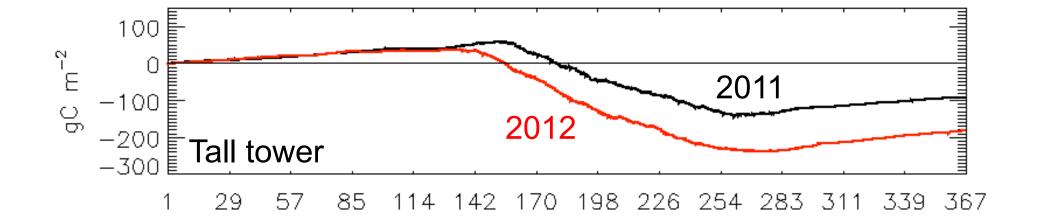
- Moisture extremes impact to regional carbon sequestration display significant seasonal lags and primarily influence monthly to seasonal uptake
- Positive correlations imply mesic forest is in-fact moisture limited, but not in the usual sense

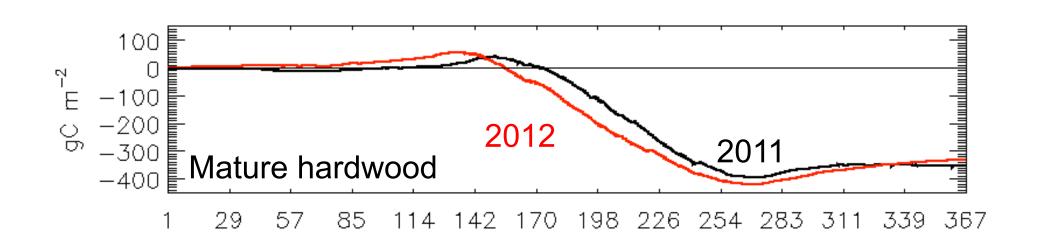
What about 2012?

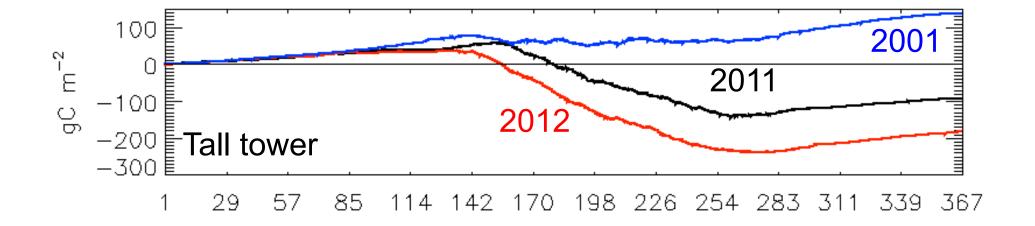


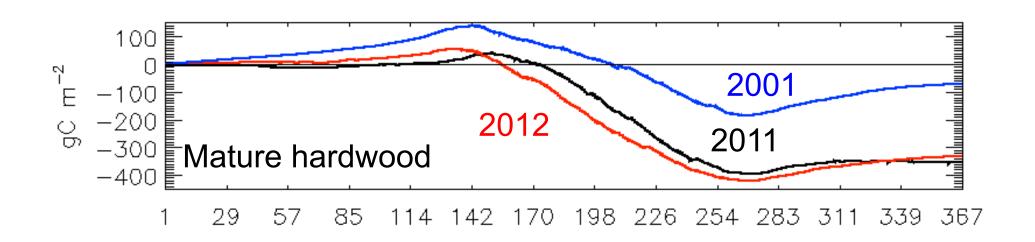
Wolf et al., in prep; Desai et al., in prep













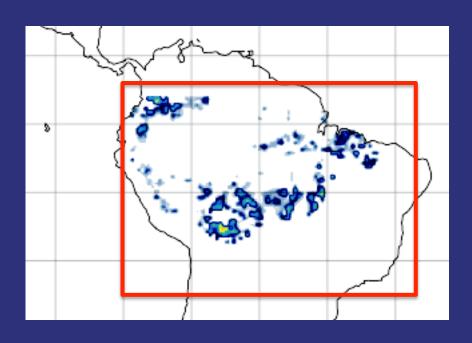
Important points 3

- Warm, dry conditions more likely promoted a longer growing season through phenology than reduced uptake by stomatal closure
- Biotic disturbances and their frequency/ extremes may be more important than climate extremes in many places





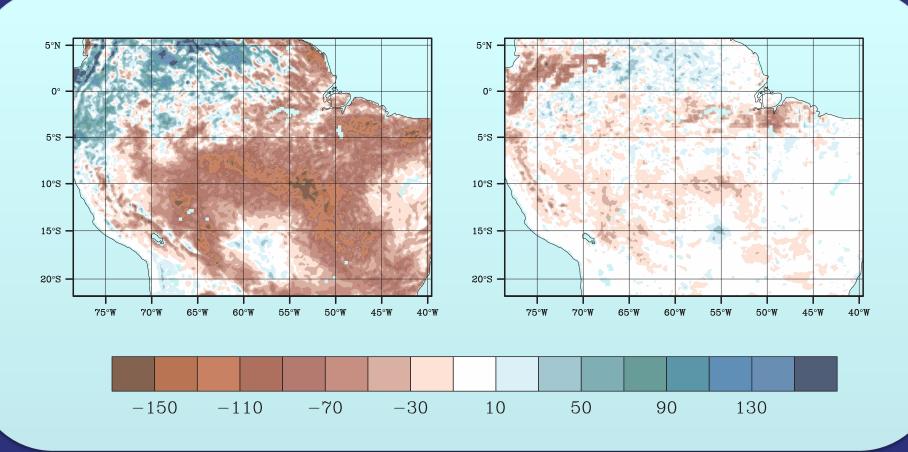
WRF-Noah Setup



Bagley, J.E., Desai, A.R., Harding, K.J., Snyder, P.K., and Foley, J.A., 2014. Drought and deforestation: Has land cover change influenced recent precipitation extremes in the Amazon? *J. Climate*, 27, 345-361, doi:10.1175/JCLI-D-12-00369.1.

- Spatial Resolution: 20km x
 20km
- Timestep: 60 seconds
- •For 2003, 2004, 2005, 2007, 2009, and 2010 the model was run from March 15 October 15 with and without deforestation
- •Total of 12 seven-month simulations completed with hourly output

Precipitation Rate (mm/month)



Dry Season Anomaly Deforestation perturbation

Amazon Rainforest Percent Changes with Deforestation

In nearly every measure the impact of deforestation is greater during drought years

	Pluvial Years
% Δ Precipitation Rate	-4.99%
% Δ Sensible Heat Flux	+.48%
% Δ Latent Heat Flux	-3.63%
% Δ Net Surface Radiation	-2.41%
% Δ Boundary Layer Height	11%
% Δ Rel. Soil Moisture Top	-3.00%
Layer % Δ Rel. Soil	
Moisture Bot. Layer	+3.50%
% Δ 2m Specific Humidity	77%
% Δ Level of free convection	+2.62%
$\% \Delta$ Lifting condensation level	+1.29%

July - September

Drought Years

-5.93%

+4.28%

-5.57%

-2.70%

+1.36%

-4.38%

+5.09%

-1.31%

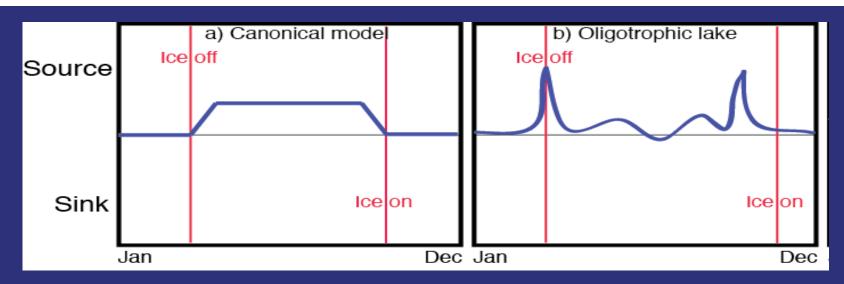
+.52%

+3.94%



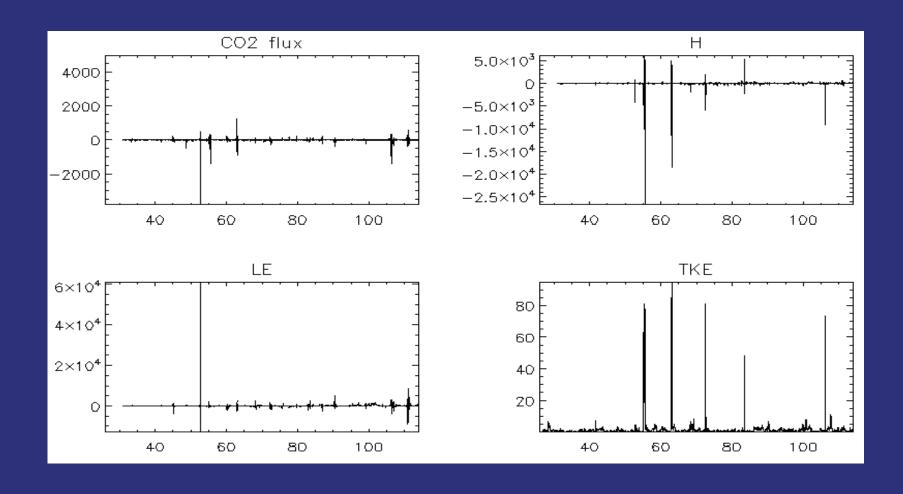
In the future...

- As models get more sophisticated and realistic, a greater number of negative (restoring) feedbacks will be successfully resolved
- However, this does not negate the very real risk of climate change on thresholds, longterm shifts, and other ecosystem state changes regardless of the feedback direction
- Further, some systems may be more sensitive than others

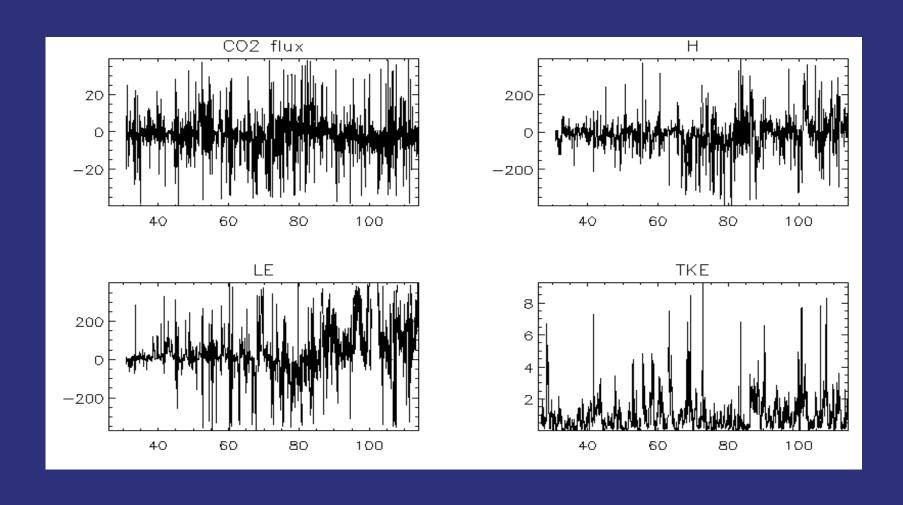




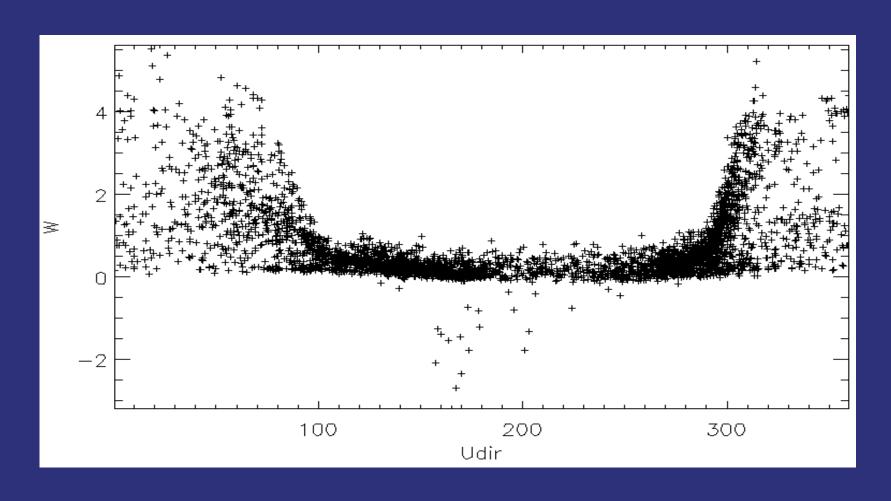
Great data?

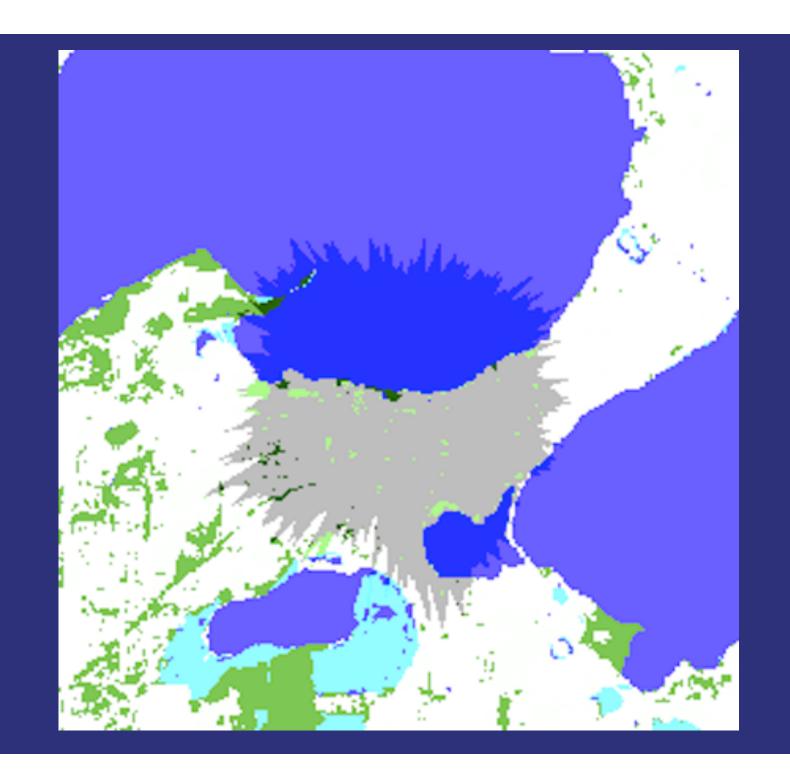


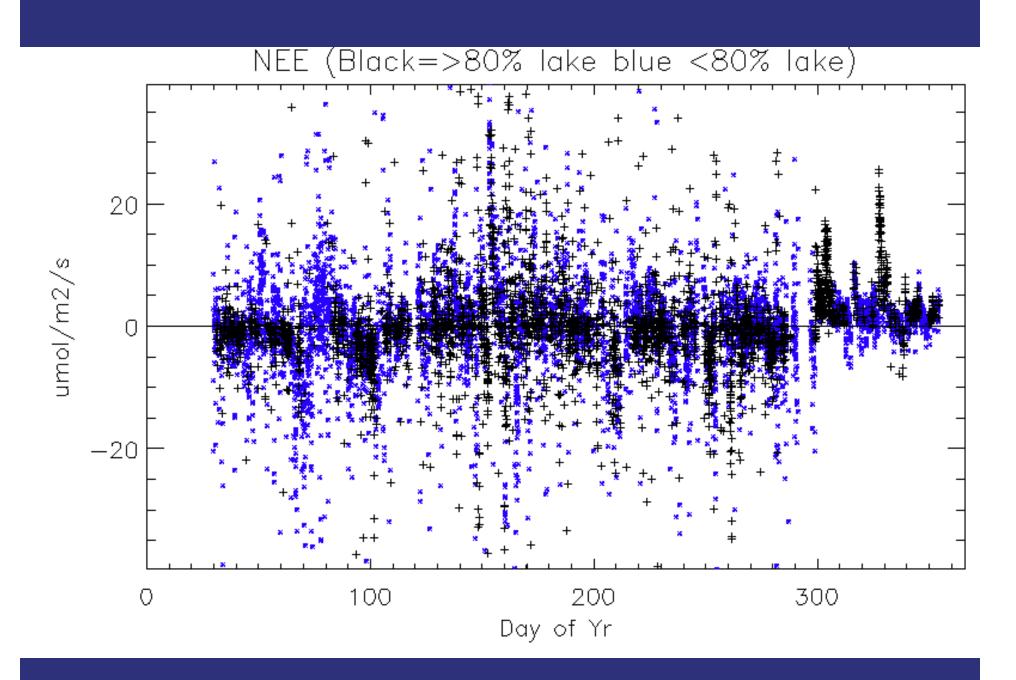
Oh – that's better

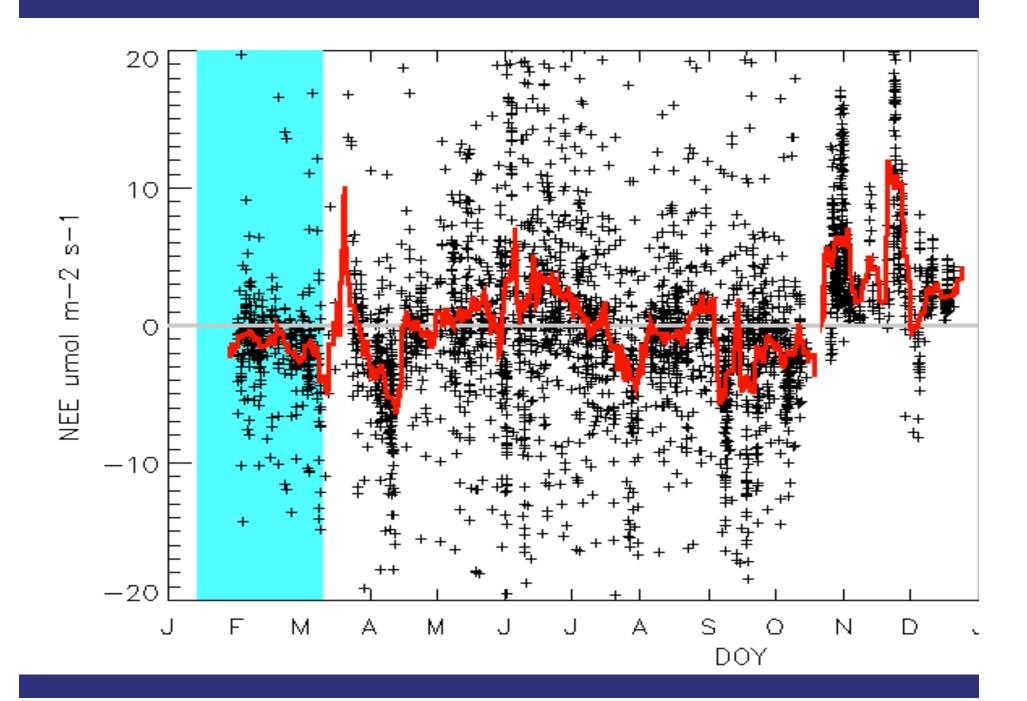


About that building flow













Thanks!

- Contributors:
 - Jonathan Thom, Ke Xu, Arlyn Andrews, Dan Baumann, Bruce Cook, Dave Moore, Britt Stephens, Justin Bagley, Ben Sulman, Malgorzata Golub, Mike Dietze, many others...
- Funding:
 - NSF, NOAA, DOE, USDA, WI Focus on Energy

